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Coastal Storm Model

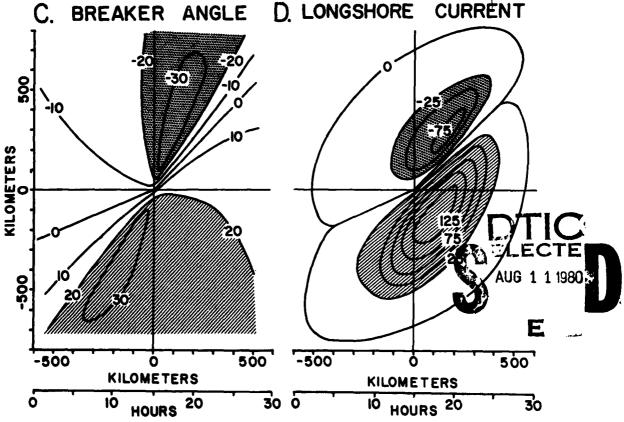
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by

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COASTAL STORM MODEL

by

William T. Fox

and

Richard A. Davis, Jr.

Technical Report No. 14, April 30, 1976

of
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Office of Naval Research



Williams College
Williamstown, Massachusetts

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ABSTRACT

A mathematical simulation model of a coastal storm has been programmed to forecast or hindcast wave and longshore current conditions at a coastal site. Storm parameters for the model are based on the size, shape intensity and path of the storm as derived from weather maps. An elliptical form is used to model the size and shape of the storm which are controlled by varying the length and orientation of the major and minor axes. Storm intensity is a function of the barometric pressure gradient which is modeled by an inverted normal curve through the storm center. The storm path is based on actual storm positions for the hindcast mode, and on projected positions assuming constant speed and direction for the forecast mode. The location, shoreline orientation and nearshore bottom slope provide input data for each coastal site.

For each storm position, the geostrophic wind speed and direction are computed at the shore site as a function of barometric pressure gradient and latitude. The geostrophic wind is converted into surface wind speed and direction by applying corrections for frictional effects over land and sea. The surface wind speed, fetch and duration are used to compute the wave period, breaker height and breaker angle at the shore site. The longshore current velocity is computed as a function of wave period, breaker height and angle and nearshore bottom slope.

The model was tested by comparing hindcast output with observed data for several coastal locations. Forecasts were made for actual storms and for hypothetical circular and elliptical shaped storms.

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Thomas Getz of Williams College developed some of the subcoutines and provided significant input in the theory and application of the computer simulation model. Douglas Rosen of the University of South Florida assisted in drafting several of the illustrations. Annie Laliberte typed and proofread the manuscript.

COASTAL STORM MODEL

INTRODUCTION

Coastal storms which provide a combination of high winds, pounding waves and rapid longshore currents are a major cause of distructive erosion along beaches and cliffs. Beaches which are generally composed of sand or cobbles are subject to sudden changes during storms. During one storm at Chesil Bank near Abbotsbury, England, the crest of a shingle beach was cut back 1.53 meters in 3 hours (Lewis, 1931). During a severe storm in July 1969 at Stevensville, Michigan, the beach and bluff were eroded back over 5.5 meters when the waves reached a height of 2 meters (Fox and Davis, 1970b). On the Oregon coast, a beach was stripped of a 2 meter thick blanket of sand and the wave cut terrace was exposed when the waves reached heights of 8 meters during late Novemeber storms (Fox and Davis, 1974).

For any operation involving the coastal zone, it is essential to make predictions of wave and current conditions during a coastal storm. General wave forecasts on a worldwide grid are available from the National Weather Service and Fleet Numerical Weather Center. These forecasts provide accurate predictions of wave conditions on the open ocean, but do not provide detailed enough predictions for the coastal zone. Therefore, a computer simulation model was developed to fill the gap between wave predictions on the open ocean and surf predictions along the coast.

The coastal storm model utilizes an ellipse to simulate a map of barometric pressure. The shape of the storm can be modified by varying the length and orientation of the major and minor axes of the ellipse. The intensity of the storm is controlled by increasing or decreasing the range in barometric pressure. The actual storm track and dimensions are read in as data for hindcasting wave and current conditions. For making forecasts, the size, shape and intensity of the storm are provided as input data for the model. In forecasting, the storm path is plotted by assuming a constant azimuth and velocity.

One of the initial steps in developing a coastal storm model is determination of the barometric pressure gradient at any point on the ground surface under the storm. The pressure gradient is then used in conjunction with the latitude to calculate the geostrophic wind speed, which in turn is used to compute the surface wind speed, wave height and longshore current velocity. In the model, it is assumed that a profile of barometric pressure along the major or minor axis of the storm ellipse can be represented by an inverted normal curve. By rotating the normal curve around the ellipse and taking the derivative, it is possible to calculate the pressure gradient at any point on the ground. From that point on, conventional methods are employed for determining the geostrophic wind speed, surface wind speed, wave height and longshore current velocity.

PREVIOUS WORK

The coastal storm model is based on a series of detailed field studies which extended from 1969 through 1975. The studies included the analysis and synthesis time-series data on barometric pressure, wind speed and direction, wave period and height and longshore current velocity for 15 days to 1 year. Topographic profiles across the beach and nearshore area were used to construct topographic maps and maps of erosion and deposition. The sites in the study are plotted by date in Figure 1 and included in the following list along with references for each study.

1969	Stevensville, Michigan	Fox and Davis, 1970a, 1970b
1970	Holland, Michigan	Fox and Davis, 1971a Davis and Fox, 1971
1971-2	Mustang Island, Texas	Davis and Fox, 1972c Davis and Fox, 1975
1972	Sheboygan, Wisconsin	Fox and Davis, 1972
1973	Cedar Island, Virginia	Davis and Fox, 1974a
1973-4	South Beach, Oregon	Fox and Davis, 1974
1974	Zion, Illinois	Davis and Fox, 1974b
1974	South Haven, Michigan	Davis and Fox, 1974b
1975	Plum Island, Massachusetts	

The models have evolved from a geometric model called the area-time prism (Davis and Fox, 1972a) through a conceptual model (Davis and Fox, 1972b) to a process-response model for Lake Michigan (Fox and Davis, 1971b and 1973). The simulation model developed for Lake Michigan was limited to the local geographic area where the storms moved directly onshore to the north of the study area. The proposed coastal storm model is an outgrowth of the earlier model but is more generalized with broader application under a wider range of storm conditions and shoreline orientations.

Several different types of computer models have been proposed for the coastal environment. Probablistic models were developed to reproduce gross coastal features such as a recurved spit on the south coast of England (McCallagh and King, 1970) and the Mississippi River Delta (McCammon, 1971). A markey process was used to simulate the sequence of bar formation and migration across a beach

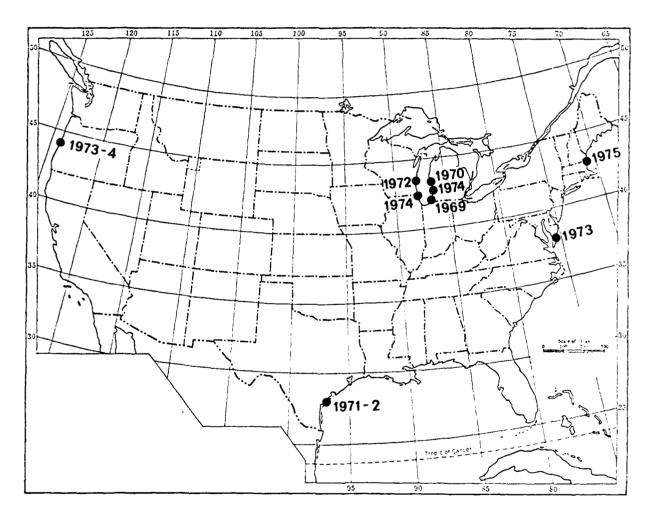


Figure 1. Location map of field sites designated by project year.

(Sonu and van Beek, 1971). A deterministic model resembling a wave tank experiment was proposed to simulate the interaction between a prograding delta and waves (Komar, 1973). A statistical model with a relatively simple beach topography to compute breaker height, longshore current velocity and wave setup (Collins, 1971) was followed by a more deterministic approach to model the nearshore circulation patterns employing monochromatic waves and more complex beach topographies (Noda, et al, 1974). An explicit finite difference model for predicting time-dependent, wave induced nearshore circulation was developed by Birkemeier and Dalrymple (1976).

On a larger scale, Resio and Hayden (1973) proposed an integrated storm model which combines three scales of atmospheric motion, large scale, synoptic scale and small scale into an estimation of a winter wave-surge climate for the mid-Atlantic coast. At a similar scale, Goldsmith, et al (1974) developed a wave climate model for the mid-Atlantic coast by using Dobson's (1967) wave refraction program to project offshore waves into the coastal zone.

The coastal storm model proposed in this report provides a link between the large-scale, seasonal wave-climate models and the dynamic surf zone models. By tracking a storm across a shoreline, the wave parameters which are output from the storm model furnish input for the surf zone models. Therefore, the proposed storm model could be combined with other computer models to provide an integrated process model for the coastal zone.

COMPUTER PROGRAMS - COASTAL STORM MODEL

Program STORM

Program STORM is a mathematical simulation model which has been programmed for the computer to forecast or hindcast wave conditions at a coastal site during a storm. The actual storm as represented by the isobars on a weather map is modeled by an elliptical storm with major and minor axes at right angles and passing through the center of the low. The size, shape, intensity and path of the storm as determined from weather maps are used to generate the surface wind pattern, wave height and period, and longshore current velocity as the storm moves across the coast.

The computer program is divided into a main program, STORM, and a series of 11 subroutines. The main program is used to read in the data, call the various subroutines for computing the wind, wave and current conditions, print out the predictions at one hour intervals. All the input and output is handled by the main program while the calculations are carried out by the various subroutines. In this way, any portion of the model can be independently tested by using a small main program to call each subroutine individually. Therefore, if any problem arises in the main program, it can be narrowed down to a particular subroutine, and that subroutine can be tested under a variety of conditions without using the main program. Also, if a portion of the simulation model is to be incorporated into another program, any of the individual subroutines can be removed and used separately with the appropriate calling arguments.

The theory and mechanics of the program will be explained in detail starting with the MAIN program and proceeding through each of the subroutines as they are called by the MAIN program. The program was written in FORTRAN IV for an 8k IBM 1130 at Williams College. A full listing of the programs with appropriate comment cards is included in Appendix A. A second version of the MAIN program was written for the Xerox 530 which includes a graphics package for a 29 inch plotter. The graphics package is used to plot barometric pressure, surface wind, onshore wind, alongshore wind, breaker height, wave period and longshore current.

Main Program - Input and Output Options

The main program is used to read the input data for the storm and shoreline conditions, call the various subroutines and print out the results at one hour intervals. A listing of the input cards is included in Appendix I with a description of each of the input variables. The program is dimensioned to make predictions up to 130 hours or 5 days and 10 hours. If a longer prediction is desired, it is necessary to increase the dimension of U(130) and V(130) to the required number of hours. In the model, 130 hours was selected because of core limitations on the 8k IBM 1130. For most of the storms, the 130 hour limitation was not a serious constriction, however, a larger dimension statement would be helpful in some cases.

The first two data cards are used to read in the title, starting time, date and input/output options. The title used for the location of the shore site can be up to 80 spaces long filling card 1. On the second card, the starting hour, ISTRT, is included in columns l and 2 followed by the date, DAY, in columns 3 to 22. The starting time is read in as an interger and must be right justified. If ISTRT is read in as 0, the program will terminate. For the input option, INAUT, in column 23, a 0 is used for metric units and a 1 is used for nautical units including nautical miles, knots, and feet per second. The output option, NAŪT, in column 24 is separate from the input option but uses the same code, O for metric units and I for nautical units. Although metric units are becoming the standard and are now required for scientific reports, it may be desirable at times to have the input or output in nautical units. With separate input and output options, it is possible to have the input in metric or nautical units and convert to the other units with the output.

The first 3 columns of card 3 are used to select the major options for the program. For the first option, INOPT, a 1 in column I will call the hindcasting mode, while a 2 in column I will call the forecasting mode. The hindcasting mode is used when storm positions are available at six hour intervals. The hourly positions of the storm are determined by a linear interpolation between the 6-hour positions. For the forecasting mode, the initial position of the storm along with a constant velocity and azimuth are used to calculate successive positions at 1-hour intervals. The variables for the hindcasting and forecasting modes are read in on card 6. The second option on card 3 is the tide prediction option, IFTID. If a 0 is punched in column 2, the tide prediction option is suppressed and card 4 is not included in the data set. The tide option is omitted for a non-tidal body of water, such as the Great Lakes. Where tide data are available from the tide tables or from observations, a 1 is punched in column 2, and the tide data are included on card 4. The longshore current equation for the simulation run is selected in option 3, LSCOP. Four different longshore current equations are included, (1) Fox and Davis (1972), (2) Longuet-Higgins (1970), (3)

Coastal Engineering Research Center (1973), and (4) Komar and Inman (1970). The longshore current equations are called in subroutines SURF and their differences will be discussed under that subroutine.

The number of storm positions, NX, are punched in columns 4 to 6 of card 3, for 6-hour intervals in the hindcasting mode, and for 1-hour intervals in the forecasting mode. For example, if the hindcasting mode is used, and 3 days of data are included, the initial position and 4 positions for each day would give a total of 13 for NX. For the forecasting mode, a 3 day forecast would use a 73 for NX, 1 for the initial position, and 72 for the 72 hour forecast. The maximum value for NX is 22 for the hindcast mode and 130 for the forecast mode.

The average basin fetch in kilometers, BNFCH, is punched in columns 7 to 12 on card 3. The average basin fetch is used as the limiting fetch in determining the wave height and period from the wind speed. Where the basin fetch is smaller than the maximum storm axis, the waves are fetch limited. However, where the fetch is significantly larger than the storm size, the average basin fetch will not be a limiting factor in determining the wave parameters. The basin fetch is considered in an offshore direction from the shore site. In the case of a large ocean, the approximate width of the ocean can be used as the basin fetch.

In columns 13 to 17 of card 3, the time interval between storm positions TINT is normally set at 1.0. The time interval refers to the printout spacing for the forecast modes. For the hindcast mode, the values are read in at 6-hour intervals, and the results are printed out at 1-hour intervals.

The minimum barometric pressure in millibars, PMIN, taken at the center of the low pressure cell is punched in columns 18 to 24 on card 3. Usually, the minimum barometric pressure is interpolated within the smallest isobar. Thus, if the smallest isobar is 1004, and the isobar spacing is 4 millibars, the minimum pressure would be estimated at 1002 millibars. The pressure at the largest encircling isobar, PMAXR, is used to determine the intensity of the storm. If the storm is circular or oval shaped, the largest isobar which encloses the storm center is used for PMAXR and punched in columns 25 to 31 of card 3. If, however, the storm has a wave form extending down from the north, a line is drawn along the storm path through the storm center to the margins of the storm. The largest isobar which the line crosses on both sides of the storm is then considered the largest encircling isobar, PMAXR. In the program, the largest encircling isobar is defined as 2 standard deviations away from the center of the storm. Therefore, the total storm radius would be 1.5 times the radius of the largest encircling isobar. The

pressure range would be 1.145 times the range within the largest encircling isobar. The latitude at the shore site, SLAT, punched in columns 32 to 36 is used in subroutine WIND to compute the geostrophic wind speed.

The geographic size of the storm is defined in terms of an ellipse with a major half-axis and minor half-axis corresponding to radius of a circle. The major half-axis, AR, of the storm ellipse measured when the storm is closest to the study site is punched in columns 37 to 42. If the storm ellipse is assymetrical, the longest half-axis on the side toward the shore location is used as the major half-axis. The minor half-axis, BR, is measured at right angles to the major half-axis through the center of the low. The major and minor half-axes are measured from the center of the low to the largest encircling isobar PMAXR. The minor half-axis, BR, is punched in columns 43 to 48 on card 3.

The orientation of the major half axis EAZ is punched in columns 49 to 54 of card 3. The orientation of the major axis is measured in degrees from true north to the northern end of the major axis ranging from -90° on the west to 90° on the east. For a front or trough related to a low pressure system, the major axis is usually several times longer than the minor axis and the orientation of the major axis lies along the line of the front. In a circular or oval storm, the major axis is usually 1 to 1.5 times as long as the minor axis. The major and minor half axis are measured when the center of the storm is at its nearest position to the shore study site.

Variables for hourly tide prediction are contained on card 4. The spring tide range in meters, ST, is punched in columns 0 to 5 and the neap tide range, TN, in columns 6 to 10. The spring and neap tide ranges and the time of the last spring high tide are available in the tide tables which are published annually by N.O.A.A. In making the hourly predictions for the model, it is necessary to punch the number of days since the last spring tide, TDAY, in columns 11 to 15. The hour of the last spring high tide, THR, preceding the start of the run is punched in columns 16 to 20. The tidal form number, FN, punched in columns 12 to 25 is used to reproduce a semidiurnal, mixed-semidiurnal, mixed-diurnal or diurnal tide with the right spacing and tidal beat. The nearshore bottom slope at low tide, SLPLO, is punched in columns 26 to 32 and the slope at high tide, SLPHI, is punched in columns 33 to 39. The nearshore bottom slope which varies with tidal elevation is used for computing longshore current velocity in subroutine SURF. The preferred method of determining bottom slope is to fit a linear surface to the nearshore map at low tide, and repeat the process at high tide. The linear slope should extend to a depth of at least twice the breaker height. For the high tide range, the foreshore slope and low tide terrace should be included in the slope calculation. Where it is not possible to fit a linear surface because of lack of data, it is possible to get an approximation of the nearshore slope by measuring the depth at some predetermined distance from the shore at low tide and at high tide. By dividing the depth by the distance, a good approximation of nearhsore slope can be estimated for low and high tides. The nearshore slope is an initial factor for the determination of the longshore current velocity, so care should be taken in estimating nearshore slope at low and high tide. It also should be pointed out that nearshore bars have a significant influence on nearshore currents and must be considered in making an estimate of the nearshore slope. The final variable on the tide prediction card is the mean tide level, TMEAN, which is the difference between mean sea level and mean low tide as reported in the annual tide tables.

Data for the shore site location including geographic coordinates, onshore direction, average bottom slope and offshore island option are punched on card 5. The shore site location is given in a X and Y coordinate system where the X-axis runs eastwest with positive X in the east direction, and the Y-axis northsouth with positive Y in the north direction. The X-Y coordinate system is measured in kilometers with the origin located at the southwest of the study site. In practice, a piece of 10 to the inch rectangular grid graph paper is laid over the weather map with the Y axis parallel to the longitude line nearest the study site. The origin of the graph paper is placed several inches to the southwest of the shore location so that the X-axis runs eastwest and the Y-axis runs north-south parallel to the latitude and longitude lines through the study site. The X and Y coordinates are read off the map in inches and converted to kilometers before they are used in the program. The X coordinate, ULOC, is punched in columns 1 to 7 and the Y coordinate, VLOC, is punched in columns 8 to 14 on card 5.

The orientation of the shoreline given by the onshore azimuth, SHAZ, and the average nearshore bottom slope, SLOPE, are punched on columns 15 to 21 and 22 to 38 respectively. The onshore direction measured in degrees in a clockwise direction from north is used to give the orientation of the shoreline. Since the storm is considered a regional feature, it is necessary to give the regional orientation for the shoreline. An east-west shoreline with land to the north would give a 0 azimuth. If a shoreline is running north-south with the land to the east and water to the west, the shoreline azimuth would be 90 degrees. Similarily, if the shoreline is running north-south with the land on the west and the water on the east, the onshore azimuth would be 270 degrees.

An option is available with the program for an offshore island which is not influenced by a large continental land mass. When a 0 is punched in column 30 of card 5, the offshore island option is suppressed and a normal continental coast or barrier island is assumed.

For a coast backed by land, land corrections are used in computing the surface wind speed when the wind blows offshore. Therefore, when the island option is used and a l is punched in column 30 of card 5, the wind is assumed to be blowing from the sea in all directions and the land correction is not used. For a barrier island which lies roughly parallel to the coast, the island option is not used because an offshore wind blows over land for a long distance before it hits the lagoon and barrier island.

The storm positions for the hindcasting and forecasting modes are punched on card 6. For option 1, the hindcasting mode, the X and Y coordinates are punched in columns 1 to 7 and 8 to 14 respectively. The X and Y coordinates are from the rectangular grid discussed for the shore site location on card 5. The coordinates for the storm are given for the initial storm position and at successive 6 hour intervals, with one pair of coordinates per card. The number of pairs of coordinates is specified by NX, the number of storm positions on card 3. For the forecasting mode, option 2, the storm positions are determined at 1 hour intervals from the storm velocity, storm azimuth and initial X and Y coordinates. The storm velocity, SVEL, in columns 1 to 7, is given in kilometers per hour. Reasonable storm velocities would vary from about 25 to 75 kilometers per hour for a slow to fast moving storm. In the forecasting mode, it is not possible to vary the storm velocity, so the initial storm velocity must be maintained for the entire forecast run. The storm azimuth or path is measured in degrees clockwise from north. As with the storm velocity, it is not possible to vary the storm azimuth in the middle of a forecast run. An azimuth of O degrees would have the storm moving due north, and a 90 degree azimuth would have the storm heading east. In the forecast mode, the initial X and Y coordinates for the storm are punched in columns 15 to 21 and 22 to 28 respectively. It is possible to make a map for each predicted variable by making a series of runs with different initial coordinates.

It is possible to run a series of models for different coastal situations by including a new data set for each model starting with the title card. To terminate the run, two blank cards are included at the back of the data deck. Since the second card of the new data set is blank, ISTRT is read in as 0 and the program will finish.

Different versions of the main program are used for making forecasts directly from the console, and for printing a map using the forecast mode. Listing of the programs are included in Appendix A along with explanations of the input options.

Subroutine LOCAT

As a storm moves across a coastline, subroutine LOCAT is used to determine the position of the coastal site relative to the storm center for each increment of time. In preparing the input data for the program, the location of the coastal site and a sequence of storm positions are plotted on a rectangular grid referred to as the map coordinate system. For the map coordinate system, the X axis points east, the Y points heading north and the origin is located to the southwest of the initial storm position. When the shore location and storm positions are plotted on a weather map, the coordinates are measured in kilometers or nautical miles, whichever are the most convenient units for the project.

In computing the geostrophic wind speed, it is necessary to determine the gradient in barometric pressure at the coastal site. Therefore, a storm coordinate system is established with the origin at the center of the storm, the XI axis parallel to the shore, the YI axis perpendicular to the shore, and the positive YI direction heading onshore. When facing the land from the sea, the positive XI direction is along the shore to the right, and the negative XI direction is to the left (Figure 2). Each time the storm moves, the origin of the storm coordinate system is also moved to the new location for the center of the storm. However, the orientation of the XI and YI axes remains the same with the XI axis parallel to the shore and the YI axis at right angles to the coast.

In the storm model, the units for the Xl, Yl coordinate system are converted from kilometers or nautical miles to storm radii by dividing by the radius of the storm. In an elliptical storm, the length of the major half axis is used in place of the storm radius.

In subroutine LOCAT, ULOC and VLOC are the map coordinates for the coastal site, and UST and VST are the map coordinates for the storm center (Figure 2). Vectors are computed parallel to the X axis (U = UST-ULOC) and parallel to the Y axis (V = VLOC-VST). The resultant vector ($Z^2 = U^2 + V^2$) gives the map distance from the storm center to the coastal site. The counterclockwise angle (ANG) between the positive X axis and the Z vector is computed by the arctangent subroutine ARCTA. The onshore azimuth (SHAZ) is the onshore direction normal to the shoreline measured in a clockwise direction from north. The angle A (A = ANG - SHAZ) is used for converting coordinates from the map system to the storm system. The shore position is then determined in the storm coordinate system for distances along the X1 axis (X1 = -Z * cos (A)) and along the Y1 axis (Y1 = Z * sin (A)).

A third coordinate system is set up for dealing with an elliptical storm. In the elliptical coordinate system, the P axis lies along the major half-axis and the Q axis lies along the minor half-

axis. The distances, P and Q, within the storm ellipse are used to locate the shore site relative to the center of the ellipse. The orientation of the storm ellipse is given by the ellipse azimuth (EAZ) which is the azimuth of the major half-axis plus or minus 90 degrees from true north.

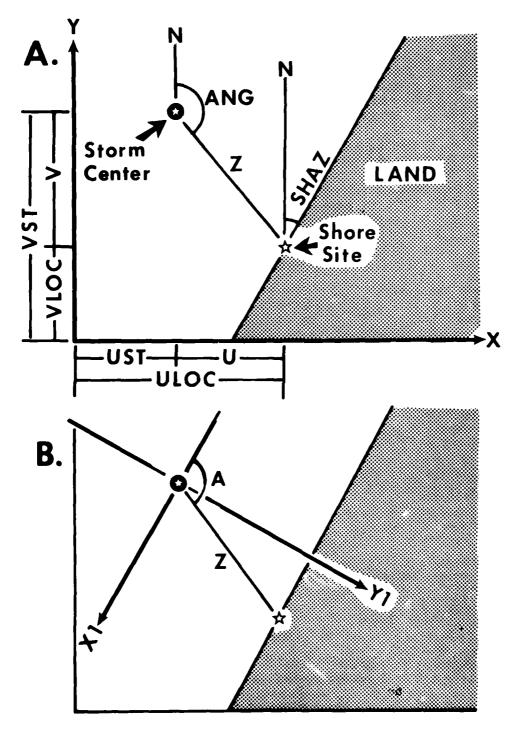


Figure 2. A. Map coordinate system (X-Y) for locating storm center and shore site, and B. Storm coordinate system (X1-Y1) with origin at storm center and X axis parallel to the shore.

Subroutine ELIPS

The wind angle and barometric pressure gradient at any position within an elliptical storm are determined by subroutine ELIPS. A is the length of the major half axis, B is the length of the minor half axis and EAZ is the angle from true north to the path of the major half axis. The shore site is located at point X1, Y1 within the storm ellipse. To determine the pressure gradient, a second ellipse is plotted with axes Al and Bl which passes through point X1, Y1. The second ellipse has the same axial ratio B/A and the same origin as the storm ellipse (Figure 3). The tangent to an ellipse is defined by equation 1.

$$\frac{X_1 X}{a^2} + \frac{Y_1 Y}{b^2} = 1 \tag{1}$$

therefore, the intersection of the tangent to the ellipse with the X axis can be found by equation 2.

$$X = \frac{a^2}{X_1} - \frac{a^2 Y_1 Y}{b^2 X_1}$$
 (2)

Equation 3 is used to generate the line normal to the tangent

$$X = X_0 + \frac{b^2 X_1}{a^2 Y_1} \quad Y.$$
 (3)

The intersection of the line normal to the tangent and the A axis is found by equation 4.

$$X_0 = X - \frac{b^2 - X_1}{a^2 - Y_1} - Y$$
 (4)

which is terms of the point X1, Y1 would be

$$X_0 = X_1 - \frac{b^2}{a} Y_1$$
 (5)

The distance X_0 is measured from the center of the ellipse to the point where the line normal to the tangent through Xl, Yl intersects the A axis. The point X2, Y2 is the intersection of the line normal to the tangent at Xl,Yl and the path of the ellipse.

To determine the gradient of the barometric pressure, it is assumed that the pressure gradient follows a normal curve along

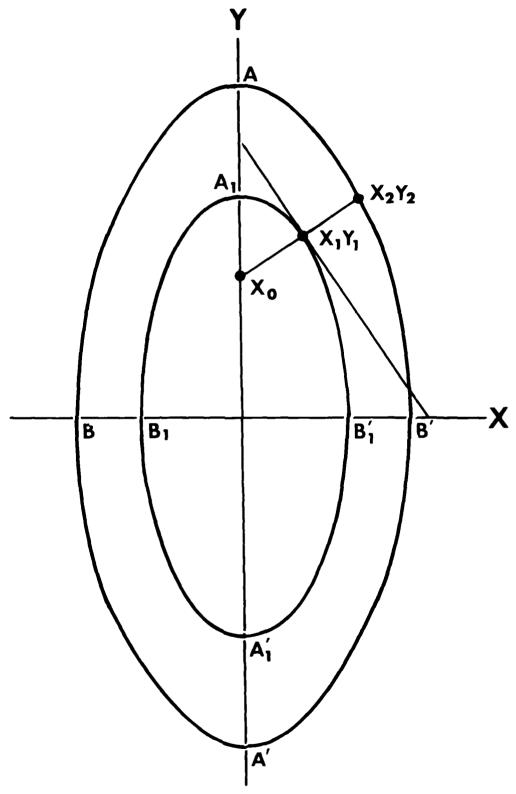


Figure 3. Location of the shore site (X1,Y1) within a storm ellipse (AB) and on a minor ellipse (A1,B1).

the major axis of the ellipse. Pl is the barometric pressure at point X_0 along the A axis. The pressure at X_0 is used in determining the pressure gradient normal to the isobar at point X1,Yl. The final pressure gradient calculation is made in subroutine WIND. XA and YA are used to plot the tangent to the ellipse at X1,Yl for determining the wind direction. The wind direction is assumed to be parallel to the tangent to the ellipse at X1,Yl and in a counterclockwise direction around the center of the ellipse.

Subroutine WIND

The geostrophic wind speed and direction for each storm position are computed in subroutine WIND. The equation for geostrophic wind speed V $_{\rm g}$ is based on latitude and barometric pressure gradient.

$$V_{g} = \frac{S}{2 \Omega \sin \phi} \frac{\Delta P}{\Delta N}$$
 (6)

where S is the specific volume, (779 cm 3 /gm), Ω is the angular velocity, (7.29 x 10^{-5} rad./sec.), ϕ is the latitude in degrees, and $\Delta P/\Delta N$ is the barometric pressure gradient normal to the isobar at the shore location (Godske et al, 1957, p. 370). The barometric pressure gradient is computed at right angles to the tangent of the ellipse through the shore site, (point XI, YI). To compute the gradient, a normal curve is constructed perpendicular to the tangent through point X1, Y1. The derivative of the normal curve is taken at point X1, Y1 to compute the barometric pressure gradient. The geostrophic wind direction is assumed to be parallel to the tangent of the ellipse at point X1, Y1 and heading in a counter-clockwise direction around the center of the ellipse. It is assumed that the small ellipses within the storm ellipse are parallel to isobars. Therefore, wind direction can be determined if the geostrophic wind is directed along the isobars with the high pressure to the right and low pressure to the left of the motion in the northern hemisphere. By means of geostrophic wind equations, the wind direction can be estimated with error of less than 10°, and speed with an error of less than 20% (Cole, 1970, p. 185).

An approxiamte relationship exists between the speed and direction of the surface wind measured at anemometer level and the upper quasi-geostrophic wind. Owing to differences in surface roughness, this relationship varies from one station to another, and also varies at a single station with stability. Thus it is rather difficult to determine the surface wind accurately from the upper quasi-geostrophic wind. Under average conditions, a rough method may be applied which makes use of the horizontal friction force, SR, near the ground (Figure 4). The equation for horizontal motion can be used for computing the friction force at a given station (Godske, 1957, p. 453).

$$sR = v_h + sv_h P + 2\alpha_z k \times v_h$$
 (7)

The horizontal component of the wind speed, v_h and the barometric pressure gradient, sv_hp , can be measured directly and inserted into the equation. Although it would be possible to determine v_h , the vertical variation in wind speed, it would involve measurements in time and space which would be vary laborious and not to reliable. In practice, v_h includes both a convective term and a local term. However, the influence of v_h is small when mean values over relatively large areas are included (Hasselberg and Sverdrup, 1915). Computations of the average of R based on time-series averages of a series of synoptic series of maps was carried out over land (Baur and Phillips, 1938) and over sea (Westwater, 1943). Based on their computations, the frictional force SR is directed backward to the right of the wind v_h , and is proportional to the wind velocity, v_h .

$$sR = bv_h \tag{8}$$

and forms an angle β with $-v_h$ as shown in Figure 4. To make corrections for wind speed and direction over land and sea, the mean values of b and β are: b = 1.9 x 10^{-4} sec $^{-1}$, β = 29° over land and b = 0.65 x 10^{-4} sec $^{-1}$, β = 50° over sea (Baur and Phillips, 1938).

Using b and β , a simple diagram has been constructed to show the configuration of forces when $v_h=0$. The balance of forces in the diagram along v_h and normal to it gives the following equalities with the angle between the geostrophic wind and the surface wind denoted by α .

s |
$$\nabla p$$
 | $\sin \alpha = sR \cos \beta = bv_h \cos \beta$
s | ∇p | $\cos \alpha = sR \sin \beta + 2\Omega_z v_h$
= $bv_h \sin \beta + 2\Omega_z v_h$ (9)

therefore, by division

$$\cot \alpha = \tan \beta + \frac{2 \Omega \sin \phi}{b \cos \beta}$$
 (10)

where \flat is the latitude. Therefore, the angle α between geostrophic wind and the surface wind can be determined directly from the latitude φ when b and β are known for a given station.

The ratio between the surface wind v_h and the geostrophic wind v_g can be derived from equations 9 and 10 and are given below according

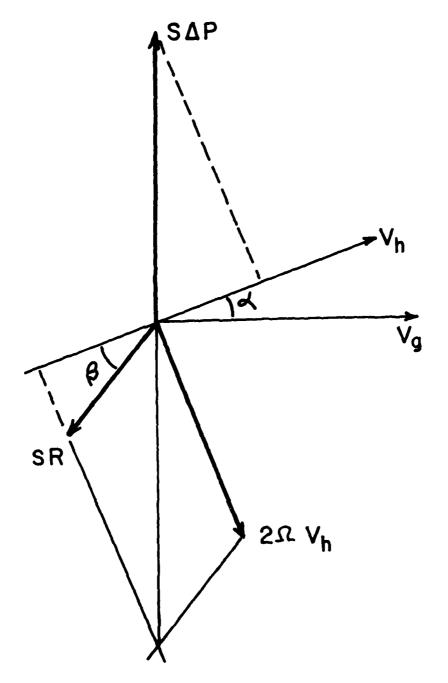


Figure 4. Orientation of the frictional force near the surface of the earth (Godske, et al, 1957, p. 453).

to Godske et al, 1957.

$$\frac{\mathbf{v}_{h}}{\mathbf{v}_{g}} = \frac{\mathbf{v}_{h}}{\frac{\mathbf{s} | \nabla p|}{2\Omega \sin \phi}} = \frac{\mathbf{s} | \nabla p| \sin \alpha 2\Omega \sin \phi}{\mathbf{b} \cos \beta \mathbf{s} | \nabla p|}$$

therefore,

$$\frac{v_h}{v_g} = \frac{2 \Omega \sin \phi \sin \alpha}{b \cos \beta}$$
 (11)

Once the angle α between the surface wind and geostrophic wind has been computed, it can be inserted into equation 11 to compute the ratio between the surface wind and the geostrophic wind over land or sea. Table 1 gives values at different latitudes for α , the angle between surface wind and geostrophic wind, and v_h/v_g , which are plotted in Figure 5. In subroutine WIND, the correction factors for computing surface wind speed and direction are computed following statement 50.

Since the values for b and β are given for wind blowing over land or over sea, intermediate values must be computed for winds blowing along the shore. Winds blowing directly onshore with a wind angle of zero would have values of b = .000065 and β = 50. In this case, the wind is blowing from the sea and the land does not have any frictional effect on the wind. In like manner, if the wind is blowing over the land in an offshore direction, values for the land, b =.000190 and β = 20 are applied. For the transition zones, a cosine transformation is used to compute the intermediate values. Within the transition zone, angle A is computed from 0 to 90° with 0 being land and 90° being sea. The new angle A is used in equations 12 and 13 to compute the transition values for b and beta.

$$b = .0001 \cdot (1.275 + .625 (sin A))$$
 (12)

$$\beta = 39.5 - 10.5 \cdot \sin A \tag{13}$$

The computed values for b and β are substituted in equations 10 and 11 to compute the surface wind speed and direction from the geostrophic wind.

The final step in subroutine WIND is to compute the effective wind speed which is carried over into subroutines FETCH and WAVES for determining effective fetch length and wave height. The effec-

tive wind speed is that which will generate waves which will in turn have an effect on the beach. If the wind is blowing directly onshore, the full force of the wind is used in generating waves which will hit the coast. However, if the wind is blowing directly offshore, small waves will be generated in the nearshore area (Resio and Hayden, 1973). Based on empirical observations, onshore winds are about three times as effective in generating local waves as offshore wind (Davis and Fox, 1974 and Owens, 1975). Therefore, a cosine transformation is used in subroutine WIND to compute effective wind speed from the surface wind speed and direction. When the wind is blowing directly onshore, the effective wind is equal to the surface wind. On the other hand, when the wind is blowing along the shore, the effective wind is equal to 2/3 of the surface wind, and when the surface wind is blowing directly offshore, the effective wind is 1/3 the surface wind. Although the values for effective wind may be rough in some cases, they seem to give good estimates where comparative wind and wave data are available.

Latitude	^α L	αs	v _h v _g ∟	v _h v _g s
0	61	40		
10	55	29		
20	49.5	23		
30	45	19	0.31	0.56
40	42	16	0.38	0.63
50	39	14.5	0.42	0.67
60	37	13	0.46	0.70
70	36	12.5	0.485	0.715
80	35	12	0.495	0.723
90	35	12		

Table 1. The angle between surface wind and geostrophic over land α_L and over sea α_s , and the ratio between surface wind speed and geostrophic wind speed over land and sea for different latitudes ϕ , according to Baur and Phillip (1938, p. 292).

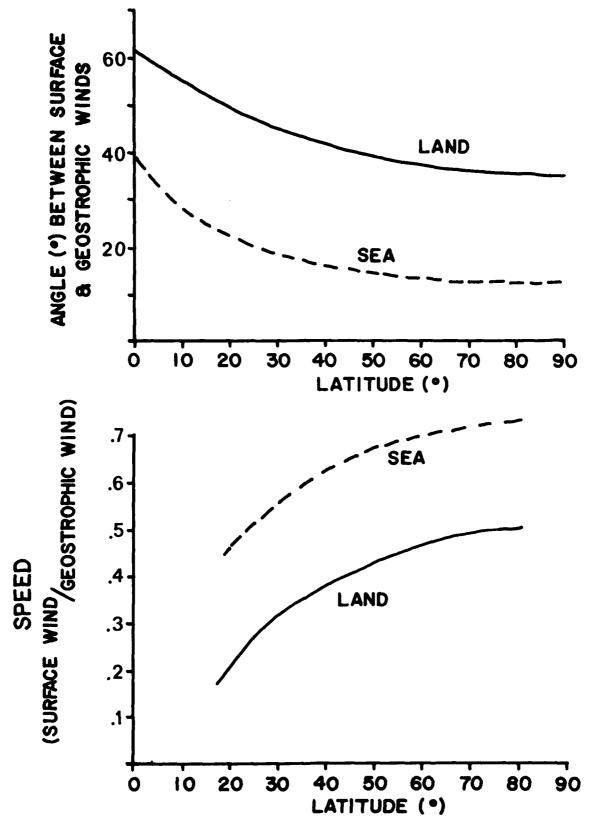


Figure 5. Angles between surface and geostrophic winds and ratio of surface to geostrophic wind speed.

Subroutine DECAY

Subroutine DECAY is used to determine the height of a given wave after it has decayed for a specified length of time. Snodgrass, et al (1966) presented empirical data on the attenuation of selected frequencies which they observed in their study of propagation of ocean swell in the Pacific. In general, they found the attenuation to be large within the limits of the wind area of the storm, and small outside the storm area. The empirical attenuation data were logarithmic coefficients reported in units of decibels per latitude degree of propagation distance. For the range of frequencies 0.06 to 0.08 Hertz, these data fit an attenuation function of the form

$$e^{-2ax}$$
 (14)

The state of the s

where a is the modulus of amplitude decay in degree $^{-1}$ = 0.1151 $_{\rm S}$, where $_{\rm B}$ is the logarithmic attenuation coefficient in decibels/degree, and where x is the propagation distance in degrees. The logarithmic attenuation coefficient versus frequency was plotted on a graph (Kaufman, 1973), and equation (15) was derived from the line on the graph.

$$\beta = 10^{\left(\frac{F-0.06}{0.0324}\right)^{-1}} \tag{15}$$

where F is the frequency of the wave being decayed.

The propagation distance x is found by multiplying the wave velocity 1.5606*T (where T is the wave period) times the time interval TINT. This distance is then reduced to degrees by multiplying it by the constant value

(using the circumference of the earth at the equator). To find the decayed wave height then, the original wave height is multiplied by the attenuation function. This decay factor was tested with several different wave heights, periods, and time intervals, giving very reasonable decay results, but there were no empirical data against which to check the results.

Subroutine ETIME

Subroutine ETIME is used to determine the amount of time (referred to as effective duration) which would be required to produce waves of a certain height by wind blowing at a given wind speed. A wave forecasting procedure developed by Sverdrup and Munk (1947), and revised by Bretschneider (1952, 1958) with additional empirical data is called the sverdrup-Munk-Bretschenider (SMB) method (C.E.R.C., 1973). The SMB curves for forecasting wave height are based on equation 16 from Bretschneider (1958).

$$\frac{gH}{U^2} = 0.283 \text{ Tanh } \left[0.0125 \left(\frac{gF}{U^2} \right)^{0.42} \right]$$
 (16)

where g is the acceleration due to gravity, H is the wave height, F is the effective fetch length, and U is the wind speed. Solving equation 16 for F gives

$$F = \frac{U^2}{g} \times \left[\frac{ARCTANH \left(\frac{gH}{0.283 \times U^2} \right)}{0.0125} \right]^{2.38}$$
(17)

Therefore, F is the effective fetch that it would take to generate waves of height H with a wind speed of U.

In terms of storm duration, the effective fetch equation is

$$F = \frac{S}{10}^{0.72} \times 10^{0.3} \times 0^{1.25}$$
 (18)

Where F is effective fetch, W is wind speed, and D is storm duration. Solving this for D, we have:

$$D = \left(\frac{F}{\frac{S}{10} 0.72 \times 10^{0.3}}\right)^{0.8}$$
 (19)

Therefore, D is the effective duration that it would take to build waves of a certain height (used to find the effective fetch) with winds of a given speed. Then, this effective duration is added to the current time increment of the storm to give a duration which takes into account the wave built in previous time increments.

Subroutine ETIME was tested by running the data from subroutine WAVES back through it to arrive at the original data. For example,

the following wave heights and wind speeds were tested using the C.E.R.C. charts and the subroutine (Table 2).

				Effective Duration			
Wave	Height	Wind Sp	Speed	From Si	MB Curves	Derived	
14	ft	80	kts	1.1	hrs	1.05	hrs
1	ft	12	kts	1.4	hrs	1.38	hrs
45	ft	80	kts	9.9	hrs	10.23	hrs
3	ft	12	kts	21.0	hrs	21.40	hrs
14	ft	30	kts	21.0	hrs	21.76	hrs

Table 2. Test for subroutine ETIME comparing effective duration available from SMB curves with derived duration for selected wave heights and wind speeds.

Subroutine FETCH

Subroutine FETCH is used to determine the maximum wind generating area and the average wind speed within that area for a circular storm on the open ocean. Three cases are considered for determining the actual storm fetch as the storm passes across a shoreline.

For a storm on the open ocean, the wind speed varies with the slope (first derivative) of a normal curve as one crosses the storm. The area of the storm with wind speed greater than one half of the maximum lies within a ring with an outside radius of 1.92 standard deviations and an inner radius of 0.32 standard deviations. This ring between 1.90 and 0.30 was divided into 13 smaller rings. The area of each ring was found and multiplied by the wind speed at the midpoint of the ring. These products were summed and then divided by the total area to give an average wind speed of 0.8847 of the maximum wind speed. The total area was found to be 7.7598 square standard deviations. Since the wind is being generated in a circular pattern, approximately one quarter of the wind is blowing along each axis of a grid with its center at the origin. Since the storm fetch area has a shape resembling an ellipse, the fetch for winds blowing in any one direction can be considered an ellipse with an area equal to one quarter of the maximum wind generating area (1.9400 square standard deviations) and centered on the maximum wind speed circle (1 standard deviations). The short half diameter was taken to be 0.7 standard deviations (1.0-0.3).

> Area = π x a x b 1.9400 = 3.1415 x a x .7 a = 0.8822 standard deviations (20)

Therefore, the maximum storm fetch length is twice that, or 1.7644 standard deviations. Converting to storm radii, the fetch is 0.5881 radii.

Where the storm crosses the shoreline three cases must be considered; Case 1, where 1/3 R > x \geq 0; Case 2, where 0.4444 R > x \geq 1/3 R and Case 3, where x > 0.4444 R, where R is the storm radius and x is the distance from the center of the storm.

For Case 1, (1/3 R > x \ge 0) (Figure 6), it is assumed that the fetch area, centered on the 1/3 R circle, swings around so that the long axis of the ellipse is always pointing at the beach. We have maximum storm fetch in this case until Y is small enough that the fetch begins to cross the beach as it continues to swing. This point is Y1. From the Case One diagram, $D^2 = X^2 + Y1^2$ and $D^2 = X^2 + Y1^2$

 $(1/3 \text{ R})^2 + (1/2 \text{ F})^2$ so we have $X^2 + Y1^2 = (1/3 \text{ R})^2 + (1/2 \text{ F})^2$ or

$$Y1 = \sqrt{(1/3 R)^2 + (1/2 F)^2 - X^2}$$
 (21)

As the storm continues to cross the beach and rotate, the fetch decreases. When the 1/3 R circle of the storm crosses the beach, the storm has swung so that the beach is in the center of the fetch area. Therefore, the storm fetch is one half of maximum at this point (Y2). As the storm continues inland, the fetch length reaches a minimum as Y goes to zero (Y3). This minimum is determined by the equation 22.

$$FMIN = \frac{1 \times 1}{\sqrt{X^2 + Y2^2}} \cdot FMAX. \tag{22}$$

This gives a diminishing minimum as |x| goes to zero. The fetch then increases to a high equal to $1/2 \cdot FMAX$ as the 1/3 R circle again crosses the beach (Y4). The fetch then declines again, finally dropping to zero at Y4, at which point the storm is completely onshore. This gives the "eye of the storm" effect where the winds drop to a minimum when the storm is at its closet point of approach and then increase again as the storm moves on, and finally drop as the storm moves away. Cosine smoothing is used to smooth out the increases and decreases in fetch.

For Case 2 (0.4444 R > $x \ge 1/3$ R)(Figure 7), again we have full fetch until the fetch area comes into contact with the beach at Yl. The fetch then decreases as the storm moves onshore, hitting zero at Y2. The reason that the fetch doesn't drop and then rise again, as in Case One, is that the eye of the storm (the part inside the 1/3 R circle) never passes over the beach.

This case ends at X=0.4444 R because D in the case two diagram = $\sqrt{(1/2 \text{ R})^2 + (1/2 \text{ F})^2} = 0.4444$ R. When $X \cdot D$, then the geometric relationships that allow us to find Y1 no longer hold, since D, the distance from the origin to the beach, can no longer equal both $\sqrt{X^2 + Y1^2}$ and $\sqrt{(1/3 \text{ R})^2 + (1/2 \text{ F})^2}$. Again consine curves are used to smooth out the changes in storm fetch.

For Case 3, (x > 0.4444 R)(Figure 8) the storm fetch area never actually crosses the beach, but the fetch decreases as the storm passes over the shore farther along the coast. Thus the storm fetch is at its maximum until the fetch area starts going onshore at Y1. Here the fetch is perpendicular to the storm path. Y1 is at zero since at that point, the whole half of the storm that is generating alongshore waves is still over the water, so we still have full fetch. As the entire storm passes onshore (Y2), the fetch drops to zero. Y2 is $-\sqrt{R^2-(0.4444 \text{ R})^2}$. Y2 is frozen, using an X value of 0.4444R.

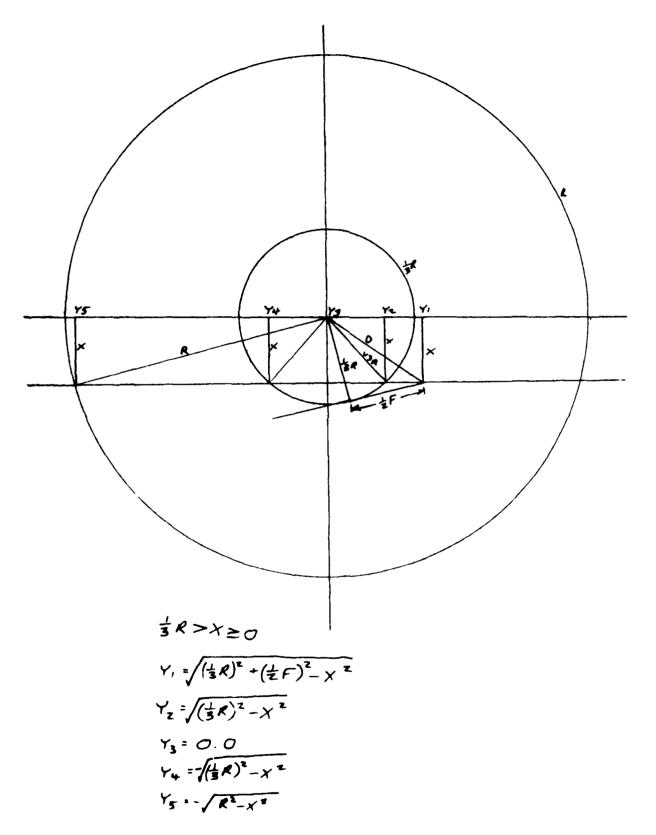


Figure 6. Case 1 - storm fetch when the distance from center of storm X is less than 1/3 storm radius, R.

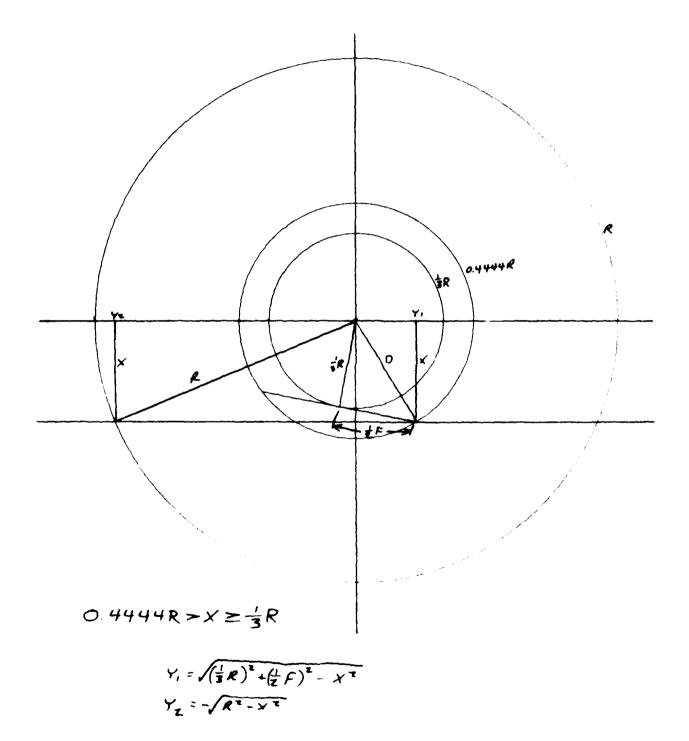
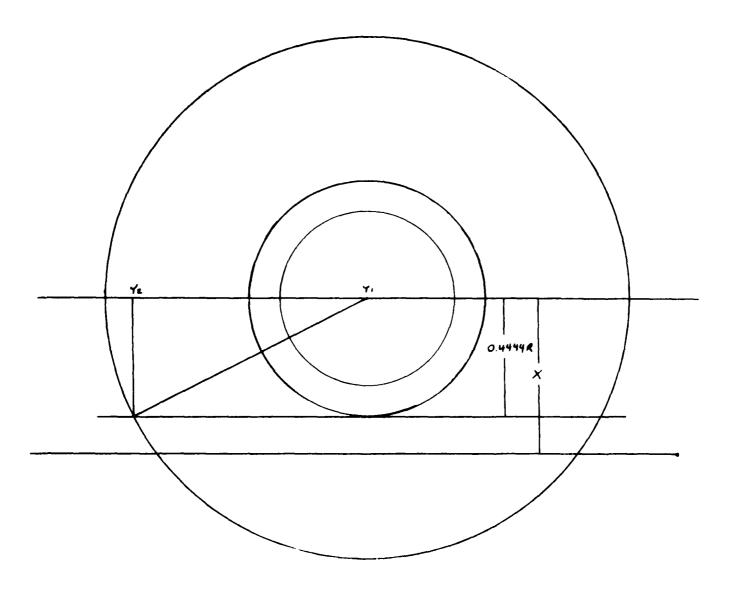


Figure 7. Case 2 - storm fetch when the distance from center of storm X is less than 0.4444 and greater than 1/3 storm radius, R.



$$Y_1 = 0 0$$
 $Y_2 : -\sqrt{R^2 - (0.4444 R)^2}$

Figure 8. Case 3 - storm fetch when the distance from center of storm $\,$ X is greater than 0.4444 storm radius, $\,$ R.

The reason that this is valid, is that from $X = 0.4444 \, R$ on out, the fetch area never actually passes over the beach, so that all these cases are essentially the same, as far as the size of their storm fetches goes. Only the distance from the storm fetch to the beach changes.

Subroutine FETCH was tested in two ways. First X and Y were varied independently, running Y from +R to -R for each value of X. This simulated storms with paths perpendicular to the shoreline. Second, X and Y were varied simultaneously by a constant amount, simulating storms crossing the shoreline at an angle. In both cases, the program produced continuous curves of fetch varying with storm location, with smooth transitions between all cases.

Subroutine WAVES

The equations for predicting significant wave height and period in subroutine WAVES are based on the Sverdrup-Munk-Bretschneider (SMB) method revised by Bretschneider (1958) and plotted as a series of curves by C.E.R.C. (1973). The SMB wave forecasting curves for fetches of 1 to 1000 miles are given in Figure 9. The wave prediction curves use the wind speed in knots, storm duration in hours, and storm fetch to calculate the significant wave height and significant wave period.

In order to use the SMB method in the model, the first task was to find equations to approximate the effective fetch from the wave prediction curves. The effective fetch is the limiting fetch which corresponds to a given wind speed and duration. The effective fetch is determined by moving to the left across the chart at the level of the wind speed until you hit the appropriated storm duration line. Then drop straight down to the fetch length axis from the intersection of the wind speed line with the duration line. This value on the fetch length axis is the effective fetch, if it is less than the actual fetch.

To develop an equation for effective fetch, the first problem is to determine the intercept of the proper duration line. To do this, the intercept of each line with the fetch length axis was plotted against the storm duration. Log scales were used on both axes since the original fetch length axis had a log scale and the duration lines themselves were spaced logarithmically. These formed a nearly linear trend and the equation for the line was found, in terms of the log axes, to be:

$$I = 10^{0.3} \times D^{1.25} \tag{23}$$

where I is the intercept and D is the storm duration.

The next step was to determine the slope of the storm duration lines given in the following equation.

$$F = \left(\frac{S}{10}\right)^{0.72} \times I \tag{24}$$

where F is the effective fetch, S is the wind speed, and I is the intercept of the duration line with the fetch length line. Combining equations 23 and 24 we get equation 25 for the effective fetch in terms of wind speed and storm duration:

$$F = (\frac{S}{10})^{0.72} \times (10^{0.3} \times D^{1.25})$$
 (25)

If this value is less than the actual storm fetch then it decreases the relevant fetch length which is used in the rest of the equations.

The SMB forecasting curves were constructed from equations 26 and 27, which were empirically derived by Bretschneider (1958).

$$\frac{gH}{U^2} = 0.283 \text{ TANH } \left(0.0125 \left(\frac{gF}{U^2} \right)^{0.42} \right)$$
 (26)

$$\frac{gT}{2pU} = 1.20 \text{ TANH } \left(0.077 \left(\frac{gF}{U^2} \right)^{0.25} \right)$$
 (27)

where g is the gravitational constant, p is PI (3.1459), H is the significant wave height, U is the wind speed, F is the effective fetch, and T is the significant wave period. Solving these equations for H and T, we get equations 28 and 29.

$$H = \frac{U^2 \times 0.283 \times TANH \left(0.125 \left(\frac{gF}{U^2}\right)^{0.42}\right)}{g}$$
 (28)

$$T = \frac{2pU \times 1.20 \times TANH \left(0.077 \left(\frac{gF}{U^2}\right)^{0.25}\right)}{g} (29)$$

The values for wind speed, storm duration, and effective fetch are then inserted into equations 28 and 29 to yield the wave height and period. For example in the case of a storm with a duration of 10 hours, a wind speed of 35 knots, and an actual fetch of 200 nautical miles, this gives us an effective fetch of 87.44 nautical miles, a wave height of 12.78 feet, and a wave period of 7.85 seconds. But suppose that we have a storm the same as the last one, but with a fetch of only 80 nautical miles. In this case the actual fetch is smaller than the computed fetch, so it remains as the effective fetch. This gives us a wave height of 12.36 feet and a wave period of 7.72 seconds.

The program has an option so that the results can either be metric or, to facilitate checking the results against the SMB curves, the results can be in nautical miles and feet. The subprogram was tested with numerous combinations of wind speeds, durations, and fetch lengths, with the results agreeing very well with the SMB forecasting curves.

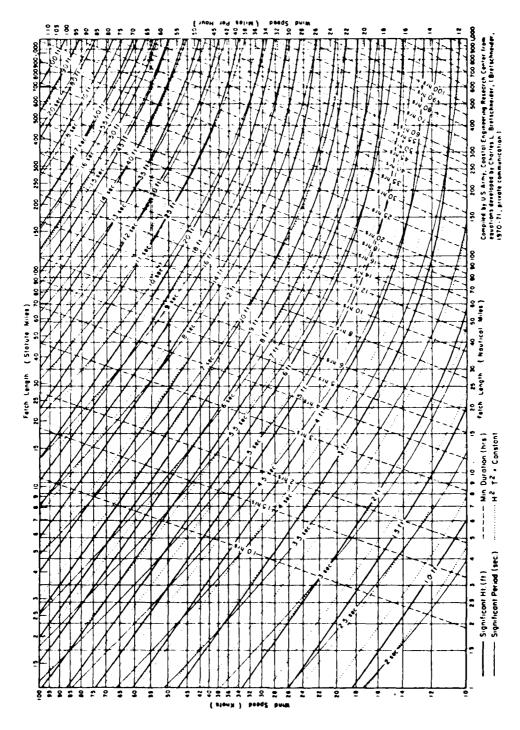


Figure 9. Deepwater wave forecasting curves as a function of wind speed, fetch length and wind druation based on the S.M.B. method.

Subroutine TIDE

Subroutine TIDE is used to determine the tide level at each hour, the spring tide range (ST), the neap tide range (NT), the number of days since the last spring high tide (TDAY), the hour of the last high spring tide (THR), and the tidal form number (FN). Four principal tidal components, $\rm M_2$ - Principal lunar, $\rm S_2$ - Principal solar, $\rm K_1$ - Lunar-solar diurnal, and $\rm O_1$ - Principal lunar diurnal are used for making a prediction of the hourly tide level. The periods of the semi-diurnal components (TM2 = 12.42 hours and TS2 = 12.00 hours) and the diurnal components (TK1 = 23.93 hours and TO1 = 25.82 hours) are constants in the subroutine. The tidal form number FN is used to classify the tides of a locality according to equation 30 (Defant, 1960, p. 306).

$$FN = \frac{K_1 + O_1}{M_2 + S_2} \tag{30}$$

The following classification based on Dietrich (1944, p. 69) is used to classify tides according to their form number.

FN = 0 - 0.25 Semi-diurnal tide FN = 0.25 - 1.50 Mixed-mainly semi-diurnal tide FN = 1.50 - 3.00 Mixed-mainly diurnal tide FN =greater than 3.0 Diurnal tide

As examples of the different types of tides, Immingham, England has a semi-diurnal tide with a form number of 0.11. San Francisco, California has a mixed, dominately semi-diurnal tide with a form number of 0.90. Manila has a mixed, dominately diurnal tide with a form number of 2.15, and Do San, Viet Nam on the Gulf of Tonkin has a pronounced diurnal tide with a form number of 19.2. The four major components are responsible for the general form of the tides and generally account for about 70 percent of the total variance. If the next three most important tidal components, N_2 , K_2 and P_1 are included, the percentage of the total variance increases to about 83% (Defant, 1960).

If some simplifying assumptions are made concerning the major tidal components, it is possible to make a good approximation of the hourly tide level from the maximum spring tide range, minimum neap tide range, and the tidal form number. First, it is necessary to assume that the diurnal components, K_1 and O_1 are approximately equal. Second, assume that the maximum spring tide range is equal

to the sum of the four major components according to equation 31.

$$ST = M_2 + S_2 - (K_1 + O_1)$$
 (31)

Next it is assumed that the neap tide range is approximately equal to the lunar components minus the solar components.

$$TN = M_2 + K_1 - (S_2 + O_1)$$
 (32)

If K_1 and O_1 are approximately equal, it follows that

$$TN = M_2 - S_2$$
 (33)

The form number is the diurnal components over the semidiurnal components in equation 30, however, if $K_1 = 0_1$, then

$$FN = \frac{2 K_1}{M_2 + S_2} \tag{34}$$

By combining the equations for the form number (Equation 30), the spring tide range, and the neap tide range, it is possible to solve for M_1 and S_2 .

$$M_{1} = \frac{ST + TN}{4(1 + FN)} \tag{35}$$

and

$$S_2 = \frac{ST - TN}{4 (1-FN)}$$
 (36)

If it is now assumed the lunar components are proportional, an approximation of the $\rm K_1$ component can be derived from the $\rm M_2$ component

$$K_1 = FN \cdot M_2 \tag{37}$$

and the solar components are related in like manner, therefore,

$$O_1 = FN \cdot S_2 \tag{38}$$

The amplitude of the maximum spring tide was taken at the last previous spring tide for each run, therefore, the phases for the four major components are considered 0 at that time. By computing the time differences from the last spring high tide to the hour for the prediction, the contribution for each tidal component can be calculated. The argument (ARG) is equal to 2 pi times the number of hours since the last spring tide. The tide is computed by adding together the contribution for each of the tidal components.

TIDE =
$$M_2 \cdot \cos\left(\frac{ARG}{T_{M_2}}\right) + S_2 \cdot \cos\left(\frac{ARG}{TS_2}\right) + K_1 \cdot \cos\left(\frac{ARG}{TK_1}\right) + O_1 \cdot \cos\left(\frac{ARG}{TO_1}\right)$$
 (39)

Although some rough approximations were made in deriving the major tidal components from the spring tide range, neap tide range and form number, the resulting tide predictions work out quite closely with the tide tables. The tide tables give the time of high and low tides for each day, and the predicted times of high and low tides fall within 1 hour using subroutine TIDE. The subroutine was tested for Plum Island, Massachusetts, Cedar Island, Virginia, Sapelo Island, Georgia and Mustang Island, Texas, and gave satisfactory predictions for each of the areas.

One of the major reasons for making tidal predictions in the model is to determine the effect that tides have on the nearshore bottom slope. The slope at low tide SLBOT, the slope at high tide SLTOP and the tide level are used to determine the intermediate slope between high and low tide.

$$SLOPE = SLBOT + TLOC *(SLOTP - SLBOT)$$
 (40)

The final tide level which is included with the output is computed by adding the relative tide level TIDX to the mean low tide level TMEAN.

Subroutine SURF

Breaker height, angle and longshore current velocity are computed in Subroutine SURF. The critical value, $\rm H_b/h_b=.78$ where $\rm H_b$ and $\rm h_b$ are breaker height and depth, are used for a breaking criterion (Munk, 1949). Applying any wave theory and assuming conservation of energy flux, Komar and Gaughan (1972) derived the relationship

$$H_b = 0.73 \text{ cm} + .383 \text{ g}^{1/5} (T H_c^2)^{2/5}$$
 (41)

where $\mathbf{H}_{\mathbf{b}}$ is the breaker height, g is gravity, T is wave period and $\mathbf{H}_{\mathbf{m}}$ is the deep water wave height.

The breaker angle α_b is computed by first finding the shallow water wave length and then taking the ratio of shallow water to deep water wave length using Snell's law to determine the breaker angle.

$$\sin \alpha_b = \sin \alpha_0 \ TANH \left(\frac{2\pi H_b}{L_b} \right) \tag{42}$$

where α_c is the deep water wave angle which is assumed to be the same as the wind angle, H_b is the breaker height and L_b is the breaker depth.

The refracted breaker height, ${\rm H}_{\rm R}$, is obtained from the refraction coefficient, KR,

$$KR = \frac{\cos (\alpha_0)}{\cos (\alpha_b)}$$
 (43)

which is multiplied times the breaker height, $H_{\rm h}$

$$H_{R} = KR \bullet H_{b} \tag{44}$$

Four different options are available for computing the long-shore current velocity. The longshore current equations by Longuet-Higgins (1970), Komar and Inman (1970), Fox and Davis (1972) and Coastal Engineering Research Center (C.E.R.C., 1973) used basic different assumptions with the same set of variables. The variation

in longshore current velocity across the surf zone and along the shore, as well as differences in nearshore topography brought about by bars and rip channels make any prediction of average longshore current velocity very difficult. However, in making predictions about the surf zone, it is essential to at least have a good estimate of the maximum longshore current velocity.

The radiation stress theory of Longuet-Higgins (1970) has been tested with laboratory data from Galvin and Eagleson (1965), and field data from Putman, Munk and Traylor (1949). The long-shore current velocity in the surf zone, $V_{\rm b}$, is a function of the bottom slope, m, the breaker height, $H_{\rm b}$, and breaker angle, $\alpha_{\rm b}$, between the wave crest and the shoreline (Longuet-Higgins, 1970).

$$V_b = M_1 m (gH_b)^{1/2} \sin 2\alpha_b$$
 (45)

where M_1 , the friction factor is:

$$M_{1} = \frac{0.694 \, \Gamma \, (2\beta)^{-1/2}}{f_{f}} \tag{46}$$

The longshore current, V_b is measured at the breaking position and Γ is a mixing coefficient with a range of 0.17 (little mixing) to 0.5 (complete mixing) with a mode at about 0.2. The depth to height ratio in shallow waver, β , is taken to be 1.2 and f_f the friction coefficient is set at 0.01. By inserting the above values in equation 46, the value for M_1 becomes 9.0. Therefore, the long-shore current equation according to Longuet-Higgins (1970) can be reduced to:

$$V_b = 9.0 \text{ m } (gH_b)^{1/2} \sin 2\alpha_b$$
 (47)

When equation 47 was applied to test sets of field and laboratory data by C.E.R.C. (1973), the data yields predictions that average about 0.43 of the measured values. The measured values were taken in the fastest field of flow shoreward of the breaker zone, whereas the predictions were made for longshore current at the line of breakers. Therefore, it has been proposed by C.E.R.C. (1973) that the Longuet-Higgins equation be multiplied by 2.3 to yeild the C.E.R.C. equation:

$$V = 20.7 \text{ m } (gH_b)^{1/2} \sin 2a_b$$
 (48)

Komar and Inman (1970) derived a longshore current equation based on radiation stress. Where the radiation stress components defined by Longuet-Higgins and Stewart (1964) is the excess flow of momentum due to the presence of waves. The Komar and Inman (1970) equation is:

$$V = C_1 U_m \frac{Tan \beta}{C_f} sin \alpha_b cos \alpha_b$$
 (49)

where V is the longshore current velocity, Tan β is the beach slope, C_f is the bottom frictional drag coefficient. U_m is the maximum horizontal component of the orbital velocity of the waves and C_1 is a dimensional coefficient of proportionality. However, Komar (1969) suggested that:

$$(Tan \beta cos \alpha_b)/C_f = constant$$
 (50)

indicating that the variation in beach slope does not produce a change in longshore current velocity. Therefore, the Komar and Inman (1970) longshore current equation becomes:

$$V = C_1 U_m \sin \alpha_b \tag{51}$$

A fourth equation developed by Fox and Davis (1972) uses empirical data subjected to linear regression analysis to predict longshore current velocity. The linear regression analysis is based on 3 sets of data collected at Stevensville, Michigan (Fox and Davis, 1970), Holland, Michigan (Fox and Davis, 1971a) and Sheboygan, Wisconsin (Fox and Davis, 1972). Each set of data consists of 360 observations taken at 2 hour intervals for 30 days of longshore current speed and direction, breaker height, period and breaker angle. Using a stepwise regression analysis, the contribution of each variable was tested separately, and then in various combinations. The ratio, H_h/T is related to the mass flux on volume of water which enters the surf zone and must be removed by the longshore current. The breaker angle, $\alpha_{\mbox{\scriptsize h}}$, defines the angle between the breaker crest and the shoreline and is therefore related to the momentum transfer in the longshore direction. Using the regression program, a series of combinations was tested for the sin of the angle including sin A, sin 2A, sin 3A, sin 4A... sin 8A. The closest fit was obtained when sin 4A was used for the angles. For the 1969 set of data from Stevensville, Michigan, the following equation,

$$V = 5.42 \left(\frac{H_b}{T}\right) \sin 4A \tag{52}$$

gave the best fit and accounted for 83.5 percent of the total sum of squares. For the 1970 data from Holland, Michigan, the coefficient of proportionality was 3.47 and the equation accounted for 78.8 percent of the total sum of squares. For the 1972 data from Sheboygan, Wisconsin, the coefficient was 2.98 and the equation accounted for 77.8 percent of the total sum of squares.

The three areas differed in the nearshore bottom slope and the occurrence of sand bars which influenced the coefficient of proportionality. The coefficient for each case was approximately equal to 100 times the bottom slope. Therefore, the longshore current velocity according to Fox and Davis (1973) is

$$V = 100 \text{ m} \left(\frac{H_b}{T}\right) \sin 4 \alpha_b \tag{53}$$

When the four longshore current equations were tested in the model, the equation by Longuet-Higgins (1970) and Fox and Davis (1973) gave very similar results for breaker angles up to about 20 degrees. For higher breaker angles, the predicted results from the Fox and Davis (1973a) equation were too low. The values for longshore current predicted by Komar and Inman (1970) and C.E.R.C. (1973) were consistantly too high. Although the four equations are available as options, it is recommended that the Longuet-Higgins (1970) equation be used for making predictions. If possible, it is best to test predictions with hindcast data from the same area.

Subroutine ENRGY

Subroutine ENRGY is used to determine the wave energy during each hour of the storm which is summed to give the total wave energy for the storm. The deep water wave energy $\rm E_{0}$ (C.E.R.C., 1973) is given by

$$E_0 = \frac{\rho g H^2 L_0}{8} = \frac{5.12 \rho g (HT)}{8}$$
 (54)

where p is the mass density of the water which is 1.94 slugs/cubic foot for fresh water and 2.0 slugs/cubic foot for salt water, H is the deep water wave height and T is the wave period. Conversion factors are included to change from foot pounds/foot to Joules/meter. The subroutine was tested using wave energy calculation from previous studies (Fox and Davis, 1971b).

Subroutine ARCTA

Subroutine ARCTA is a customized arctangent subroutine for finding the angle in radians from the arctangent of a function (Louden, 1967, p. 119). The library arctangent function ATAN accepts as an argument the tangent of an angle (sin/cos) and produces as output the angle in radians. Since the tangent of an angle repeats itself every 180 degrees, it is not possible to use the library function ATAN to determine a full range of angles from 0 to 360 degrees. To compute the correct angle for all possible combinations of X and Y, it is necessary to test for positive, near zero and negative X, and positive near zero and negative Y. The IF statements accomplish these test and produce an angle in radians ranging from 0 to 2 pi.

HINDCAST ANALYSIS WITH COASTAL STORM MODEL

Hindcast Tests of Model

The coastal storm simulation model can be used to hindcast wind, wave and current conditions at a shore site during the passage of a coastal storm. Hindcast analysis differs from forecast analysis discussed in the previous section because exact storm positions are known in hindcasting, whereas a constant azimuth and storm velocity are used in forecasting. The results of hindcast analysis at several sites are included in Appendix C. On the Great Lakes, the sites include Holland and Stevensville, Michigan, and Sheboygan, Wisconsin. On the east coast of the United States and Canada, sites include the Magdalen Islands on the Gulf of Saint Lawrence; Plum Island, Massachusetts; Cedar Island, Virginia and Sapelo Island, Georgia. Mustang Island, Texas was studied on the Gulf Coast. On the west coast of the United States, hindcasts were made for Monterey, California and South Beach, Oregon.

Sites were selected for hindcast analysis which had weather and wave data available for several storms. Several of the sites were studied by Davis and Fox using time series analysis from 1969 through 1975. Other sites were chosen in which there was good beach profile data which could be correlated with wave and current conditions during a storm.

Stevensville, Michigan, July 1969

A storm which passed over Lake Michigan in late July 1969 has been choosen as an example of hindcast analysis. When the storm passed over, a 30 day time-series study was being conducted at Stevensville, Michigan by Fox and Davis (1970a and b). Stevensville is located on the southeastern shore of Lake Michigan about 11 kilometers south of Benton Harbor, Michigan. The shoreline is oriented roughly north-south with an average nearshore slope of about 0.033.

The storm which affected the Stevensville area was tracked from 2000 on July 26, 1969 through 0800 on July 30 (Table 3). The size, shape, intensity and path of the storm were interpreted from weather maps for July 26 through 30, 1969. When the storm was closest to the coastal site at Stevensville, the barometric pressure at the center of the low was estimated as 994 millibars. The pressure at the largest encircling isobar was 1012 millibars, and therefore, the maximum pressure included in the storm was 1014.6 millibars. The storm had an elliptical shape with the length of the major half axis equal to 960 kilometers and the minor half axis

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equal to 700 kilometers. The orientation of the major half axis was 30 degrees west of north. For this particular example, the equation by Fox and Davis (1972) was used to compute the long-shore current velocity.

The position of the shoreline at Stevensville is given in the X-Y coordinate system with X equal to 1043 and Y equal to 486 kilometers. The X axis runs east-west and the Y axis north-south with the origin located to the southeast of Stevensville. The latitude at the shore location is 42° north and the onshore azimuth is 90 degrees. The nearshore slope from the shoreline out across the nearshore bars is 0.033. The average fetch distance for the southeastern shore of Lake Michigan is about 200 kilometers.

The storm positions were plotted at 6 hour intervals in kilometers on the X-Y coordinate system (Table 3). The initial position of the storm at 2000 on July 26 was to the northwest of the shore site with X equal to 333 and Y equal to 1229 kilometers. The storm passed over the shoreline about 0230 on July 28. At that time, the storm center was located 178 kilometers north of the shore site (664-486=178 kilometers). When the storm tracking was complete, the final storm position at 0800 on July 30 was X equal to 1907 and Y equal to 1830 kilometers. In general, the storm made a loop swinging down from the northwest, passing eastward across the shore, then moving off to the northeast.

The X1, Y1 coordinate is oriented with the X1 axis parallel to the shore and the Y1 axis normal to the coast (Figure 2). The origin of the X1, Y1 coordinate system is at the center of the storm with the positive XI direction to the right and the positive YI direction toward the coast. The X1, Y1 coordinate system is used to locate the shore position with reference to the center of the storm. The units of the X1, Y1 coordinate system are in terms of The storm radius is 1.5 times the length of the storm radii. major half axis which is measured from the center of the storm to the largest encircling isobar. Using an inverted normal curve to simulate the storm cross-section the largest encircling isobar is defined as 2 standard deviations from the center of the storm. The full radius would extend out 3 standard deviations from the center of the low. As the storm approaches shoreline, the Yl value decreases, and it becomes negative after the storm has passed over the coast. When the storm is to the north of shore site the X1 values are positive. Therefore, the storm at Stevensville remained to the north of the shore site for the entire run.

In the hindcast analysis, the barometric pressure decreases from 1012.4 millibars on July 26 to a minimum of 995.2 at 2200 on July 27, then increased to 1014.6 on July 30, 1969 (Table 3 and Figure 10). In the actual barometric pressure record at Stevens-ville, the pressure dropped to 1000.2 millibars (29.54 inches) at

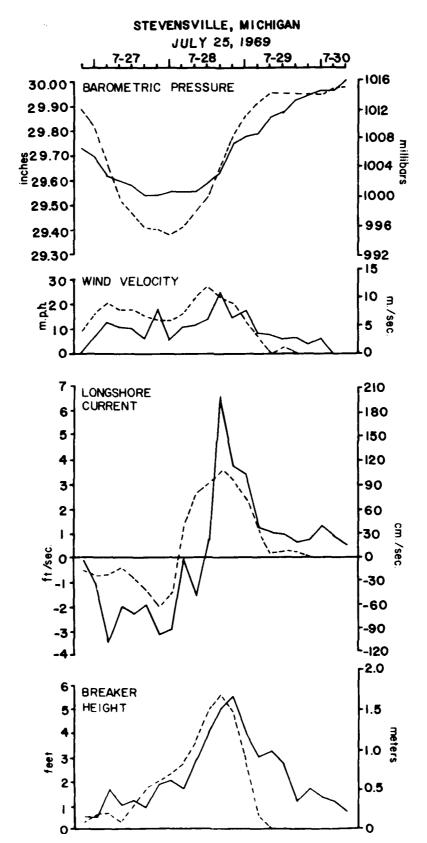


Figure 10. Observed and hindcast curves for barometric pressure, wind velocity, longshore current and breaker height at Stevens, ville, Michigan, July 26-30, 1969.

2000 on July 27. Since small scale weather maps were used for determining the storm position and pressure values, more accurate results could be obtained by using large scale maps available at 3 hour intervals. The plot for barometric pressure is a function of accurate plotting of the storm positions and a careful estimate of the size and intensity of the storm. Therefore, a correspondence of the observed and hindcast curves for barometric pressure is not a test of the predictive capabilities of the model, but is more a test of the accuracy of the weather maps and of the plotting ability of the operator.

In the wind observations at Stevensville, the maximum wind speed during the storm was 12.4 meters/second (28.4 knots) at 1600 on July 28 (Fox and Davis, 1970a). For the hindcast, the maximum wind speed was 12.0 meters/second at 1400 on July 28. The overall pattern for the observed and hindcast winds are also quite similar (Figure 10).

The observed longshore current velocity at Stevensville reached -116 centimeters/second (northerly) at 2200 on July 27, and 215 centimeters/second (southerly) at 1800 on July 28 (Fox and Davis, 1970a). The maximum hindcast values were -62.9 centimeters/second at 2100 on July 2/ and 101.9 centimeters/second at 1700 on July 28 (Figure 10 and Table 3). The instantaneous longshore current values were much higher in the observed than the hindcast values, however, the minimum and maximum values occurred within one hour of the hindcast values. The observed longshore current during the storm was exceptionally high and may be accounted for by the well developed trough between the nearshore bar and the shore which channeled the current along the coast. The value used for nearshore slope could also be too low in the hindcast. When the longshore current curve was smoothed using the 15 term Fourier plot, the hindcast more closely resembles the observed curve for longshore current (Fox and Davis, 1970a and b).

The observed and hindcast curves for breaker height are very close (Figure 10). For the observed curve, the maximum breaker hieght of 1.82 meters (6 feet) occurred at 1800 on July 28. For the hindcast, the maximum height was 1.83 meters, also at 1800 on July 28. The overall shapes of the observed and hindcast curves for breaker height are also very similar but the hindcast curve drops off more rapidly than the observed curve (Figure 10).

Additional Hindcast Examples

Additional examples of the hindcast tests are presented in Appendix B and Figures 26, 27, 28 and 29. Since the input parameters are given in Appendix B and comparative plots are in Figure 26 through 29, a full discussion will not be included for each of the examples.

The observed data for the Holland, Michigan examples was extracted from a 30-day time series study during July 1970 (Fox and Davis, 1971a). Holland is located on the southeastern shore of Lake Michigan about 96 kilometers north of Stevensville. The hind-cast values for longshore current and breaker height were quite close for July 3, 1970 (Figure 11), but were somewhat low for July 18, 1970 (Figure 12).

The observed data for the Sheboygan, Wisconsin examples come from a 30-day time series study conducted at Sheboygan during July 1972 (Fox and Davis, 1973a). Sheboygan is located on the western shore of Lake Michigan about 72 kilometers north of Milwaukee, Wisconsin. At this location, the storms moved in a northeast direction and generally offshore. Therefore, the characteristic reversal of longshore current direction observed on the eastern shore of Lake Michigan was not observed at Sheboygan. The curves for July 16 show generally low observed and hindcast curves for longshore current and breaker height with pronounced peaks on July 17, 1972 (Figure 13). The curves for July 22 show a substantial drop in barometric pressure, but very low values for longshore current and breaker height (Figure 14). There is a reversal in longshore current direction on July 24 for both the observed and hindcast curves. Since the winds were blowing predominately offshore, the waves and longshore currents are quite subdued.

Several additional tables of hindcast results are given in Appendix B. The output for the Atlantic, the Gulf and the Pacific coasts of the United States includes a variety of conditions for storms of varying sizes, shapes and intensities. Tidal predictions are also included for the oceanic sites where the tide tables are available. Some of the hindcast results followed quite closely with the observed data, while at other places, the fit was not as good as expected.

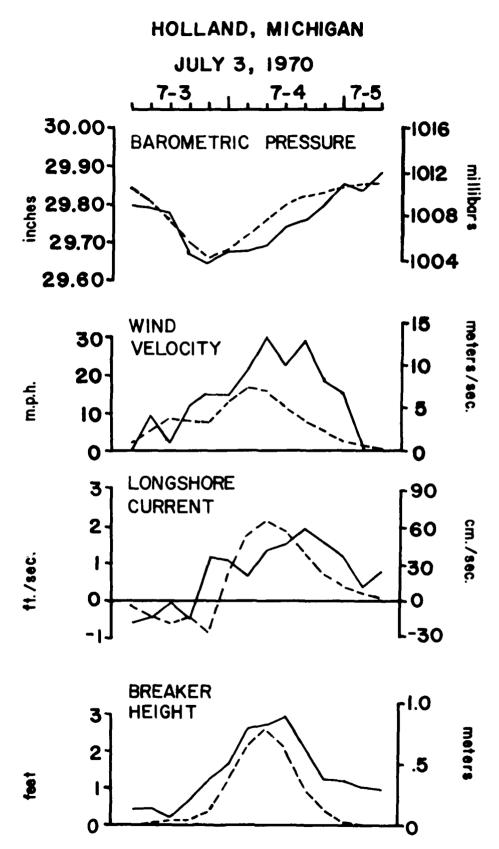


Figure 11. Observed and hindcast curves for Holland, Michigan, July 3-5, 1970.

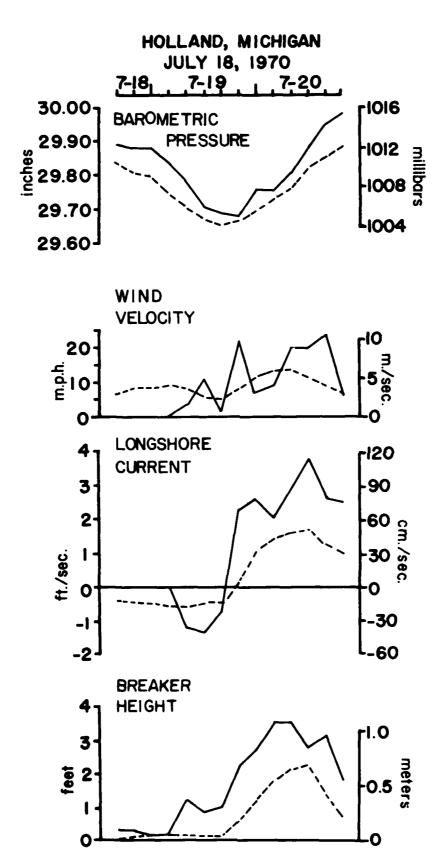


Figure 12. Observed and hindcast curves for Holland, Michigan, July 18-20, 1970.

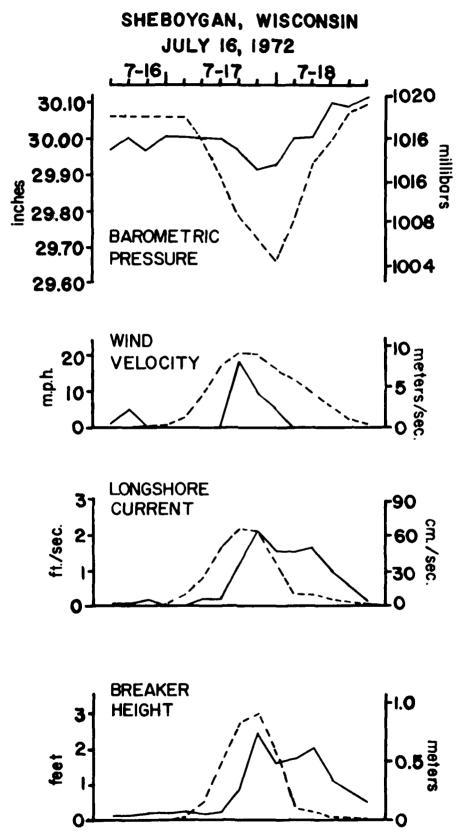


Figure 13. Observed and hindcast curves for Sheboygan, Wisconsin, July 16-18, 1972.

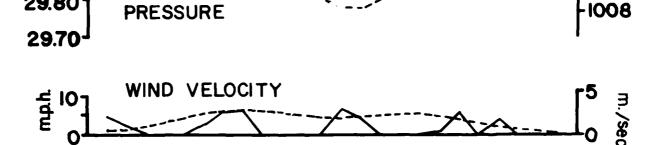
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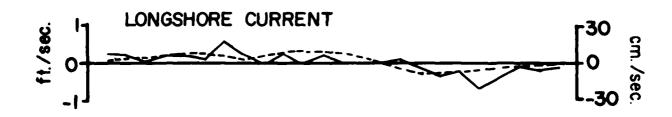
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Observed and hindcast curves for Sheboygan, Wisconsin, Figure 14. July 22-26, 1972.

FORECAST ANALYSIS WITH THE COASTAL STORM MODEL

Short-term Forecasts

Short-term predictions of wave height and longshore current velocity can be made using the forecast mode of the coastal storm model. Given the initial coordinates of the storm, its size, shape and intensity, along with the position and orientation of the shoreline, it is possible to predict the wind, waves and currents as the storm passes over the coast. In the forecast mode, the storm azimuth and velocity are used to plot a straight storm path as the storm proceeds toward the shore. In general, the forecast is run for 72 hours which is the usual limit for short-term weather prediction. Examples of output using the forecast mode are given in Appendix C.

In the forecast mode, it is assumed that the size, shape and intensity of the storm remain constant, as well as the direction of the storm path and the speed of the storm along the path. The operators experience with weather prediction plays an important role in estimating the speed and path of the storm. For 12 to 24 hours, the speed and path may remain fairly constant, but for longer periods of time, the storm may veer off on another path or change its speed along the path. Since it is almost impossible to predict the path of a storm for several days, a series of diagrams have been devised for predicting wave and current conditions for a storm with a given size, shape and intensity, but without a fixed storm path.

Circular Storm Test

A circular storm test is used as an example to explain how the forecast mode works with the coastal storm model. The circular storm test is based on a series of intense storms which crossed over the Oregon coast during late fall of 1973 (Fox and Davis, 1974). The size, shape and intensity of the storm, and the orientation of the shoreline are similar to those encountered at South Beach, Oregon in November 1973 (Fox and Davis, 1974). The computer listings for the circular storm are given in Appendix C.

For the circular storm, the barometric pressure at the center of the low was set at 1000 millibars. The pressure at the largest encircling isobar was placed at 1020 millibars to give a range of 20 millibars within the central portion of the storm. The plot of barometric pressure for the circular storm model is generated by rotating an inverted normal curve around its center. Therefore, the barometric pressure surface has a basin shape with the low pressure at the center, the steepest pressure gradient at one standard deviation out from the center, and gradually reaches a

maximum pressure at 3 standard deviations away from the center. Since the outer margin of a storm often interfers with other high or low pressure systems, it is assumed that the largest encircling isobar occurs at 2 standard deviations away from the center, or 2/3 of the total storm radius. Therefore, the total storm radius would be 1.5 times the radius measured at the largest encircling isobar, and the maximum pressure at the margin of the storm would be 1.145 times the pressure range from the center to the largest isobar. In the circular storm test, the pressure at the 2 standard deviations is 1020 and the maximum pressure included in the storm is 1022.9 millibars at 3 standard deviations.

The size of the storm is determined by measuring the lengths of the major and minor half axes of the storm. For an elliptical or wave-shaped storm, the length of the major half axis is measured from the center of the low to the largest isobar, where it is farthest from the storm center. The length of the minor half axis is measured at right angles to the major half axis from the storm center to the largest isobar. A circular storm exists when the major and minor half axes are equal. For the circular storm test, lengths of the major and minor half axes are 300 kilometers. When the major and minor half axes have different lengths, the orientation of the major half axis is plotted in degrees from north. Therefore, for the circular storm test, the orientation of the major half axis is 0 degrees.

For the circular storm model, the storm velocity is set at 40 kilometers/hour with a storm azimuth of 90 degrees. This means that the storm will proceed from its initial position along a path 90 degrees east of north at 40 kilometers/hour. Therefore, if the storm is tracked for 30 hours, it will move a distance of 1200 kilometers.

For the shore position coordinates in the circular storm test, X is set at 1000 and Y is set at 0 kilometers. In the X-Y coordinate system used for plotting storm position and shore location on a weather map, the shore site is located 1000 kilometers east along the X axis and O kilometers north along the Y axis. The shore latitude for the test case is 43 degrees north, the approximate latitude of the Oregon coast. The onshore azimuth is 90 degrees which indicates that the coastline runs north-south with land to the east and sea to the west, the same orientation as the Oregon coast. The nearshore slope is at 0.033 which is close to the average nearshore slope in the different areas studied. A value of 1000 kilometers was used for the average fetch which would indicate open ocean. If the average fetch is greater than the storm radius, the fetch distance will not have an effect on the wave and current calculations. However, if the fetch distance is smaller than the storm radius, the waves will be fetch limited when the average fetch is less than the effective fetch.

In compiling the output for the circular storm test, the storm outlined above was tracked along 12 paths normal to the shoreline and parallel to each other (Figure 15). The location of the shore site, orientation of the shoreline, and three storm tracks are given in Figure 15A. The diagram extends 1200 kilometers in a north-south direction, 600 kilometers to the north (positive) and 600 kilometers to the south (negative) of the study site which is located in the center of the diagram. The diagram also extends 1200 kilometers in an east-west direction, 600 kilometers offshore (negative) and 600 kilometers onshore (positive) of the shore site. A time scale is included along the bottom of the diagram to indicate the length of time in hours from the initial tracking of the storm to any position along the storm path. The speed of the storm was set at 40 kilometers/hour, therefore, if the storm started at the left edge of the diagram, its center would pass over the shoreline after 15 hours and would move off the right side of the diagram after 30 hours.

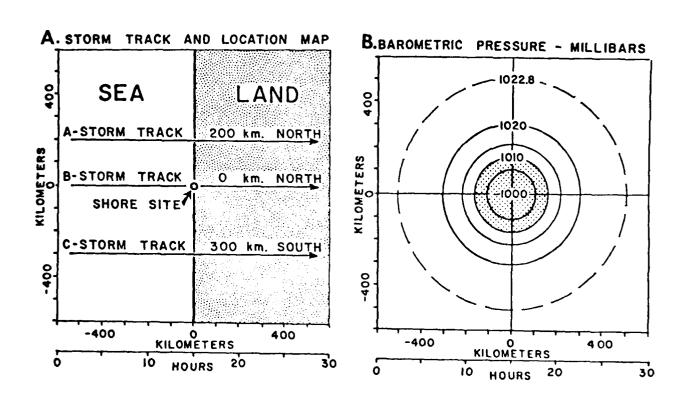


Figure 15. Map of storm tracks and time-distance plot of barometric pressure for a circular storm.

Barometric Pressure

The barometric pressure diagram (Figure 15B) is set up with the same coordinate system as the storm track diagram (Figure 15A). The barometric pressure which would be recorded at the shore site is plotted at the storm location as the storm moves across the diagram. After 12 storms were tracked normal to the shore at 100 kilometer spacings along the shore, the barometric pressure values plotted along the storm tracks were contoured to produce the barometric pressure diagram (Figure 15B).

The use of the barometric pressure diagram can be explained by examining a series of profiles through the diagram (Figure 16). Three storm tracks are plotted on the pressure diagram indicating storms which moved from west to east across the shoreline. Barometric pressure profiles are shown when the storm track is 200 kilometers north, directly over the shore site, and 200 kilometers south of the site (Figure 16). When the storm passes 200 kilometers north of the shore site, the pressure at the shore location drops from 1022.8 millibars to 1013.4, then increases again to 1022.8. When the storm passes directly over the shore location, the pressure at the shore site drops from 1022.8 millibars to 1000.0, then increases to 1022.8. When the storm passes 200 kilometers south of the shore site, the pressure profile is identical to the profile which was made 200 kilometers north of the shore site. For a circular storm, the barometric pressure diagram is symmetrical, so that profiles cut through the storm a given distance north or south of the shore site will result in identical patterns.

It should be emphasized that the values plotted at the storm locations are for observations recorded at the shore site when the storm follows along a given path. Therefore, the pressure profile in Figure 16 are profiles of the pressure at the shore site when the storm passes to the north, over the site, or to the south. Although the diagram for barometric pressure is identical to a weather map with storm center located directly over the shore, it should not be interpreted in that way. The X axis represents time, and distance along the X axis is used to show the storm position at a given time. With the other diagrams, such as wind speed and breaker height, it is impossible to use the time distance diagram as a map with the storm center located at the shore site.

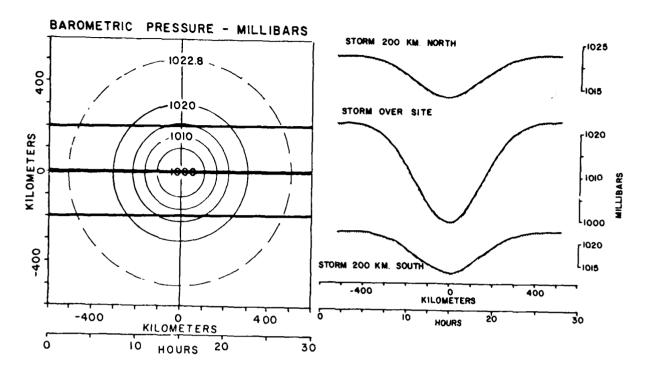


Figure 16. Time-distance plot of barometric pressure and pressure profiles 200 km north, over the site, and 200 km south of the site.

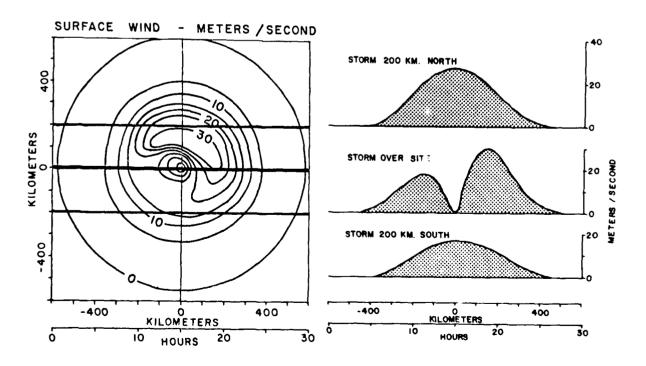


figure 17. Time-distance plot of surface wind speed and three profiles of wind speed in a circular storm.

Wind Speed and Direction

The time distance diagram for surface wind speed is not symmetrical because the maximum wind speeds occur when the storm track is between 100 and 200 kilometers north of the study site (Figure 17). When the storm passes 200 kilometers north of the study site, the surface wind speed reaches 28 meters/second at the shore site. However, when the storm passes directly over the shore site, the surface wind speed reaches 18.8 meters/second as the storm approaches, drops down to zero as the center of the storm passes over the coast, then increases to 30.8 meters/second. When the storm track is 200 kilometers to the south of the shore site, the surface wind speed reaches 17.0 meters/second. The highest wind velocities are recorded when the storm passes to the north, and after the storm has passed over the coastline.

To understand the wind pattern during a coastal storm, it is necessary to consider both wind speed and direction. Time-distance diagrams are plotted for surface wind speed, wind direction, onshore component, and alongshore component of the wind in a circular storm (Figures 18A, B, C and D). The surface wind speeds in the storm were computed by the geostrophic wind equation (Equation 6) with corrections applied for speed and direction (Equations 10 and 11) to account for the frictional effects of land or sea (Figure 5). For example, at 40° north latitude, the angle between the surface wind and the geostrophic wind would be 42° if the wind is from the land, and 16° if the wind is from the sea (Figure 5 and Table 1).

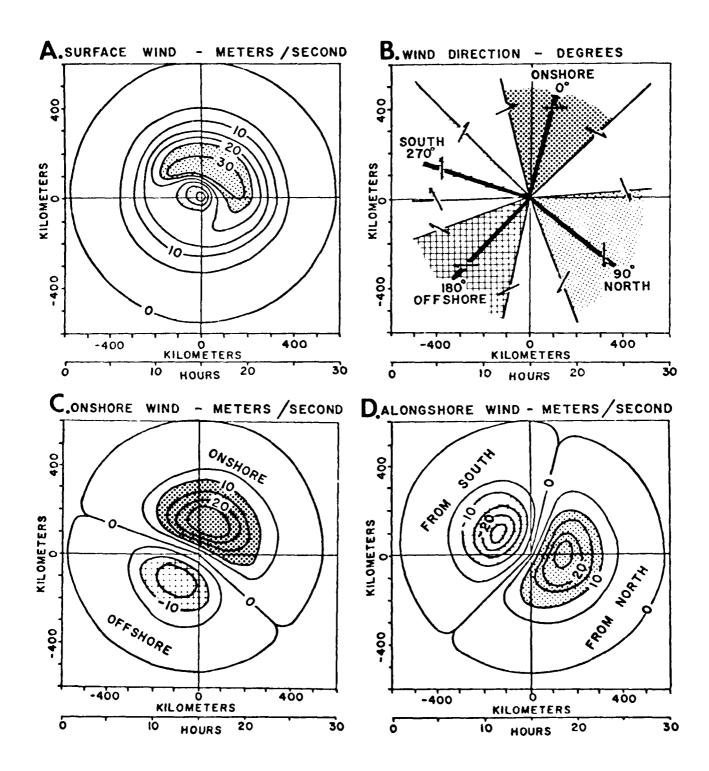
When the storm passes to the north of the shore site, the wind blowing around the storm center in a counterclockwise direction is generally onshore at the shore site (Figure 18B). The contour lines indicating wind direction radiate out from the center of the diagram and the arrows along each contour line point in the direction the wind is blowing along that line (Figure 18B). The dark lines on the diagram indicate the major wind directions with onshore winds = 0° , north winds = 90° , offshore winds = 180° and south winds = 270°. When a storm follows a path 200 kilometers north of the study site, the wind direction at the study site starts out from the south (274°) , slowly shifts over to onshore (0°) , and ends up out of the northwest (52°). When a storm passes 200 kilometers to the south of the shore site, the wind starts off blowing offshore (209) and then shifts over to the north (71°). Different patterns are used to show area where the winds are blowing generally onshore, from the north, offshore are from the south. These patterns related to similar areas for the onshore and alongshore components of the wind (Figures 18C and D).

Diagrams for the onshore and alongshore components of the surface wind speed are given in Figures 18C and D. The onshore component of the surface wind is obtained by taking the cosine of the

wind direction times the wind speed, and the alongshore component is produced by taking the sine of the wind direction times the wind speed. Storms which follow a path to the north of the shore site generally have a strong onshore wind component, while those following a path to the south of the shore site are predominately offshore (Figure 18C). The dividing line between the onshore and offshore components of the wind follows north-south wind direction lines (0° to 180°)(Figure 18B). The onshore wind reaches a maximum of 28.2 meters/second along a storm path 100 kilometers north of the shore site. The offshore wind reaches 18.6 meters/second when the storm passes 100 kilometers south of the shore site and before its center moves across the coast. The onshore wind speed is greater than the offshore wind speed because the friction is less when the wind is blowing over the water.

The time distance diagram for alongshore wind component indicates that the winds are from the south as the storm approaches the coast and shift over to the north after the storm passes (Figure 18D). To the north of the shore site, the shift in the alongshore component from south to north takes place after the storm has passed over the coast, but to the south, the shift takes place before the storm reaches the coast. The boundary line between the north and south components of the longshore wind (Figure 18D) follows the onshore-offshore line $(90^{\circ}-270^{\circ})$ in the wind direction diagram (Figure 18B). The reversal from south to north wind is very abrupt near the center of the storm, and more gradual near its margin. When the storm passes to the north, the low in barometric pressure reaches a minimum when the storm passes over the coast (Figure 17), but the shift in wind direction from south to north does not take place until several hours after the low has passed. This lag in wind direction reversal behind the low in pressure was observed during storms at Holland and Stevensville, which passed to the north of the study areas (Fox and Davis, 1970 and 1971a). Also, the maximum wind speeds were observed after the lows had passed and the wind shifted over to the north.

Figure 18. Time-distance plots of A - surface wind speed, B - wind direction, C - onshore wind, and D - alongshore wind in a circular storm.



Wave Period and Breaker Height

The diagrams for wave period and breaker height are quite similar (Figures 21A and B), and both resemble the plots for surface wind and onshore wind (Figures 18A and C). Three profiles were plotted across the breaker height diagram to show what the height would be at the shore site as the storm moves across (Figure 19). If the storm moves 200 kilometers north of the shore site, the breaker height will reach 5.30 meters 3 hours after the storm crosses the coast. If the storm passes directly over the shore site, the breaker height will reach .86 meters as the storm approaches, drop down as the center of the storm passes, then reaches 3.63 meters 5 hours after the center passes over the coast. The decrease in wave height as the storm center passes directly over the shore site corresponds to the zero wind velocity at the center of the storm. While wind velocity may drop to zero at the storm center, the zero wave height is probably an artifact of the computer model and does not occur in nature. Residual waves would most likely remain in the area and could be built model if so desired.

The surface wind speed is not used directly for determining breaker height and wave period, because strong onshore wind is effective in generating waves which will reach the coast, and a strong offshore wind tends to subdue existing waves. On the Texas coast during studies made at Mustang Island, offshore wind was accompanied by a sharp drop in breaker height (Davis and Fox, 1972c). On Lake Michigan, where a single storm system was studied as it moved offshore at Zion, Illinois and onshore at South Haven, Michigan, breaker height was over 2 times as great where wind was blowing onshore than where it was blowing offshore (Davis and Fox, 1974b). Therefore, the effective wind speed was used in determining wave height and period in place of the surface wind. For an onshore wind, the effective wind is equal to the onshore wind. However, for an offshore wind, the effective wind is about one third of the surface wind speed. For a wind blowing along the shore, the effective wind speed is two thirds of the surface wind speed. A cosine transformation was used to produce a smooth gradient in effective wind from onshore through alongshore to offshore.

The plot for wave period closely resembles the plot for breaker height with the maxima to the north of the shore site and displaced landward of the shoreline (Figure 21A). The maximum wave period of 9.5 seconds occurs at the same time as the maximum breaker height, 3 hours after the storm has passed over the coast. The plot for wave period has a broad relatively flat area surrounding the maximum, while the plot for breaker height is much steeper, reaching a peak and rapidly dropping off after the peak has been passed. The wave height and periods forecast in the model correspond closely to those encountered on the Oregon coast during November 1973 (Fox and Davis, 1974).

Breaker Angle and Longshore Current Velocity

The plots of breaker angle and longshore current velocity (Figures 21C and D) are similar in many respects to the plot for alongshore wind (Figure 18D). The boundary line which separates the north and south components of the wind is the same as the boundary which separates the north and south breaker angles and longshore currents. For field studies conducted at Holland and Stevensville, Michigan, there was also a close correspondence between the longshore component of the wind and longshore current velocity (Fox and Davis, 1970a and b, and 1971a).

The breaker angle is defined as the acute angle between the wave crest and the shoreline as the wave passes over the nearshore bar. In deep water the wave direction is roughly parallel to the wind direction, and the wave crests are about normal to the wind. The dominate wind direction is often used in wave refraction computer programs to determine the deep water wave angle (Dobson, 1967). As a wave enters shallow water, the celerity decreases and the wave crest is refracted so that it becomes closer to parallel to the beach. Snell's Law of geometrical optics is used for computing the refraction coefficient and breaker angle in the surf zone.

In the breaker angle diagram, the area to the left of the zero line has breaker angles open to the north, and to the right of the zero line, the breaker angles are open to the south (Figure 21C). The largest breaker angles are about 32 degrees when the wind is blowing directly out of the north or the south. As the wind direction swings around from alongshore to onshore or offshore, the breaker angles decreases from 30 degrees to zero (Figures 18B and 21C). When the wind is blowing directly onshore or offshore, the breakers are parallel to the beach and breaker angle is zero.

The plot for longshore current velocity is very similar to the plot for the alongshore component of the wind (Figure 18D and 21D). Three profiles are plotted when the storm passes 200 kilometers north, over the shore site and 200 kilometers south (Figure 20). When the storm path is 200 kilometers north, the longshore current velocity reaches 97.4 centimeters/second to the north, reverses direction after the low has passed, and increases to 81.6 centimeters/second to the south. When the storm passes directly over the study site, the current reverses from 63.0 centimeters/second to the north to 129.7 centimeters/second to the south. However, when the storm path is 200 kilometers south of the shore site, the current to the north is only 4.8 centimeters/second and the southward current is 58.1 centimeters/second. When the storm passes to the south, the reversal in current direction takes place before the low in barometric pressure passes the shore. The maximum long-

Figure 19. Time-distance plot of breaker height and three profiles of breaker height in a circular storm.

Figure 20. Time-distance plot of longshore current and three profiles of longshore current in a circular storm.

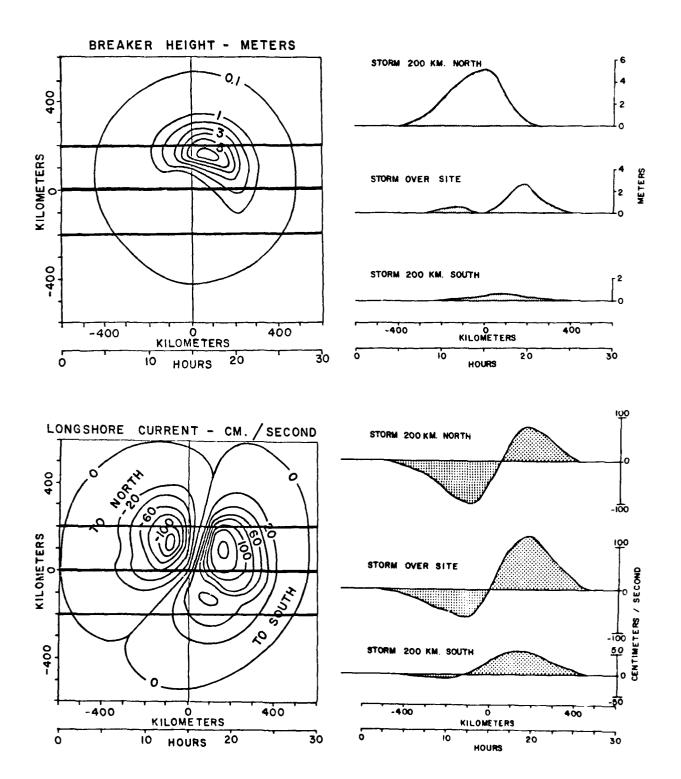
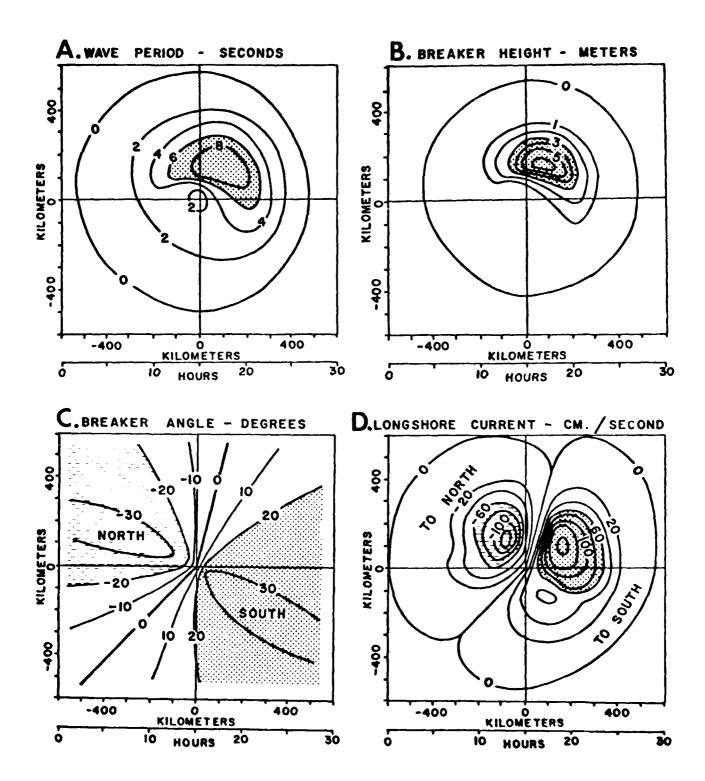


Figure 21. Time-distance plot of A - wave period, B - breaker height, C - breaker angle, and D - longshore current in a circular storm.



shore current was 144 centimeters/second to the south 4 hours after the storm passed over the shoreline on a path 100 kilometers north of the shore site (Figure 20).

The longshore current velocity is a function of nearshore slope, breaker height and breaker angle (Longuet-Higgins, 1970). The influence of both breaker angle and breaker height can be seen in the plot for longshore current velocity (Figures 19B, C and D). Other longshore current equations were tested which gave similar patterns, but different absolute velocities.

The circular storm test illustrates the general patterns which emerge in barometric pressure, wind, waves and longshore currents as a storm passes over a coast. If the shape and path of the storm, and the orientation of the shoreline are held constant while the size or intensity of the storm are varied, the same patterns will persist, but the absolute values will change for each of the variables. However, if the shape or path of the storm are changed, the patterns as well as the absolute values will change for each of the variables. In the next section an elliptical storm is used to demonstrate the effect which a change in storm shape would have on the wind, waves and currents.

Elliptical Storm Test

With the coastal storm model, it is possible to vary the size and shape of the storm while holding the intensity constant. The elliptical storm test provides a good example of an oval shaped storm which has its long axis extending to the north-northeast (Figure 22). In both the circular and elliptical tests, the shoreline orientation and nearshore bottom slope are the same.

In the circular and the elliptical storms, the minimum barometric pressure at the center is 1000 millibars, and the pressure at the largest encircling isobar is 1020 millibars (Figures 16 and 22). In the circular storm, the major and minor axes are the same, 300 kilometers. In the elliptical storm, however, the major axis (500 kilometers) is twice the length of the minor axis (250 kilometers). The major axis in the elliptical storm is oriented 30° east of north. Therefore, a low pressure trough extends in a north-northeast direction with the lowest value at the storm center.

The time-distance plot of barometric pressure is identical to a weather map made when the storm center is over the shore site (Figure 22). When the storm track is located to the south of the shore site, the low pressure trough reaches the shore site before the low pressure center passes over the coast. However, when the storm track is to the north of the shore site, the low pressure center reaches the coast before the trough passes over the shore site. Therefore, the long axis of the storm marks the time when the low pressure trough passes over the shoreline.

Although the range in barometric pressure is the same in the circular storm and the elliptical storm, the pressure gradient is steeper in the constricted part of the elliptical storm. The pressure gradient is a function of the size of the storm and the range in barometric pressure. In the circular storm, both the major and minor axis have lengths of 300 kilometers, and therefore the pressure gradient is equal on all side of the storm. In the elliptical storm, the major axis is 500 kilometers and the minor axis is 250 kilometers. Therefore, along the minor axis the pressure gradient is steeper, while it is more gentle along the major axis.

The time-distance plot of surface wind speed has an elliptical shape with the high winds concentrated on the right side of the diagram (Figure 23A). The high wind speeds are a function of the steeper pressure gradient along the minor axis and differences in surface friction over land and sea. With the higher pressure gradient, the wind speed reaches 36.8 meters/second in the elliptical storm, while in the circular storm, it only reaches 30.9 meters/second. The winds greater than 20 meters/second are split into

two areas in the elliptical storm, a major area down the right side of the storm, and a minor area in the northwest quadrant. At the center of the storm along the major axis, the surface wind speed drops down to zero.

The pattern for wind direction in the elliptical storm is significantly different from the pattern in the circular storm (Figures 18B and 23B). In the circular storm, the wind direction contours radiate out from the center and are rotated in a clockwise direction from 14° to 45°. The zero wind direction indicates onshore wind and 180° is an offshore wind. The elliptical storm is constructed along the minor axis and extended along the major axis. Therefore, the wind direction contours are gathered around the major axis which is pointed 30° east of north. Similarly, the contours are spread out from the minor axis. In the elliptical storm the zero contour extends 45° east of north, and the 180° contour extends 67° west of south. The 270° contour indicating south winds points 17° east of north and the 90° contour for north winds extends 25° west of south.

The difference in wind direction pattern results in major changes in the onshore and alongshore wind patterns in the elliptical storm (Figures 23C and D). In the circular storm, the boundary line between onshore and offshore winds runs in a generally east-west direction with onshore winds when the storm track is to the north and offshore winds when the storm track is to the south of the study site. In the elliptical storm, on the other hand, the boundary between onshore and offshore winds has shifted so that it runs generally north-south with onshore winds to the east and offshore winds to the west (Figure 23C). As the storm approaches the coast the winds are offshore, and after the storm has passed over the shoreline the winds shift to onshore. When the storm passes to the south of the shore site, the shift from offshore to onshore winds takes place shortly after the low pressure trough passes over, but when the storm track is to the north, the shift in wind direction shortly preceeds the low pressure trough. The maximum offshore wind is 21.4 meters/second as the storm approaches, and the maximum onshore wind is 30.2 meters/second after the storm has passed over the coast.

The boundary for the alongshore wind which separates the north wind from the south wind extends generally in a northwest-southeast direction (Figure 23D). It resembles the alongshore wind diagram for the circular storm, but the axes are rotated about 30° in a clockwise direction (Figures 18D and 23D). The maximum south wind of 21.1 meters/second occurs when the storm track is to the north of the shore site and after the storm has passed over the coast. For the north wind, the maximum of 28.1 meters/second occurs along a storm track 200 kilometers south of the shore site after the storm has passed the coast.

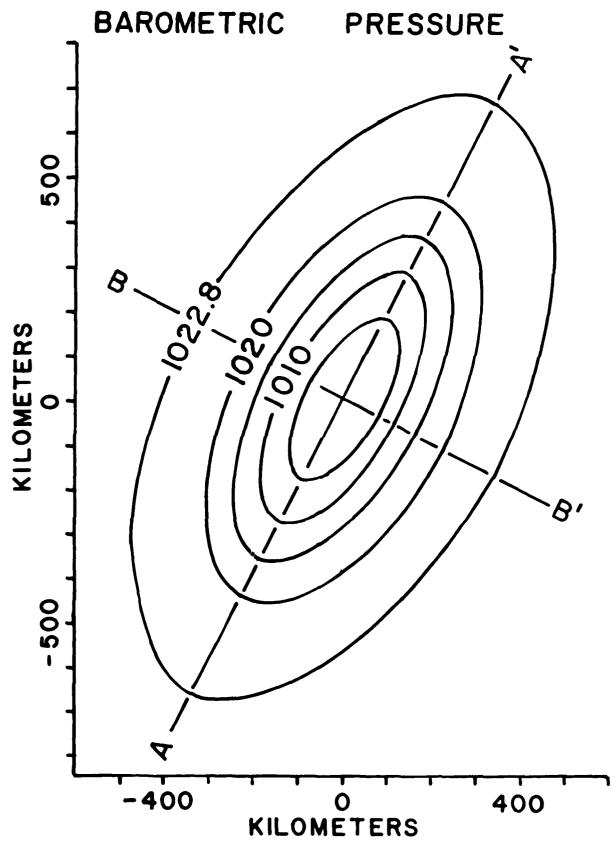


Figure 22. Time-distance plot of barometric pressure in an elliptical storm.

Figure 23. Time-distance plot of A - surface wind speed, B - wind direction, C - onshore wind and D - alongshore wind in an elliptical storm.

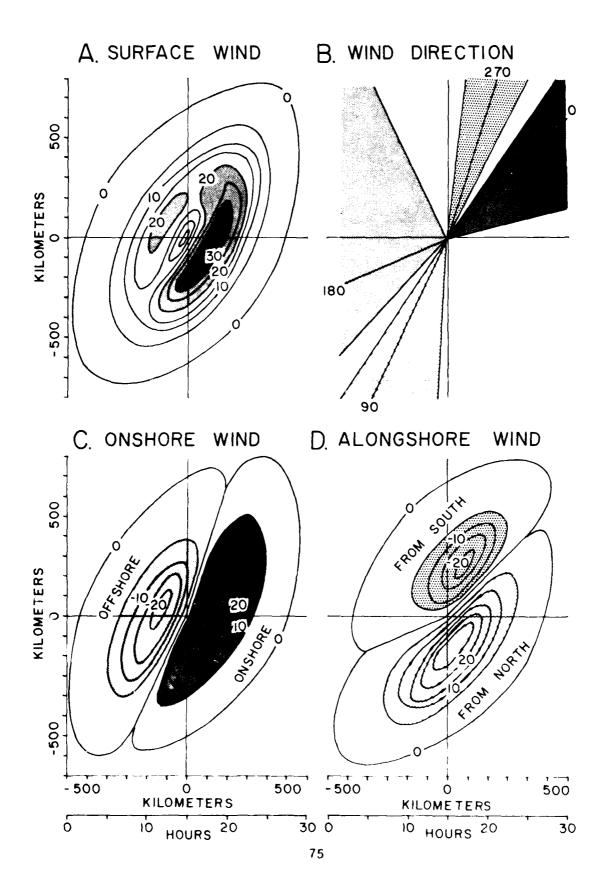
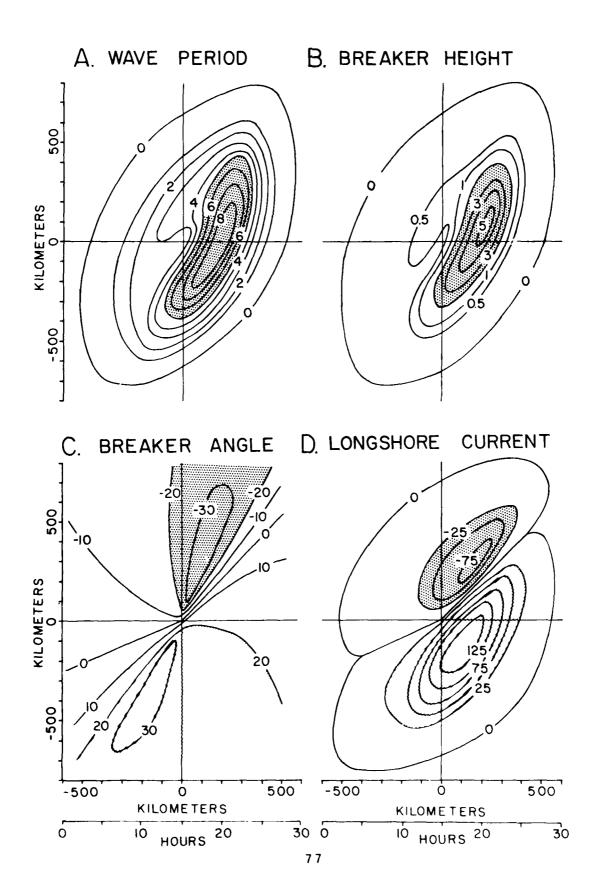


Figure 24. Time-distance plot of A - wave period, B - breaker height, C - breaker angle and D - longshore current in an elliptical storm.



The plots for wave period and wave height in the elliptical storm are very similar and closely resemble the plots for wind speed and onshore wind (Figures 23A and C, and 24A and B). Both wave period and breaker height reach their maxima after the storm trough has passed over the coast and the wind has shifted from offshore to onshore. In the elliptical storm, the maximum wave period is 8.9 seconds and the greatest breaker height is 5.1 meters.

The greatest breaker angles for the elliptical storm occur when the wind is blowing out of the south (270°) or the north (90°) (Figures 22B and 23C). When the storm track is to the north of the shore site, the greatest northerly breaker angles (32.3°) are present just before the low pressure trough passes over the coast. However, when the storm track is to the south of the study site, the largest southerly breaker angles (32.4°) occur just after the storm trough has passed the shore.

The plot for longshore current in an elliptical storm is very similar to the plot for alongshore wind (Figures 23D and 24D). The boundaries between north and south winds and north and south currents follow the same line and the maxima are in the same position. The maximum northward flowing longshore current (44 centimeters/second) occurs when the storm is on a track 300 kilometers north of the shore site, while the maximum southward flowing current (143 centimeters/second) occurs on a storm track 100 kilometers south of the shore site.

In summarizing the comparison between a circular storm and an elliptical storm of the same intensity, the differences in barometric pressure and wind direction influence the other environmental parameters. In the circular storm, the highest surface winds are found when the storm track is to the north of the shore site while in the elliptical storm, the maximum winds occur in a north-northeast trending zone to the right of the low pressure trough. The boundary between offshore and onshore winds lies generally east-west for the circular storm and north-south for the elliptical storm. For alongshore winds, the boundary between north and south winds is rotated about 30° in a clockwise direction in the elliptical storm. In both the circular and elliptical storm, the wind direction contours radiate out from the center and are rotated clockwise due to surface friction, but in the elliptical storm, the contours are gathered around the major axis. Wave period and breaker height in the circular storm resemble wind speed and form pods to the north of the shore site, while in the elliptical storm, period and breaker height form linear trends to the east of the shore site. For both storms, the longshore current follows the same pattern as the alongshore wind.

CONCLUSIONS

A mathematical model has been developed and programmed for a computer to forecast barometric pressure, wind, waves and longshore currents during passage of a storm across a coastal site. The following set of conclusions can be drawn from the coastal storm model.

- 1. The shape of a coastal storm can be approximated with an elliptical model by specifying the lengths of the major and minor half axes and the orientation of the major axis.
- 2. The barometric pressure profiles along the major and minor axes of the ellipse are represented by a series of inverted normal curves.
- 3. Geostrophic wind speed and direction at any point on the earth's sufrace under a storm are computed from the latitude and barometric pressure gradient.
- 4. Geostrophic wind speed and direction are used to compute surface wind speed and direction over land or sea.
- 5. Wave period, height and direction are calculated from the wind speed, fetch and duration as the storm passes over the coast.
- 6. Longshore current speed and direction are computed from wave height, period and direction, and nearshore bottom slope.
- 7. Wave and current hindcast data can be used to test the model when the size, shape, intensity and track of the storm are known.
- 8. If the storm azimuth and velocity are assumed to be constant, forecasts of wind, wave and longshore currents can be made for a coastal storm.

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APPENDIX A. LOASTAL STURM PRUGRAMS

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// JOB
// FOR
*IOCS(CARD+1132 PRINTEK)
  *COMMON U(130) +V(130)
                 COMMON U(301+V(130)

DIMENSION TITLE(201+DAY(5)

DIMENSION A1(21+A2(21+A3(21+A4(21+B1(2)+B2(21+C1(21+C2(21+C4(2)+C4(2)+C4(2)+B1(2)+B2(21+C1(21+C2(21+C4(2)+C4(2)+B1(2)+B2(21+C1(21+C2(2)+C4(2)+C4(2)+B1(2)+B2(21+C1(21+C2(2)+C4(2)+C4(2)+B1(2)+B2(21+C1(2)+C4(2)+C4(2)+B1(2)+B2(21+C4(2)+C4(2)+C4(2)+B1(2)+B2(21+C4(2)+C4(2)+C4(2)+C4(2)+B1(2)+B2(21+C4(2)+C4(2)+C4(2)+C4(2)+B1(2)+B2(21+C4(2)+C4(2)+C4(2)+B1(2)+B2(21+C4(2)+C4(2)+C4(2)+B1(2)+B2(21+C4(2)+C4(2)+C4(2)+C4(2)+B1(2)+B2(2)+C4(2)+C4(2)+B1(2)+B2(2)+C4(2)+C4(2)+B1(2)+B2(2)+C4(2)+C4(2)+B1(2)+B2(2)+C4(2)+C4(2)+B1(2)+B2(2)+C4(2)+C4(2)+B1(2)+B2(2)+C4(2)+C4(2)+B1(2)+B2(2)+C4(2)+C4(2)+C4(2)+B1(2)+B2(2)+C4(2)+C4(2)+C4(2)+C4(2)+B1(2)+B2(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C4(2)+C
                     CARD 1 - TITLE
                      CARD 2 - STARTING DATE: TIME AND INPUT-OUTPUT OPTIONS
                                              COLS. 1-2 ISTRT- STARTING HOUR COLS. 3-22 DAY - STARTING DATE
                                             COL. 23 INAUT- INPUT OPTION

O • METRIC UNITS

1 • NAUTICAL MILES AND FEET
                                             COL. 24 NAUT - OUTPUT OPTION
                                                                                                                0 = METRIC UNITS
1 = NAUTICAL MILES. FEET AND KNOTS
                     CARD 3 - STORM PARAMETERS
                                                                                   STORM INPUT UPTIONS - BLE CAND 6

1 - HINDCASTING - STORM POSITION AT 6 HOUR INTE

2 - FORECASTING - INITIAL POSITION, AZIMUTH AND

VELOCITY

TIDE PREDICTION DOTION

0 - TIDE PREDICTION NUT INCLUDED - UMIT CARD 4

1 - TIDE PREDICTION INCLUDED - SEE CARD ~
                                              COL. 2
                                                                                      LONGSHORE CURRENT EQUATION CPTION

1 - FOX AND DAVIS: 1972

2 - LONDUET-HISDINS: 1970

3 - C.E.R.C..: 197:

4 - KOMAR AND INMAN: 1970
COLS, 4-6 NX - NUMBER OF STURM POSITIONS
SIX HOUR INTERVALS FOR HORCASTING
ONE HOUR INTERVALS FOR FORECASTING
                                             COLS. 7-12 BRFCH- AVERAGE BASIN FETCH IN KM. (NAUT. MI.)
COLS. 13-17 TINT - TIME INTERVAL BETWEEN STORM PUSITIONS
NORMAL SETTING IS 1.0 HOURS
                                             CCLS. 1H-24 PMIN - MINIMUM HARCMETRIC PRESSURE IN MILLIBARS CCLS. 25-31 PMARK- PRESSURE AT LARGEST ENCIRCLING ISUBAR CCLS. 32-36 SCAT - CATITUDE AT 5 ORE SITE
                                            COLS. 37-42 AH = MAJOR HALF-AXIS (E)FECTIVE CONG HADIUS) 
COLS. 43-45 BR = MINOR HALF-AXIS (E)FECTIVE SHORT HADIUS) 
- MINOR HALF-AXIS (E)FECTIVE SHORT HADIUS) 
- MINOR HALF-AXIS (E)FECTIVE CONG HADIO HALF AXIS PLUS OF MINOR 90 DEGREES FROM NUMTH.
                    CARD 4 - TIDE FREDICTION - UPTION FROM CARD 3
                                            COLD. (-5 ST - SPRING TIDE RANGE IN FEET COLD. (-1) TN - NUMBER THOS HANGE IN FEET COLD. 11-14 TLAY - NUMBER OF LASS FROM LAST SPRING TIDE COLD. 15-12 THE - HOUR OF LAST HIGH SPRING TIDE
                                            CLEDE C. TO FINE - TIDE FORM NUMBER

DEFINE 25 - SEMIDIUMNAL TIDE

EN TO CEN - MIRED SEMIDIUMNAL TIDE

CONTROL OF MIRED COMPACTIVE

GREATER THAN SE - DIUMNAL TIDE
                                           CLUB 26-92 SCHOL- NORMHORE BOTTOM SCOPE AT LOW TIDE COCS. 13-45 SCHOL- NEARCH WE HAT M SCHE AT HIGH TIDE COCS. 4 -42 TOPAN- NEAR TOE CEREL IN FEET.
                     LANGER - SMOKE SATE OF ATTON
                 Fig. which \gamma = a_{\rm CM} as the a shaw fostings at a word interval at many that
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COLS. 8-14 AZI - STORM AZIMUTH - CLOCKWISE FROM NURTH
COLS. 15-21 VIII - INITIAL X-COURDINATE IN NAUTICAL MILES
COLS. 22-28 VIII - INITIAL Y-COURDINATE IN NAUTICAL MILES
CNK+1.85319
                            CKN=.53961
CFM=.3048
CMF=3.2806
                            CKTS=1.9425
CCMF #.032808
                          CUCUL = +737561
CPG+33+8639
   c
1 CNK=1.0
             (FM=1.0
2 1F1NAUT) 11-11-12
11 CKN=1.0
CMF=1.0
CKTS=1.0
(CMF=1.0
CJOUL=1.0
CPU+16

12 IF .15TR1-EU-0+01 GO TO 1000

*RITE:3+0181[STR1+504+

916 FURMAT(1X+'NUN HIGHNS AT HOUR '+12+' ON '+544)

*HEAD(2+901]INOPT+1FTID+USCOP+NX+BNFC(+TINT+PMIN+PMAX*+SLAT+AR1+BK1

> 1544
                           CPU=1.0
         1 +EAZ
901 FORMAI(311+13+66+0+F5+1+2F7+1+F0+1+366+0)
HNECH+BNEC1+ZNK
                       SHEBRISON.
c
                       BAHOMETRIC PRESSURE AT LANGEST ENCINCUING ISOBAR IS ASSUMED TO BE AT THE STURM. TO FIND STANDARD DEVIATIONS FROM THE CENTER OF THE STURM. TO FIND THE AUTUAL STURM SIZE, THE MADUR AND MINOR ARES ARE MULTIPLIED BY AND THE MATHEM PRESSURE IS MULTIPLIED BY LANG TO DETERMINE AT THE MARGIN OF THE STORM.
       16-16-10 - 5-5-4

+ MLAULE-970- SIFTI-TDAY-THRIPTS-SCPEO-S.PHI-TMEAI

920 FORMATISH-1-38-1-23

17-51-1-1-1

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BZ ATTTETTHER
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     AR AND BY ARE HADLE OF THE STORM OUT TO THE STANDARD DEVIATIONS... MULTIPLY BY 1.5 TO YIELD THE FULL LENGTHS.
      HEADEBH

HEADEZERO - UND EVUCCESHAZESUMHEETSUNDEBLOR

93 FORMATINETERSTEELT

I MAREDMAXENCOM
                      PHINOPHINEBCOR
                      JUDENJE /CPCNK

/_ /C+/L /CPCNK

(F (IN)/P(=1) 1000+10+30
                     CALL SUBROUTINES FOR MINDCASTING AND FURECASTING STORM PLSITIONS
        1. TALL HINDONESA)
33-73 NO
37 CALL FURE(NESANN
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40 DO 42 I=1+NX
U(1)=U(1)=CNK
42 V([]=V,]]=CNK
                                                         ULOC=ULOC/A
VLOC=VLOC/A
U1=U(1)=A++005
                                                            V1=V(1)*A+.005
                                                                         WRITE SHORE AND TIDE DATA
                                                         UL1=UL+CKN++05
                                            ULIBURECH++05

BRC1-BRCH-BCKN

#RITEE13-970JUL1-VL1-A1(NN)+A2(NN)+A3(NN)+A4(NN)+SLAT+SHAZ+SLOPE+

1 BRFC1-A1(NN)+A2(NN)+A3(NN)+A+(NN)

FORMATI' SHOKE - POSITION COONDINATES - X *'+F7-1+' Y *'+F7-1+

1 A44

2 *'-FE-0+' DEUNEE5'/9X+'NEARSHORE SLOPE *'+F6-0+' AVERAGE FEICH
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UPT=U111=A+CK+++5
VPT=V111=A+CK+++5
#SRF=#SURF=CKTS
UN=ONSH=CKTS
AL=ALSH=CKTS
EF=EF#ND=CKTS
                                                HII = HE IGT = CMF
HBI = HB = CMF
TID1 = TIDX = CMF
VI = VLS = CCMF
                     IMOUR=IMOUN-11-IXITINT)

IF (IMOUR-67-24) IMOUR=IMOUR-24

50 CONTINUE

ELSCT=ENMEP

ESUM=ESUM=CUOUL

EN=EN=CUOUL

MRITE(3-5-80) ESUM+CI(NN)+C2(NN)+C3(NN)+ELSCT+C1(NN)+C2(NN)+C3(NN)

5-49 FORMAT(IX-//*1X+15-18AVE ENERGY IN THE BREAKER ZONE ="+E10-3+3A4//

IX+15-+10-11 LONG-SHORE CURRENT ENERGY ="+E10-3+3A4//

WRITE(3+5421 EP+C1(NN)+C2(NN)+C3(NN)+ENSC(INN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3(NN)+C3
     1344/71x+TE
GC TO 8080
1000 CALL EXIT
END
// OUP
*DELETE
*STORE #S U
                                                                             STRMX
  // JOB
// FOR
*IOCSICARD:1132 PRINTER)
*ONE WORD :NTEGERS
*LIST SOURCE PROGRAM
SUBROUTINE FORE(NA:A:NAUT)
COMMON U(1301:V:133)
                                                                                                                                                                                                                                                                                                                     FORE
                                                                                                                                                                                                                                                                                                                                                                                       3-86-001
                                        FORECASTING - COMPUTE STORM POSITIONS FROM INITIAL POSITION. VELOCITY AND AZIMUTH
                 TINT+1.0
RAD=57.2958
READ(2+903) SVFL+AZI+U(1)+V(
903 FORMAT(4F7-0)
 READ(2+003) SVEL+AZI+U(1)+V(

903 FORMAT(4F7+0)

IF(NAUT) 1:1-2

1 WRITE(3:910) SVEL+AZI

910 FORMAT(9X-'STORM VELOCITY *'+F4+0+' KILOMETERS/HQUR'/

1 9X+'STORM AZIMUTH *'+F4+0/'

90 TO 3

2 WRITE(3:911) SVEL+AZI

911 FORMAT(9X-'STORM VELOCITY *'+F4+0+' KNOTS'/

1 9X+'STORM AZIMUTH *'+F4+0/'

5 CONTINUE

U(1)+U(1)/A

V(1)+U(1)/A

DIST+SVEL+TINT/A

AZM+90-AZI

IF(AZM) 31-32+32

31 AZM+AZM+360

32 DO 35 I=2+XA

U(1)+U(1-1)+UIST+COS(AZM/RAD)

V(1)+V(1-1)+DIST+SIN(AZM/RAD)

35 CONTINUE

RETURN

END

// DUP

**DELETE FORE

**STORE #5 UA FORE
                                                                    FORE
W5 UA FORE
   *STORE
// JOB
// FOR
*IDCS(CARD+1132 PRINTER)
*IDCS(CARD+1132 PRINTER)
*ONE WORD INTEGERS
*LIST SOURCE PROGNAM
SUBPOLITINE MINDINAVAL
COMMON ULIBOTAVISO)
DIMFNSION X(3074YISO)
                                                                                                                                                                                                                                                                                                                  HIND
                                                                                                                                                                                                                                                                                                                                                                      3-86-001
                                      HINDCASTING - COMPUTE STORM POSITIONS AT 1 HOUR INTERVALS FHOM 6 HOUR POSITIONS ON WEATHER MAPS
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V(11)=Y(1)
              V(I)=Y(I)
DIFU=(X(I+1)-X(I))/6.
DIFV=(Y(I+1)-Y(I))/6.
DO 20 J=1.5
M=I(+J
U(M)=U(M-I)+DIFU
       V(MI=V(M-1)+01FV
20 CONTINUE
              II=NX#6-5
U(II)=X(NX)
V(II)=Y(NX)
NX=II
              RETURN
              END
  // DUP
                           HIND
WS UA HIND
  *STORE
 // JOB
// FOR
*IOCS(CARD+1132 PRINTER)
*ONE WORD INTEGERS
*LIST SOURCE PROGRAM
SUBNOUTINE LOCATIUST+VST+ULOC+VLOC+SHAZ,X+Y+EAZ+P+J)
                                                                                                                               3-86-001
 0000
             U AND V COORDINATES ARE READ INTO THE PROGRAM IN KILOMETERS AND CONVERTED TO UNITS PROPURTIONAL TO THE MAJOR AXIS. THE U AND V COORDINATE SYSTEM IS A RECTANSULAR GRID WITH U POINTING LAST, V POINTING NORTH. AND THE ORIGIN LOCATED TO THE SOUTHWEST OF TH SHORE LOCATION.
             SUBROUTINE LOCAT IS USED TO ESTABLISH THE X AND Y COORDINATE SYSTEM WITH THE CHICIN AT THE CENTER OF THE STORM AND POSITIVE Y PUINTING OFFSHURE, POSITIVE Y POINTING OFFSHURE, POSITIVE X TO THE RIGHT FACING CHARGES, AND NEGATIVE X TO THE LEFT FACING
              ASHUR:
              RADA57.2958
             RADASTALSHB
URUST-ULCC
V=7CCC-VST
J=5QR31U=P2+V+P23
CACL AACTA ANUUSV)
AFATW-SHAZ/RAC
AFATW-SHAZ/RAC
AFATW-SHAZ/RAC
LEREAZE90-55HAZ
LEREAZE90-55HAZ
              AZ=A+D1F/RAD
P=-Z*CCS(AZ)
              RETURN
              END
 // DUP
*DELETE
*STORE
                       LOCAT
 // JOB T
 // JUG CSTRM 2
*LOCALCSTRM.LOCAT.WIND.DECAY.ETIME.FETCH.WAVES.SUHF.ETRUY.TIDE.HIND.
*LUCALCSTRM.LOCAT
*LUCALFOREC
HOLLAND: MICHIGAN
1JULY 7: 1970
101 24 150: 1:0
1000: 0: 90:
                                                                  42.
                                          •033
                                 90.
// JOB
// FOR
*IOCSICARD+1132 PRINTER)
*ONE WORD INTEGERS
*LIST SOURCE PROGRAM
                                                                                                          ELIPS 3-86-00:
              SUBROUTINE ELIPSIA. B. XI. YI. FMIN. PMAX. FI. ERAD. XA. YA. DZA.
             SUBROUTINE ELIPS IS USED TO DETERMINE THE WIND ANGLE AND PRESSURE GRADIENT AT ANY PULNT WITHIN AN ELLIPTICAL STORM.
             RAD=57+2958
R=8/A
AA=A/A
C
               A AND B ARE THE MAJER AND MINOR AXES OF THE STORM ELLIPSE POINT XI-YI LIES ON A SECOND ELLIPSE WITH AXES AT NO BI.
             A1-5-RT(x1--2--1--2/R--2)
              *D*#1+81**2/A1**2*#1
               <. IS THE INTERECTION OF THE A AXIS WITH A LINE NORMAL TO THE
TANGENT OF THE SECOND ELLIPSE AT 21.71.
             X2++3.(14-8A++2/AA++2)
Y2+5URT:8A++2+(1-x2++2/AA++2)
               R2 AND 12 ARE AT THE INTERSECTION OF THE LINE NORMAL TO THE TANGENT OF THE SECUND ELLIPSE. AND THE STORM ELLIPSE.
        16 (417 1+2+2
1 42+442
2 COMMINGE
```

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į

41 €

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D1*SURT((x1-x0)**2 + Y1**2)
D2*SURT((x2-x0)**2 + Y2**2)
D2A=D2**A
IF(D2A+x7T-A) U2A*A
                  ERAD=D1/D2
C1+3.#AO
PDIF=PMAX=PMIN
                  PINC=PC1F#EXP1-21##2/2.1
PI=FMAX-PINC
                  P1 IS THE BAROME'RIC PRESSURE AT XO COMPUTED ON A NORMAL CURVE
ALONG THE MAJOR AXIS. THIS PRESSURE VALUE IS USED IN DETERMINING
THE PRESSURE GRADIENT NURMAL TO THE ISOBAR AT POINT ALEXI.
                 1FEX1:Eu-0:31 X1#+.c.

REX1:Eu-0:31 X1#+.c.

XAEA1*#27:1

97-01*#17:1
                  COMPUTE THE TANGENT TO THE ELLIPSE AT X1 AND Y1 TO DETERMINE THE WIND DIRECT. No.
          [F(X].d1.0.0.4AND.Y1.01.0.0] 60 TO 11

[F(X].d1.0.3.AND.Y1.01.0.0] 60 TO 12

[F(X]..10.0.3.AND.Y1.01.0.0] 60 TO 12

[F(X]..10.0.3.AND.Y1.0.0] 60 TO 11

[CA--YA--YA--YA--YA--Y1.0.0] 60 TO 12

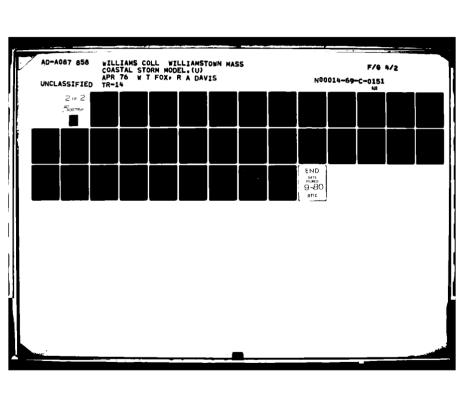
12 YA--YA--YA--YA--YA--YA--YA
          20 CUNTINUE
  *DELETE
*STURE
                              ELIPS
*5 UA ELIPS
SUBRUITINE WIND IS USED TO DETERMINE THE GEOSTUPHIC WIND SPEED AND THECTION FROM THE BARGMETRIC PRESSURE GRADIENT, THE SURFACE WIND SPEED AND SELECTION ARE DETERMINED UVER LAND AND SEA FROM COPPLETIONS ARE LITED TO THE GEOSTHUPHIC WIND, THE ONSHORE AZIMUTH IS USED TO THIS INSTITUTE AND ALBOMSHURE COMPONENTS OF THE WIND, THE CITED THE WIND STEEL THE WIND SPEED IS USED TO FIND THE WAVE HEIGHT, PERIOD, AND LENSHOUGH COMPONENTS VECTORING.
                 RAU=51.2718
11 1719.
U181A -0010729
                 THE BANDMETHIC PHYSLUNE SHABJENT ALONG A NORMAL CURVE IS USED TO CALLULATE THE SELSTHOPHIC WIND SPEED.
                  COR#2. #UMEGA#SIN(SEAT/RAD)
                 Z=3. MERAD
PDIF=PMAX= /1
                PDIFLOMAX--1
4 1MC=#21F*etAP(-2**2/2**
4.0Max-PINC
DETERMENT CODA
AGETREST CORP*(DPDN/5000*)
CALL ARCTA(ANG,XA,YA,
DIFFERZ**00--5042
1F3(1**31-2700) DIF=DIF=360*
AANG=*FAD**ANG-DIF
                  1F - MANG + E F + O + O ) - WANG + WANG + 360 + 1F - MANG + GT + 360 + WANG + WANG + 360 +
                 H AND BETA ARE CORRECTION FACTORS USED TO DETERMINE SURFACE WIND SPEED AND DIRECTION FROM GEOSTROPHIC WIND OVER LAND AND SEA
                  F - CANS - H = +270,40 AND BETA + 29
- + 56A - B + +223 65 AND BETA + 50
                  0. = UNOMINE - POSITIVE Y

*G. * A. Nash Mt. * MODITIVE X

UNO. * OFFUNDEE - NEGATIVE Y

CT. * A. NASHUME - NEGATIVE Y
                 ligati, telepatiki
Nasiena
                   F 4 ISLANT OFFICE & AND BETA VALUES FOR THE SEA ARE USED
         A U.SIN TRANSF, RMATION IS USED IN A 30 DEGREE HANSITION ZONE
N EACH STOLE OF THE SHORE INEX FOLL LAND AND BEAUTHER OF ONE AH
USED OUTSTOLE OF THANSITION ZONE OF PROPRETIONATE (ZEC) I ARE OUT
THE LATER AND SHALL PROFITIONS WITHOUT THANSITION ZONE
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29 SINS=SIN(SANG/RAD)

IF(SINS=-5) 35.30.30

30 B=.00019

BETA=29.

GO TO 50

35 IF(SINS+-5) 36.40.40

36 B=.000065

BETA=50.

GO TO $0

40 IF(SANG-90.) 41.41.42

41 SANA=SANG*3.

GO TO 48

42 IF(SANG-180.) 43.43.44

43 SANA=180.-3.*(180.~SANG)

GO TO 48

44 IF(SANG-170.) 45.45.46

45 SANA=180.-3.*(SANG-180.)

GO TO 48

46 SANA=360.-3.*(SANG-180.)

GO TO 48

48 SANA=360.-3.*(360.~SANG)

48 B=.0001*(1.275+0.625*SIN)
     48 B-000]=(1.275+0.625*SIN(SANA/RAD))
BETA-39.5-10.5*SIN(SANA/RAD)
50 BETR-8ETA/RAD
COTA-SIN(BETR)/COS(BETR)+COR/(B*COS(BETR))
ALPHA-ATAN(1.6/COTA)*RAD
IF(NX) 51.52.52
51 SHANG=KANG+90.-ALPHA
             SANG-SANG-ALPHA
     SANG=SANG=ALPHA
NX=1
GO TO 27
52 CONTINUE
VNCOR=COR=SIN(ALPHA/RAD)/(B+COS(BETR))
WSURF=VHCOR=WSGEO
000
              ONSHORE AND LONGSHORE COMPONENTS OF THE SURFACE WIND
           ONSH=WSURF *COS (SHANG/RAD)
            ALSH=WSURF#SIN(SHANG/RAD)
            EFWND IS THE EFFECTIVE WIND SPEED USED IN DETERMINING WAVES AN OFFSHORE WIND IS ASSUMED TO BE .333 TIMES AS EFFECTIVE IN GENERATING WAVES AS AN ONSHORE WIND.
            EFWND=WSJRF+(.6667+.3333+COS(SHANG/RAD))
      IF(SHANG) 55.60.60
55 SHANG=SHANG+360.
60 CONTINUE
      60 CONTINUE
IF(SHANG=360.) 70.70.65
65 SHANG=SHANG-360.
70 CONTINUE
RETURN
END
// DUP
                        WIND
+STORE
// JOB
// FOR
*IOCS(CARD*)132 PRINTER)
*ONE WORD INTEGERS
*LIST SOURCE PROGRAM
                                                                                                    DECAY
                                                                                                                          3-86-001
            SUBROUTINE DECAYITINT . PEROD . HEIGT!
            SUBROUTINE DECAY IS USED TO FIND THE DECAY IN WAVE HEIGHT AS THE WAVES MOVE AWAY FROM THE STORM CENTER.
C... FIND LOGARITHMIC ATTENUATION COEFFICIENT
            FREQ=1./PEROD
            ATCOF=10.00((FREQ-0.06)/(0.0324))-1.)
IF(ATCOF.GT-1.0) ATCOF=1.0
IF(ATCOF.LT.0.1) ATCOF=0.1
C... FIND PROPAGATION DISTANCE IN DEGREES C
            DIST=1.5607=PEROD+TINT+360./40074.
C ... FIND DECAYED WAVE HEIGHT
            HEIGT=HEIGT+(2.7183)++(-2.+(0.1151+ATCOF)+DIST)
            RETURN
END
 // DUP
DELETE
STORE
                        WS UN DECAY
// JOB
// FOR
+IOCS(CARD+1132 PRINTER)
                                                                                                     FTIME
                                                                                                                           3-86-001
 #ONE WORD INTEGERS
#LIST SOURCE PROGRAM
SUBROUTINE ETIME(WSNEW+OLDM+HOURS+DRATN)
0000
            SUBROUTINE ETIME IS USED TO DETERMINE THE EFFECTIVE FETCH AND DURATION FOR THE NEW WIND SPEED FROM THE LAST PREVIOUS WIND SPEED.
C ... CONVERT M TO FT AND M/SEC TO FT/SEC
```

```
c
               H=OLDH/0.3048
WS=WSNEW/0.3048
             FIND EFFECTIVE FETCH USING NEW WIND SPEED AND OLD WAVE HEIGHT
                U=((32.10H)/(0.283*WS**2.))
                SUM=U
DO 10 K=1:51:2
1=K+2
                SUM-SUM+UP# I / I
   10
                CONTINUE
                F=(w$802./32.1)#(SUM/0.0128)#02.38
              CONVERT FROM FT TO NM AND FT/SEC TO KTS
               F=F/6076.
WSKTS=WS#3600./6076.
  C ... FIND EFFECTIVE DURATION USING DERIVED EFFECTIVE FETCH
               EFDUR=(F/((WSKTS/10+)**0.72*10***0.3))**0.8
DRATN=HOURS+EFDUR
                RETURN
               END
   // DUP
                                             ETIME
                            WS UM ETIME
   *STORE
                                                                                                              FETCH
  // JOB
                                                                                                                                      3-86-001
   *IOCS(CARD+1132 PRINTER)
  *ONE WORD INTEGERS
*LIST SOURCE PROGRAM
SUBROUTINE FETCH(A+B+XA+YA+F)
               SUBROUTINE FETCH IS USED TO FIND THE EFFECTIVE FETCH AS A STORM MOVES ONSHORE OR OFFSHORE. THE FETCH LENGTH IS DETERMINED FROM THE FETCH AREA WHERE THE WIND IS BLOWING TOWARD THE SHORE SITE. THE EFFECTIVE FETCH AREA IS AN ELLIPSE WITHIN WHICH THE WIND IS BLOWING AT GREATER THAN HALF THE MAXIMUM WIND SPEED. THE MAXIMUM EFFECTIVE FETCH IS THE MAJOR AXIS OF THE ELLIPSE.
               R=(A+B)/2.
FMAX=0.5881*R
P1+3.14159
RAD=57.2958
              XAD=7:2758

X=XAPA

Y=YAPA

IF (ABS(X)-GE-(0-)-AND- ABS(X)-LT-(1-/3-PR)) GO TO 100

IF (ABS(X)-GE-(1-/3-PR)-AND- ABS(X)-LT-1D-646464PR)) GO TO 200

IF (ABS(X)-GE-(0-4446444PR)) GO TO 300

GO TO 400
 C ... X GREATER THAN OR EQUAL TO ZERO AND LESS THAN 1/3 R
             Y1=SGRT((1./3.*R)**2.+(1./2.*FMAX)**2.-X**2.)
Y2=SGRT((1./3.*R)**2.-X**2.)
Y3=0.0
Y4=-Y2
Y5=-SGRT(R**2.-X**2.)
IF(Y.GE**1) GO TO 110
IF(Y.GE**1) GO TO 110
IF(Y.LT**1.AMD.*Y.GE.Y2) GO TO 120
IF(Y.LT**2.AMD.*Y.GE.Y3) GO TO 130
IF(Y.LT**3.AMD.*Y.GE.Y4) GO TO 140
IF(Y.LT**3.AMD.*Y.GE.Y4) GO TO 150
IF(Y.LT**3.AMD.*Y.GE.Y5) GO TO 150
 100
            120
 130
              GO TO 500
RATIO-ABS(XI/SURT(X++2++4++2+)
             NOTE ... THIS IS THE SAME RATIO AS 130 SINCE YZ = Y4
             FMIN=RATIO=1./2.0FMAX
RANGE=1./2.0FMAX=FMIN
CURVE=:1.0COS((Y=Y4)/(Y3=Y4)0PI))/2.0
F=CUNVE@RANGE=FMIN
GJ 70 300
NATIJ=(Y=Y5)/(Y4=Y5)
150
            NATION(Y-Y-)/(Y-Y-)
CURVE-(11+COS(11-MATIO)-PI))/2.
F-CURVE-11-/2.#FMAX
UD TO 500
F-U-0
GO TO 500
100
C ... X GREATER THAN OR EQUAL TO 1/3 R AND LESS THAN 0.44444 R
             Y1=5QRT((1,-/2.0FMAX)002.0(1,-/3.0R)002.0-x002.)
Y20-8QRT(R002.0-x002.)
200
```

```
IF(Y+GT+Y1) GO TO 210
IF(Y+LE+Y1+AND+Y+GT+Y2) GO TO 220
IF(Y+LE+Y2) GO TO 230
GO TO 400
F=FMAX
 210
          T=FMAX

GO TO 500

RATIO=(Y1-Y)/(Y1-Y2)

CURVE=(1.+COS(RATIO+PI))/2.

F=FMAX+CURVE
          GO TO $00
F=0.0
GO TO $00
 230
C C... X GREATER THAN OR EQUAL TO 0.444444 R
 300
          Y1=0.0
          Y1=0.0

Y2=SQRT(R**2,-(0.4444*R)**2.)

IF(Y4CT.*Y1) GO TO 310

IF(Y4LE.*Y1.AND*Y5CT.*Y2) GO TO 320

IF(Y4LE.*Y2) GO TO 330

GO TO 400

F=FMAX
310
          F=FMAX
GO TO 500
RATIO=(Y1-Y)/(Y1=Y2)
CURVE=(1:+COS(RATIO+P1))/2:
F=FMAX*CURVE
          GO TO 500
F=0.0
GO TO 500
330
C ... ERROR CONDITION
 400
          WRITE(3.410)
 410
           FORMAT(1x+'LOGIC ERROR ENCOUNTERED IN SUBROUTINF FETCH...')
           F=0.0
          RETURN
IF(X+LT+(0+)) F=0+5+F
RETURN
          END
// DUP
*DELETE
*STORE
                    FETCH
WS UA FETCH
// JOB
// FOR
#IOCS(CARD:1132 PRINTER)
*ONE WORD INTEGERS
*LIST SOURCE PROGRAM
SUBROUTINE WAVES(SPEED:TIME:FETCH:EFETC:MEIGT:PEROD:IOUT)
                                                                                                        3-86-001
          SUBROUTINE WAVES IS USED TO DETERMINE WAVE PERIOD AND HEIGHT FROM WIND SPEED- DURATION AND FETCH-
C... UNIT CONVERSION PACKAGE...KM TO NM...M/SEC TO KNOTS
          FCHNM=0.5400#FETCH
SPDKT=1.9439#SPEED
C... EFFECTIVE FETCH PACKAGE
          EFETC=(SPDKT/10+)**0+72*10+**0+3*TIME**1+25
IF(FCHNM+LT+EFETC) EFETC=FCHNM
C ... UNIT CONVERSION PACKAGE...NM TO METERS...KNOTS TO M/SEC
          EFETC=EFETC/0.540041000.
SPDMS=SPDKT/1.9439
C... WAVE PARAMETER PACKAGE
          CONST= {9.8062*EFETC}/SPDMS**2.
HEIGT=\SPDMS**2.*0.283*TANH\(0.0125*CONST**0.42)\/9.8062
PEROD=\2.*3:\4\59*SPDMS*1.*20*TANH\(0.077*CONST**0.25)\/9.8062
EFETC=EFETC/\2000.*
           IF(100T.EQ.1) RETURN
IF(100T.EQ.2) GO TO 200
C ... ERROR CONDITION
          WRITE(3-100)
FORMAT(1X-'ILLEGAL OUTPUT OPTION IN SUBROUTINE WAVES')
100
          RETURN
C ... NON-METRIC OUTPUT OPTION
C
200 EFETC=EFETC=0.9400
HEIGT=MEIGT=3.2808
RETURN
END
// DUP
DELETE WAVE
                    WAVES
 STORE
// JOB
// FOR
**IOCS:(CARD+1132 PRINTER)
**ONE WORD INTEGERS
**LIST SOURCE PROGRAM
```

```
SUBROUTINE TIDE (ST.TN.FN.ISTRT.THR.TDAY.TIDX.SLTOP.SLBOT.SLOPE.I. 1 TMEAN)
           PERIODS FOR THE MAJOR DIURNAL AND SEMIDIURNAL TIDAL COMPONENTS
           TM2=12.42
           IMC=12-02
TS2=12-00
TK1=23-93
TO1=25-82
MR=24-9TDAY+FLOAT([STRT]+FLOAT([-1]-THR
ARG=6-2832*HR
           COMPUTE THE AMPLITUDE OF THE TIDAL COMPONENTS FROM THE SPRING AND NEAP TIDE RANGES AND THE TIDAL FORM NUMBER.
           FM2=(ST+TN)/(4.*(1.+FN))
$2=(ST-TN)/(4.*(1.+FN))
FK1=FN*FM2
          TIDE LEVEL IS THE SUM OF THE FOUR MAJOR TIDAL COMPONENTS AT EACH HOUR.
        TIDX = FM2*COS(ARG/TM2) + S2*COS(ARG/TS2)+ FK1*COS(ARG/TK1) + 1 O1*COS(ARG/TO1)
 0000
          BOTTOM SLOPE IS COMPUTED FROM THE TIDE LEVEL. AND THE SLOPE AT HIGH AND LOW TIDES.
          TLOC = (TIDX+ST/Z*) / ST
SLOPE = SLBOT + TLOC * (SLTOP-SLBOT)
TIDX=TIDX+TMEAN
 // DUP
 *DELETE
                   WS UA TIDE
// JOB
// FOR
// FOR
*IOCS(CARD+1132 PRINTER)
*ONE WORD INTEGERS
*LIST SOURCE PROGRAM
SUBROUTINE SURF(SHANG*PEROD*HEIGT*SLOPE*HB*BRANG*VLS*DEPTH*LSCOP)
*** THE CONGENIES CURRENT VELOCITY
          P1=3-14159
R4D=57-2958
         MADD3742958
6=980-62
HTCM=100.*HEIGT
HBCM=.383*G**0*2*(PEROD*HTCM**2)**0.4
DEPTH=1.2*HBCM
WLO=G/(2.*PI)*PEROD**2
C
C
C
         DISPERSION EQUATION USED TO DETERMINE SHALLOW WATER WAVE LENGTH
         DO 10 I=1+20 WL1=WL0=TANH(2+PI+DEPTH/WL1)
     10 CONTINUE
         SHELL'S LAW USED TO DETERMINE BREAKER ANGLE AND HEIGHT DUE TO WAVE REFRACTION.
         SINA SIN(SHANG/RAD) TANH(2. PI TEPTH/WL1)
COSA SORT(1. SINA P2)
CALL ARCTA(BANG) COSA SINA)
RRANG-RAD BANG
         PRANGERADEDANG
REFRC=SGRT (COS (SHANG/RAD)/COS (BANG))
HB=REFRC=HBCM
         OPTION OF FOUR DIFFERENT LONGSHORE CURRENT EQUATIONS.
         GO TO (1.2.3.4) . LSCOP
         LONGSHORE CURRENT - FOX AND DAVIS: 1972
      1 VLS=100.*SLOPE*MB/PEROD*SIN(4.*BRANG/RAD)
GO TO 5
C
C
C
         LONGSHORE CURRENT - LONGUET-HIGGINS. 1970
      2 VLS - 9.00SLQPE+SQRT( G+H8 )+S1N(2.+BRANG/RAD)
GO TO 5
         LONGSHORE CURRENT - C.E.R.C. . 1973
      3 VLS =20-7*SLOPE*SQRT( G*H8 | 1*SIN(2.*BRANG/RAD) GO TO 5
         LONGSHORE CURRENT - KOMAR AND INMAN, 1970
4 Gl=G/100.
VL$=100.=Gl=MEIGT ==2/8.=Gl=PERGD/(4.=P[]=SIN(2.=BRANG/RAD)=
1 COS(2.=BRANG/RAD)
3 H8=HB/100.
RETURN
END
// DUP
```

```
// JOB
// FOR
•IOCS(CARD+1132 PRINTER)
                                                                                                               ENRGY
                                                                                                                                      3-86-001
    ONE WORD INTEGERS

*LIST SOURCE PROGRAM

SUBROUTINE ENRGY(H+T+TIME+ISALT+E+IUNIT)
   SUBROUTINE ENRRY(HoToTIME.ISALTOE.IUNIT)

C SUBROUTINE ENERGY IS USED TO DETERMINE WAN
HOUR. THIS ENERGY IS DERIVED FROM THE ENER
AND SUMED TO FIND THE TOTAL ENERGY IN THE

C STAND SUMED TO FUND THE TOTAL ENERGY IN THE

C STAND NUMBER OF WAVES IN GIVEN TIME PERIOD

C SUBROUTINE ENRRY(HoToTIME.ISALTOE.IUNIT)
                 SUBROUTINE ENERGY IS USED TO DETERMINE WAY! ENERGY DURING EACH HOUR. THIS ENERGY IS DERIVED FROM THE ENERGY WITHIN A SINGLE WAVE AND SUMED TO FIND THE TOTAL ENERGY IN THE STORM.
                 WAVES-TIME-3600./T
   C... FIND PROPER MASS DENSITY OF WATER (O-FRESH-1-SALT)
                IF(ISALT.EQ.0) DENSE=1.94
IF(ISALT.EQ.1) DENSE=2.0
IF(ISALT.GT.1) GO TO 500
   C CONVERT WAVE HEIGHT FROM METERS TO FEET
                HT=H/0.3048
   C C... FIND ENERGY FOR A SINGLE WAVE
               E=5.12+DENSE+32.1+(HT+T)++2./8.
  C... FIND TOTAL ENERGY
   C ... IF OUTPUT IS DESIRED IN FT-LBS/FT.RETURN C
                IF(IUNIT.EQ.O) RETURN
IF(IUNIT.NE.1) GO TO 500
  C... CONVERT FT-LBS/FT TO JOULES/METER
              E=E+1.35582/0.3048
RETURN
  C ... ERROR CONDITION
              WRITE(3,600)
FORMAT(1x, 'ILLEGAL OPTION IN SUBROUTINE ENRGY')
RETURN
  500
              END
  // DUP
*DELETE
*STORE
                           ENRGY
WS UA ENRGY
// JOB
// FOR
// FOR
PONE WORD INTEGERS
PLIST SOURCE PROGRAM
SUBROUTINE ARCTA(ANGLE.X.Y)

WIGHT ROUTINE IS USED
TO SIN AND COST
                                                                                                             ARCTA
                                                                                                                                    3-86-001
            ARCTANGENT ROUTINE IS USED TO DETERMINE THE ANGLE FROM \boldsymbol{x} and \boldsymbol{y} coordinates. Or sin and cosin.
C ANGLE = 0.0

IF(ABS(X)-.001)2.9.9

9 IF(X) 1.3.3

3 IF(Y) 4.9.5

4 ANGLE = 6.2831825

GO TO 5

1 ANGLE = 3.1815926

5 ANGLE = ANGLE+ATAN(Y/X)

RETURN
2 IF (ABS(Y)-.001) 8.10.10

10 IF(Y)6.7.7

6 ANGLE-4.7123889

RETURN
7 ANGLE = 1.5707963
8 RETURN
6 RETURN
7 OPELETE ARCTA
95TORE WS UA ARCTA
```

DELETE STORE

SURF WS UA SURF

```
// JOB
// FOR
+10CS(CARD+1132 PRINTER+TYPEWRITER+KEYBOARD)
                  $(CARO-1132 PRINTER-TYPEWRITER-KEYBOARD)
WORD INTEGERS
COMMON U(130)+V(130)
DIMENSION TITLE(20)+DAY(5)
DIMENSION A1(2)+A2(2)+A3(2)+A4(2)+B1(2)+B2(2)+C1(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C2(2)+C
c
                        STORM PROGRAM LIMITED TO FORECAST MODE
                         TYPEWRITER INPUT FOR THE STORM PARAMETERS
                        CARD 1 - TITLE
                       CARD 2 - STARTING DATE: TIME AND INPUT-OUTPUT OPTIONS
                                                  COLS. 1-2 ISTRT- STARTING HOUR COLS. 3-22 DAY - STARTING DATE
                                                                          23 INAUT- INPUT OPTION
                                                  COL.
                                                                                                                           0 = METRIC UNITS
1 = NAUTICAL MILES AND FEET
                                                                                               NAUT - OUTPUT OPTION

1 = NAUTICAL MILES. FEET AND KNOTS

Q = METRIC UNITS
                                                   COL.
                                                                          24
                       CARD 3 - STORM PARAMETERS
                                                                                                TIDE PREDICTION OPTION
0 = TIDE PREDICTION NOT INCLUDED = OMIT CARD 4
1 = TIDE PREDICTION INCLUDED = SEE CARD 4
                                                  COL.
                                                                      2
                                                                                                LONGSHORE CURRENT EQUATION OPTION
                                                  COL.
                                                                          3
                                                                                                LONGSMORE CORRENT EQUATION

1 = FOX AND DAVIS, 1972

2 = LONGUET-HIGGINS, 1970

3 = C.E.R.C.+, 1973

4 = KOMAR AND INMAN, 1970
                                                                                                       NX - NUMBER OF STORM POSITIONS
SIX HOUR INTERVALS FOR HINDCASTING
ONE HOUR INTERVALS FOR FORECASTING
                                                  COL5. 4-6
                                                 COLS. 7-12 BNFCH- AVERAGE BASIN FETCH IN KM. (NAUT. MI. COLS. 13-17 TINT - TIME INTERVAL BETWEEN STORM POSITIONS NORMAL SETTING IS 1.0 HOURS
                                                 COLS. 32-36 SLAT - LATITUDE AT SHORE SITE
                      CARD 4 - TIDE PREDICTION - OPTION FROM CARD 3
COLS. 0-5 ST - SPRING TIDE RANGE IN FEET
COLS. 6-10 TN - NEAP TIDE RANGE IN FEET
COLS. 11-15 TOAY - NUMBER OF DAYS FROM LAST SPRING TIDE
COLS. 16-20 THR - MOUR OF LAST HIGH SPRING TIDE
                                                 COLS. 21-25 FN
                                                                                                              - TIDAL FORM NUMBER
                                                                                                                              1 DAL FORM NUMBER
0.0 TO .25 - SEMIDIURNAL TIDE
0.25 TO 1.5 - MIXED SEMIDIURNAL TIDE
1.5 TO 3.0 - MIXED DIURNAL TIDE
GREATER THAN 3.0 - DIURNAL TIDE
                                                 COLS. 26-32 SLPLO- NEARSHORE BOTTOM SLOPE AT LOW TIDE COLS. 33-39 SLPHI- NEARSHORE BOTTOM SLOPE AT HIGH TIDE COLS. 40-46 TMEAN- MEAN TIDE LEVEL IN FEET
                      CARD 5 - SHORE SITE LOCATION
                                                 COLS. 1-7 ULOC - X-COORDINATE IN NAUTICAL MILES
COLS. 8-14 VLOC - Y-COORDINATE IN NAUTICAL MILES
COLS. 15-21 SHAZ - ONSHORE DIRECTION - CLOCKWISE FROM NORTM
COLS. 22-29 SLOPE - AVERAGE NEARSHORE BOTTOM SLOPE
COL. 30 ISLND- 0 - CONTINENTAL COAST OR BARRIER ISLAND
                           THE FOLLOWING INFORMATION IS TYPED IN FROM THE CONSOLE FOR EACH
                          STORM SIMULATION RUN.
                                                                                               PMIN - MINIMUM BAROMETRIC PRESSURE IN MILLIBARS
PMAXR- PRESSURE AT LARGEST ENCIRCLING ISOBAR
AR - MAJOR HALF-AXIS (EFFECTIVE LONG RADIUS)
BR - MINOR HALF-AXIS (EFFECTIVE SMORT RADIUS)
EAZ - ORIENTATION OF MAJOR HALF AXIS PLUS OR
MINUS 90 DEGREES FROM NORTH:

SVEL - STORM VELOCITY IN KNOTS
AZI - STORM AZIMUTH - CLOCKWISE FROM NONTH
XIII - INITIAL X-COORDINATE FOR THE STORM.
Y(I) - INITIAL X-COORDINATE FOR THE STORM.
                                             TYPE
                                                  IN
                      CNK=1.85319

CKN=.53961

CFM=.3048

CMF=3.2808

CKTS=1.9425

CCMF=.032808

CJOUL = .737561
```

```
c
RAD=57.2958
8080 READ (2.915) TITLE
915 FORMAT(20A4)
READ(2.917) ISTRT.DAY.INAUT.NAUT
       READ(2-917) ISTRT-|
917 FORMAT([2-5A6-21])
HN=NAUT+]
IF([NAUT) 1:1-2
1 CNK=1-0
CFM=1-0
2 IF(NAUT) 11-11-12
           11 CKM=1.0
CMF=1.0
CKTS=1.0
CCMF=1.0
      CKTS=1=0
CMP=1=0
CJQUL=1=0
12 IF (ISTRT.EQ=0=0) GO TO 1000
READ(2+901) INOPT+IFTID+LSCOP+NX+BNFC1+TINT+SLAT
901 FORMAT(31)=13+F6+0+F5+1+14X+F5+1)
BNFCH+0BNFC1=CKK
IFIIFTID) 5-5+6
4 READ(2+902) 51+T1+TDAY+THR+FN+SLPLO+SLPHI+TMEA1
920 FORMAT(5F5-2+3F7-2)
ST-S1=0FFM
TN=11=CFM
TN=11=CFM
TNEAN=THEA1+CFM
9 READ(2+907) ULC1+VLC1+SHAZ+SLOPE+ISLND
907 FORMAT(4F7-0+12)
ULC1=ULC1=CHK
VLC1=VLC1=CHK
VLC1=VLC1=CHK
945 WRITE(1+947)
947 FORMAT(4*TYPE IN THE MINUMUM AND MAXIMUM BAROMETRIC PRESSURE IN MI
1L18ARS - FORMAT AAAA-BBBB+')
READ(6+948) PMIN+PMAXR
948 FORMAT(2F5-0)
        948 FORMAT(275.0)

wRITE(14.955)
955 FORMAT! TYPE IN THE LENGTH OF THE MAJOR AND MINOR HALF AXES OF TH
1E STORM - FORMAT XXXX.YYYY.')
        READ(6.956) AR1.BR1
956 FORMAT(2F5.0)
       956 FORMAT(275-0)
WRITE(1.980)
960 FORMAT(1 TYPE IN THE ORIENTATION OF THE MAJOR HALF AXIS IN DEGREES
1 FROM NORTH - FORMAT XXX.')
READ(6.961) EAZ
961 FORMAT(F3-0)
                    BAROMETRIC PRESSURE AT LARGEST ENCIRCLING ISOBAR IS ASSUMED TO BE AT TWO STANDARD DEVIATIONS FROM THE CENTER OF THE STORM. TO FIND THE ACTUAL STORM SIZE. THE MAJOR AND MINOR AXES ARE MULTIPLIED BY 1.9 AND THE MAXEMUM PRESSURE IS MULTIPLIED BY 1.145 TO DETERMINE THE PRESSURE AT THE MARGIN OF THE STORM.
                      BR=BR1 + CNK
RANGP=PMAXR-PMIN
                      PMAX -RANGP-1.145+PMIN
                     AR AND BR ARE RADII OF THE STORM OUT TO TWO STANDARD DEVIATIONS... MULTIPLY BY 1.5 TO YIELD THE FULL LENGTHS.
                      A-1.5-AR
                   A=1.59AR
B=1.59BR
CALL FOREC(NX.A.SVEL.AZI)
DO 42 I=1.nX
U(1)=U(1)=CNK
V(1)=V(1)=CNK
ULOC=ULC1/A
VLOC=ULC1/A
VLOC=VLC1/A
                      U1=U(1)*A+.005
V1=V(1)*A+.005
                     UL-ULOC-A
                        WRITE SHORE AND TIDE DATA
     ULI=UL=CKN++05

VLI=VL=CKN++05

BNFCI=BNFCM=CKN

WRITE:39+01 TiTLE

916 FORMAT(|M|//olx-20Ac/)

WRITE:39+09:ISTRT+OAY

18 FORMAT(|X*,"NUM BEGINS AT MOUR '+12+' ON '+5Ac)

WRITE:39-09: PRIN-PMAXR+PMAX

909 FORMAT(|X0*,"STORM - BAROMETRIC PRESSURE AT CENTER OF LOW -*'+F7+1+

1 MILLIBARS'/94+'PRESSURE AT LARGEST ENCIRCLING ISOBAR -*'+F7-1+

2 MILLIBARS'/94+'MAXIMUM PRESSURE INCLUDED IN STORM -*'+F7-1+

31 MILLIBARS'/9

ARI=ARGCKN++05

BRI=BRECKN++05
 918
                     #R1=#R=CKN+=09
#R1=E(3=926| AR1=A1(NN)=A2(NN)=A5(NN)=A6(NN)=BR1=A1(NN)=A2(NN)=
      WRITE(3-926) ARI-ALINNI-AZINNI-AZINNI-AZINNI-AAINNI-AZINNI

AZINNI-AAINNI-AEAZ

926 FORMATIGN.'LENGTH OF MAJOR MALF AXIS =' -F7-1-6A6/

19X-'LENGTH OF MINOR MALF AXIS =' -F7-1-6A6/

29X-'DRIENTATION OF MAJOR AXIS =' -F7-1-' DEGREES FROM NORTH'/I

JFINAUTI601-601-602

601 WRITE(3-9101 SVEL-AZI

910 FORMATIGN.'STORM VELOCITY =' -F6-0-' KILOMETERS/MOUR'-

1 9X-'STORM AZIMUTH =' -F6-0-'
       1 9%:"STORM AZIMUTH "":F4.0/)
GO TO ADS
602 WRITE(5:911) SVEL:AZI
911 FORMAT:9%:"STORM VELOCITY "":F4.0" KNOTS":
1 9%:"STORM AZIMUTH "":F4.0/)
609 CONTINUE
```

```
3 = '.F6.0.4AA/)

IF(IFTID) 59.59.58

58 S1=ST+CMF

T1=TN+CMF
       WRITE(3.927) S1.T1.B1(NN).B2(NN).SLPLO.SLPHI
927 FORMAT(' TIDES - SPRING TIDE RANGE -'.F6.2.' NEAP TIDE RANGE -'
1F8.2.2A4 /9X.'SLOPE AT LOW TIDE -'.F6.3.' SLOPE AT HIGH TI
                                                                                                                                                                                         SLOPE AT HIGH TIDE
      1F8.2.244 /9X'SLOPE AT LOW TIDE =".F6.3." SLOPE AT MI

1 =".F6.3/1

IF(FN.LE..25) WRITE(3.941) FN

941 FORMAT(9X."SEMIDIURNAL TIDE - FORM NUMBER IS'.F6.2/)

IF(FN.GT..25.AND.FN.LE.3.0) WRITE(3.942) FN

942 FORMAT(9X."MIXED SEMIDIURNAL TIDE - FORM NUMBER IS'.F6.2/)

IF(FN.GT.1.5.AND.FN.LE.3.0) WRITE(3.943) FN

943 FORMAT(9X."MIXED DIURNAL TIDE - FORM NUMBER IS'.F6.2/)

IF(FN.GT.3.0) WRITE(3.944) FN

944 FORMAT(9X."DIURNAL TIDE - FORM NUMBER IS'.F6.2/)

95 WRITE(1.830)
       59 WRITE(1.830)
830 FORMAT() PLEASE CHECK TO SEE THAT THE INFORMATION ON THE PRINTER I
15 CORRECT' / ' IF THE PROGRAM IS ALL SET TO GO. PRESS THE 1 KEY. O
2THERWISE. PRESS THE 2 KEY!)
       READ(6.831) NNN
831 FORMAT(11)
IF(NNN-1) 590,590,945
    831 FORMAT(11)

1F(NNN-1) 590,590.995

900 MRITE(3.900)

00 FORMAT(1N0.' HOUR' .T11.'X',T18.'Y',T24.'X1',T31.'Y1',T37,'BARO.'.

1T45.'WIND'.T52.'SURF.'.T59,'ONSH'.T66.'ALSH'.T72.'EFFLCT.'.T81.

2'WAVE'.T88.'WAVE'.T98.'BREAKER'.T1.09.'LSC'./LX.T36.'PRESS.'.144.

3'ANGLE'.T52.'WIND'.T59.'WIND'.T66.'WIND'.T73.'WIND '.T83,'H'.T90.

4'T'.T97.'H'.T101.'ANGLE'.T108.'VELOC.')

1F(IFID.GT.0) WRITE(3.951)

951 FORMAT(1H0.T117,'T1DE')

1F(NAUT) 800.800.800

800 WRITE(3.801)

801 FORMAT(1H0.T10.'KM'.T17.'KM'.T24.'RAD'.T31.'RAD'.T38.'MB'.T45.

1 'DEG'.T53.'M/S'.T60.'M/S'.T67.'M/S'.T74.'M/S'.T83.'M'.T89.'SEC'

2 .T97.'M'.T102.'DEG'.T108.'CM/SEC')

1F(IFID.GT.0) WRITE(3.802)

809 WRITE(3.802)

820 FORMAT(1H0.T10.'MM'.T17. 'NM'.T24.'RAD'.T31.'RAD'.T38.'MB'.T45.

1 'DEG'.T53.'KTS'.T60.'KTS'.T74.'KTS'.T74.'KTS'.T82.'FT'.T89,'SEC'

2 .T90.'FT'.T102.'DEG'.T108.'FT/SEC')

1F(IFID.GT.0) WRITE(3.822)

1F(IFID.GT.0) WRITE(3.822)
      IF(IFTID+GT+0) WRITE(3+822)
822 FORMAT(IH++T118+'FT')
815 WRITE(3+952)
       952 FORMAT(10x)
                        CALL SUBROUTINES TO DETERMINE LOCATION AND COMPUTE WIND. WAVES, AND TIDES AT EACH LOCATION
                 DO 50 I=1.NX

CALL LOCAT(U(I).V(I).ULOC.VLOC.SHAZ.X.Y.EAZ.P.U)

CALL ELIPS(A.B.P.U.PMIN.PMAX.P1.ERAD.XASTASDZA)

CALL WIND(X.Y.P.PMIN.PMAX.SLAT.WSURF.ALSH.ONSH.SHANG.ERAD.P1.A.XA.

1YA.SHAZ.EAZ.DZA.ISLND.EFWND)

IF(I.EU.) GO TO 510

CALL DECAY(IINT.PEROD.HEIGT)

HDECY-HEIGT

CALL ETIME(EFWND.HEIGT.TINT.DRATN)

GO TO 520

DRATN-TINT

FSUM-0.0
                     ESUM=0.0
         EN=0.0
EP=0.0
IMOUR=ISTRT

CALL FETCH(A.B.x.y.*STFCH)
TLFCH**STFCH
IF(BNFCH**LT**STFCH)TLFCH**BNFCH
CALL WAVES(EF WND.DRATN*TLFCH**EFETC**HEIGT**PEROD**1)
IF(IFTID) 36*56*55
55 CALL TIDE(ST**TN**FN**)ISTRT**TN**TDX**SLPMI**SLPLO**SLOPE*1**TMEAN)
36 CALL SURF(S**MANG**PEROD***HEIGT**SLOPE**HB**BRANG**VLS**DEPTM**LSCOP)
IF(BRANG**180**L) 49**49*48
BRANG**RBANG**180**L
                     EN=0.0
520
**
                      BRANG-BRANG-360.
                     COMPUTE WAVE AND LONGSHORE CURRENT ENERGY
                     EL$+.0257+(0.6+VL$1++2.+DEPTH++2./SLOPE
                      IF(VLS) 45.46.47
                     EN-EN-ELS
GO TO 46
EP-EP-ELS
                      CONTINUE
                     CALL ENRGY(HB.PEROD.TINT.1.E.1)
ESUM-ESUM+E
UPT-U(1)*A*CKN+.5
                      VOTAVI I I BARCK Ne.5
                     WSRF-WSURF#CKTS
ON-ONSH-CKTS
AL-ALSH-CKTS
                     EF-EFWND+CKTS
HT1=HE1GT+CMF
HB1+HB+CMF
                      V1-VLS-CCMF
```

```
WRITE(3,530) [HOUR:UPT:VPT:X:Y:P:SHANG:WSRF:ON:AL:EF:HT1:PEROD:

1 HB1:BRANG:V1

30 FORMAT(3X:[2:2F7.0:2F7.2F6.1:4F7:1:F7.1:F8.2:F7.1:F7.2:2F7.1)

IF(IFTID:GT.0) WRITE(3:531) TIDX

531 FORMAT(1H:-114X:F5.2)

IHOUR:IHOUR:IFIX(TINT)
       530
                         IF(IHOUR-GT-24)IHOUR-IHOUR-24
50 CONTINUE
ELSCT-EN-EP
                 ELSCT=EN+EP

ESUM=ESUM=CJOUL

ELSCT=ELSCT=CJOUL

EP=EP=CJOUL

EN=EN=CJOUL

WRITE(3-540) ESUM-C1(NN).C2(NN).C3(NN).ELSCT.C1(NN).C3(NN).C3(NN)

540 FORMAT(1X-1/-1X-75'MAVE ENERGY IN THE BREAKER ZONE ='.El0.3.3A4.//

1.11x.T5,'TOTAL LONG-SHORE CURRENT ENERGY ='El0.3.3A4.//

WRITE(3,542) EP-C1(NN).C2(NN).C3(NN).EN-C2(NN).C2(NN).C3(NN)

542 FORMAT(1X-T5.*TOTAL POSITIVE LONG-SHORE CURRENT ENERGY ='.El0.3.

1344-//.1x.T5.*TOTAL NEGATIVE LONG-SHORE CURRENT ENERGY ='.El0.3.

2344)
     13A4-//six-T5, TOTAL NEGATIVE LONGSHORE CURRENT ENERGY ='.E10.3, 23A4)

WRITE(3,992)

992 FORMAT(IM1)

WRITE(1990)

990 FORMAT(' IF YOU WANT TO TRY ANOTHER SIMULATION RUN. TYPE IN A 1.

10THERWISE TYPE IN A ZERO')

READ[6,991] NEND

991 FORMAT(I1)

IF(MENDA-EG-1) GO TO 945

GO TO 8080

1000 CALL EXIT

END

// DUP

DELETE CSTRM
      DELETE
STORE
                                                                                  CSTRM
WS UA CSTRM
// JCB
// FOR
*IDCS(CARD+)132 PRINTER+TYPEWRITER+KEYBOARD)
*ONE WORD INTEGERS
*LIST SOURCE PROGRAM
*LIST SOURCE PROGRAM
SUBROUTINE FOREC(NX+A+SVEL+AZI)
COMMON U(130)+V(130)
**COMPUTE STORM POSITIONS**
                                                                                                                                                                                                                                                                                                                                       FOREC
                                                                                                                                                                                                                                                                                                                                                                                                             3-86-001
                                           FORECASTING - COMPUTE STORM POSITIONS FROM INITIAL POSITION VELOCITY AND AZIMUTH
               TINU=1+0
RAD=57,2958
WRITE(1+975)
975 FORMAT(' TYPE IN THE STORM VELOCITY AND AZIMUTH - FORMAT VVV+AAA.'
  975 FORMAT(* TYPE IN THE STORM VELOCITY AND AZIMUTH - FORMAT VVV.AAAA
1)

READ(6:976) SVEL:AZ[
976 FORMAT(2F4:0)

MPITE(1:970)

970 FORMAT(* TYPE IN THE INITIAL X AND Y COORDINATES OF THE STORM -
1 FORMAT XXX.YYYY.*)

READ(6:971) U(1):V(1)

971 FORMAT(2F5:0)

U(1):U(1):A

V(1):U(1):A

V(1):U(1):A

AZM=90.-AZ[

IF(AZM) 31:32:32

31 AZM=AZM=360.

32 DO 35 1:2:AX

U(1):U(1-1):DIST*COS(AZM/RAD)

V(1):U(1-1):DIST*SIN(AZM/RAD)

35 CONTINUE

RETURN

END

// DUP

**POLETE**

FOREC

**STORE**

**FOREC

     *DELETE
                                                                               FOREC WS UA FOREC
     -STORE
```

APPENDIX B. HINDCAST STORM DATA AND OUTPUT

```
MARONALN ISLANDS - EAST SIDE

JOURN 100 000 11 0976 1008 48, 860. 860. 0.0

186 137 10. 214 5.0 0.020 1

185 137 10. 214 5.0 0.020 1

185 138 138 1194

138 138 128 1194

138 138 128 1194

138 138 128 1194

140 195 128 1290

190 130 128 130 1290

111 130 000 12 1290

112 130 000 12 1290

112 130 000 12 1290

113 13 129 1290

114 13 1290

115 13 1290

116 13 1290

117 120 777

180 122 120 120

1170 777

180 122 120 120

122 120 120

122 120 120

122 120 120

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124 120 120

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                                                                                                                                                   DATA LIST FOR HINDCAST EXAMPLES
| Struchyllich Michigh
| 1001 | 1200 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100 | 100
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HOLLAND . MICHIGAN

RUN BEGINS AT HOUR 1 ON JULY 3: 1970

STORM - BAROMETRIC PRESSURE AT CENTER OF LOW = 1003.0 MILLIBARS
PRESSURE AT LARGEST ENCIRCLING ISOBAR = 1010.0 MILLIBARS
MAXIMUM PRESSURE INCLUDED IN STORM = 1011.0 MILLIBARS

LENGTH OF MAJOR HALF AXIS = 550.0 KILOMETERS
LENGTH OF MINOR HALF AXIS = 200.0 KILOMETERS
ORIENTATION OF MAJOR AXIS = 30.0 DEGREES FROM NORTH

LONGSHORE CURRENT EQUATION FROM FOX AND DAVIS. 1972

SHORE - POSITION COORDINATES - X = 1015.0 Y = 560.0 KILOMETERS

SHORE LATITUDE = 42. ONSHORE AZIMUTH = 90. DEGREES NEARSHORE SLOPE = 0.033 AVERAGE FETCH = 150. KILOMETERS

HOUR	×	¥	×ı	۲1	BARO. PRESS.	WIND	SURF.	UN5 4	AL SH WIND	EFFLCT.	WAVE	WAVE T	8.H	EAKER ANGLE	LSC VELUC.
	KM	KM	RAD	RAD	MB	CEG	M/S	M/S	M/S	M/S	н	SEC	M	DEG	CM/SEC
1	968.	1030.	0.56	0.05	1010.98	210.9	0.2	-0.1	-0.1	0.0	0.00	0.0	0.00	-14.9	-0.94
2	984.	1016.	0.55		1010.95	211.9	0.4	-0.3	-0.2	0.1	0.00	0.1	0.00	-15.7	-1.65
3	1001.	1003.	0.53		1010.89	213.1	0.7	-0.6	-0.3	0.2	0.00	0.1	0.00	-16.5	-3.24
4	1017.	989.	0.52		1010.79	214.5	1.2	-0.9	-0.6	4.4	0.00	0.2	0.00	-17.3	-5.2.
5	1033.	976. 962.	0.50		1010.43	216.3	1.8	-1.4	-1.1	0.7	0.01	0.4	0.01	-18.2	-7.85
7	1066.	949.	0.47		1010.08	221.3	2.6 3.6	-2•1 -2•7	-1.6 -2.4	1.0	0.02	0.5	0.01		-11.03
8	1071.	932.	0.45		1009.75	222.8	4.4	-3 - 2	-3.0	1.8	0.05	0.9	0.05		-17.77
9	1077.	914.	0.42		1009.36	224.6	5.2	-3.7	-3.6	2.2	0.07	1.0	0.06		-20.84
10	1082.	897.	0.40		1008.91	226.9	5.9	-4.0	-4.3	2.6	0.09	1.2	0.08		-23.53
11	1087.	880.	0.38	-0.08	1008.41	229.7	6.4	-4.1	-4.9	2.9	0.12	1.4	0.11	-23.9	-25.68
12	1093.	862.	0.36		1007.87	233.4	6.7	-4.0	-5.4	3.1	0.15	1.5	0.12		-26.84
13	1098.	845.	0.34		1007.31	238.4	6 - 8	-3.5	-5.8	3.3	0.17	1.0	0.14		-26.61
14	1099.	828.	0.32		1006.84	242.0	6.7	-3 - 1	-5.9	3.4	0.19	1.7	0.15		-26.12
16	1099.	810. 793.	0.28		1006.37	246.6	6.5	-2.6	-6.0	3.5	0.21	1.6	0.15		-24.35
17	1101.	776.	0.26		1305.49	252.6	6.3	-1.8 -0.9	-6.0	3 · 6 3 · 7	0.23	1.9	0.14		-21.00 -15.26
18	1101.	758	0.24		1005.10	271.4	6.1	0.1	-6.1	4.1	0.27	2.1	0.05	-32.0	-6.27
19	1102.	741.	0.21		1004.76	284.3	6.1	1.5	-5.9	4.6	0.31	2.2	0.17		-21.74
20	1112.	736.	0.21		1004.70	306.2	6.0	3.5	-4.8	5 • 2	0.36	2.4	0.30		-41.51
21	1122.	730.	3.20	-C.13	1004.72	330.4	6.2	5.4	-3.0	5.9	0.42	2.6	0.42		-47.24
22	1132.	725.	0.20		1004.83	350.0	7.2	7.1	-1.2	7.2	0.52	2 • 8	0.55	-5.3	-22.92
23	1143.	720.	0.19		1005.02	3.1	8.7	8 • 7	0.4	8.7	0.67	3.2	0.70	1.6	8.45
24	1153.	714.	0.18		1005.28	11.7	10.2	10.0	2.0	10.1	0.85	3.6	0.88	6.2	34 - 1 3
l 2	1163.	709. 706.	0.18 0.17		1005.61	17.5 20.8	11.5	11.0	3.5	11.4	1.05	4.0	1.08	9.3	53.71
3	1181.	703.	0.17		1006.27	23.3	13.0	11.6	4.4 >•1	12.7	1.43	4.7	1.27	10.9	6 6.3 3 7 6.6 7
4	1189.	700.	0.17		1006.63	25.3	13.4	12.1	5.7	13.0	1.60	5.0	1.62	13.2	85.07
5	1198.	698.	0.16		1007.00	26.9	13.5	12.1	6.1	13.0	1.74	5.2	1.75	14.0	91.60
6	1207.	695.	0.16		1007.36	28.2	13.5	11.8	6.4	12.9	1.86	5.4	1.67	14.6	97.07
7	1216.	692.	0.16		1007.73	29.4	13.2	11.5	6.4	12.6	1.96	5.5	1.95	15.1	101.04
8	1225.	691.	0.15		1008.08	30.2	12.7	10.9	6.4	12.1	2.02	5 • 6	2.01		103.62
10	1234.	690. 689.	0.15		1008.41	30.9	12.0	10.3	6.2	11.5	2.06	5 • 7	2.04		105.25
ii	1253.	688.	0.15		1008.72	31.5 32.0	11.3	9.6	5.9	:0.7	1.92	5.5	1.90		102.31
12	1262.	687.	0.15		1009.28	32.5	9.6	8.9 8.1	5.5 5.1	9•9 9•1	1.71	5.2	1.70	16.3	97.18 91.36
13	1271.	686.	0.15		1009.53	32.9	8.7	7.3	4.7	b • 2	1-29	4.5	1.28	16.7	85.00
14	1278.	686.	0.15		1009.70	33.2	8.0	6.7	4.4	7.6	1.14	4.3	1.12	16.8	19.90
15	1285.	685.	0.15		1009.86	33.5	7.3	6.1	4.0	6.9	3.99	4.0	V.98	16.9	74.63
46	1292.	685.	0.15		1010.00	33.7	6.7	5.5	3.7	6.3	0.85	3.7	0.64	17.0	69.25
17	1299.	684.	0.15		1010.13	34.0	6.0	5.0	3.3	5.7	0.72	3.4	0.71	17.1	63.62
19	1313.	684. 683.	0.14		1010.24	34.2	5 4 4 4 • 8	4.5	3.0	5 - 1	0.61	3 . 1	0.60	17.1	58.37
20	1321.	684.	0.14		1010.45	34.5	4.2	4 • 0 3 • 5	2.7	4 • 5	0.50	2.9	0.49	17.2	52.98
21	1329.	684.	0.15		1010.54	34.7	3.7	3.0	2.4	3.4	0.40	2.6	0.39	17.2	47.04
25	1337.	685.	0.15		1010.62	34.8	3.1	2.6	1.8	3.0	0.24	2.3	0.23	17.1	41.32 35.90
23	1346.	685.	0.15		1010.69	34.9	2.7	2.2	1.5	2.5	0.18	1.7	0.17	17.0	33.84
24	1354.	486.	0.15		1010.75	35.1	2.3	1.8	1.3	2.1	0.13	1.5	J.13	17.0	26.19
1	1362.	686.	0.15		1010.79	35.2	1.9	1.5	1.1	1 - 8	0.09	1.3	0.09	16.9	22.00
2	1369.	687.	0.15		1010.83	35.2	1.7	1.3	0.9	1.6	0.07	1.1	0.07	16.8	19.10
•	1375.	688. 689.	0.15		1010.85	35.3	1.4	1 • 2	0.8	1.3	3.05	1.0	0.05	16.7	16.48
•	1388.	690.	0.15		1010.88	35.4 35.4	1.2	1.0	0.7	1.1	0.04	0.8	0.04	16.6	14.14
6	1395.	691.	0.15		1010.90	35.5	1.0	0.8	0.5	1.0 0.8	0.02	3.7	0.03	16.0	12.08
7	1401.	692.	0.16		1010.93	35.5	0.7	0.7	0.5	0.8	0.02	0.5	0.02	16.5	8.69
	•									•••	0.01		5.01		,

WAVE ENERGY IN THE BREAKER ZONE - 0-153E 10 JOULES
TOTAL LONG-SHORE CURRENT ENERGY - 0-157E 10 JOULES
TOTAL POSITIVE LONG-SHORE CURRENT ENERGY - 0-156E 10 JOULES

TOTAL NEGATIVE LONG-SHORE CURRENT ENERGY . 0.4386 07 JOULES

HOLLAND, MICHIGAN

RUN BEGINS AT HOUR 1 ON JULY 7, 1970

STORM ~ BAROMETRIC PRESSURE AT CENTER OF LOW = 1002+0 MILLIBARS
PRESSURE AT LANGEST ENCIRCLING ISOBAR = 1010+0 MILLIBARS
MAXIMUM PRESSURE INCLUDED IN STORM = 1011+1 MILLIBARS

LENGTH OF MAJOR HALF AXIS = 455.0 KILOMETERS
LENGTH OF MINOR HALF AXIS = 244.0 KILOMETERS
ORIENTATION OF MAJOR AXIS = 35.0 DEGREES FROM NORTH

LONGSHORE CURRENT EQUATION FROM FOX AND DAVIS. 1972

SHORE - POSITION COORDINATES - X = 1015.0 Y = 560.0 KILOMETERS

SHORE LATITUDE = 42. ONSHORE AZIMUTH = 90. DEGREES NEARSHORE SLOPE = 0.033 AVERAGE FETCH = 150. KILOMETERS

HOUR	. x	Y	*1	71	BARU. PRESS.	WIND ANGLE	SURF, WIND	ONSH W1ND	ALSH WIND	EFFECT.	WAVE H	WAVE T	8 F	LAKER ANGLE	LSC VELUC.
	KM	KM	RAD	RAD	мв	DEG	M/S	M/S	M/S	M/S	M	\$EC	M	DEG	CM/SEC
1	452 •	949.	0.56		1011.15	193.6	0.0	-0.0	-0.0	0.0	0.00	0.0	0.00	-6.5	-0.00
2	475. 497.	949.	0.56		1011.15	193.9	0.0	-0.0	-0.0	0.0	0.00	0.0	0.00	-6.6	-0.30
2	520	948.	0.56		1011.15	194.3	0.0	-0.0	-0.0 -0.0	0.0	0.00	0.0	0.00	-7.0	-0.00
5	543.	948.	0.56	3.49	1011.15	195.0	0.0	-0.0	-0.0	0.0	0.00	0.0	0.00	-7.1	-0.00
6	565 .	948.	0.56		1011.15	195.4	0.0	-0.0	-0.0	0.0	0.00	0.0	0.00	-7.3	-0.00
7	588 *	948.	0.56		1011.15	195.8	0.0	-0.0	-0.0	0.0	0.00	0.0	0.00	-7.5	-0.00
8	615.	942.	0.55		1011.15	196.2	0.0	-0.0	-0.0	0.0	0.00	0.0	0.00	-7.7	-0.00
9	642.	936.	0.55	0.54	1011-15	196.7	0.0	-0.0	-0.0	0.0	0.00	0.0	0.00	-7.9	-0.01
10	669• 697•	930. 925.	0.54		1011.15	197.2 197.8	0.0	-0.0	-0.0	0.0	0.00	0.0	0.00	-6.2	-0.03
12	724.	919.	0.52		1011.15	196.4	0.0	-0.0	-0.0	0.0	0.00	0.0	0.00	-8.5	-0.07 -0.15
13	751	913.	0.51		1011.14	199.1	0.1	-0.1	-0.0	0.0	0.00	0.0	0.00	~9.4	-0.31
14	773.	902.	0.50	0.35	1011.12	149.6	0.2	-0.1	-0.0	0.0	0.00	0.0	0.00	-4.6	
15	795.	890.	0.48		1011.09	200.2	0.3	-0.3	-0.1	0.1	0.00	0.0	0.00	-10.1	-1.14
16	817. 840.	879.	0.46		1011.03	200.8	0.6	-0.5	-0.2	0.2	0.00	0 • 1	0.00	-10.6	-1.50
18	862.	868. 856.	0.45		1010.94	201.6	1.0 1.5	-0.9	-0.3	0.3	0.00	0.2	0.00	-11-1	-3.23
19	884 .	845.	0.41		1010.57	203.6	2 • 2	-1.4 -2.1	-0.5 -0.9	0.5	0.00	0.3	0.00	-11.6	-7.22
50	908.	836.	0.40		1010.27	205.2	3.2	-2.9	-1.3	1.1	0.02	0.6	0.02		-10.12
21	932.	827.	0.39	0.12	1009.85	207.2	4.2	-3.8	-1.9	1.5	0.03	0.7	0.03		-13.60
22	955.	817.	0.37		1009.32	209.7	5 • 4	-4.7	-2.6	2.0	0.06	0.9	0.06		-17.77
23	979. 1003.	808. 799.	0.36		1008.67	213.1	6.5	-5.4	-3.5	2.5	0.08	1.1	0.08		-22.26
1	1027.	790	0.35	-0-01	1007.94	217.8	7+3 7+8	-5.8 -5.6	-4,5 -5.5	2.9 3.3	0.11	1.3	0.11		-20.65 -29.67
ž	1039	775.	0.31		1006.43	229.5	8.0	-5.2	-6.1	3.6	0.18	1.7	0.16		-31.50
3	1052.	760.	0.29	-0.05	1005.70	236.6	7,9	-4.3	-6.6	3 . 8	0.21	1.0	0.17		-30.01
4	1064.	744.	0.27	-0.07	1005.03	246.6	7.5	-2.9	-6.8	4.0	0.24	1.9	0.17		-25.92
5	1077.	729. 714.	0.24		1004.46	261.0	7,4	-1.1	-7.3	4.5	0.28	1	0.12		-15.95
7	1102.	699	0.22		1004.02	279.3	8.0	1.3	-7.9 -7.3	5.8	0.36	2.3	0.10 0.30		-18.39
8	1107.	692.	0.19		1003.67	308.5	8,2	5.1	-0.4	7.1	0.56	2.9	0.48		-53.43
9	1112.	685.	0.18		1003.63	318.5	8.2	6.1	~5.4	7.5	0.64	3 . 1	0.60		-62.10
10	1117.	678.	0.17		1003.62	328.3	B • •	7.1	-4.4	7.9	0.73	3.4	0.72	-10.2	-03.02
11	1121.	671. 664.	0.16		1003.65	337.3	8.7	8 - 1	-3.3	8 • 5	0.82	3.6	0.84		-50.58
13	1131.	657.	0.15		1003.71	345.4	9.3 9.9	9.0 9.8	-2.3 -1.3	9.2	0.93 1.04	3.8	0.95		-42+06 -25+01
14	1134.	652	0.13		1003.66	356.8	10.4	10.4	-0.5	10.4	1.16	4.0	1.22		-11.07
15	1138.	646.	0.12	-0.17	1003.98	0.8	10.9	10.9	0.1	10.9	1.28	4.5	1.35	0.4	2.97
16	1141.	640.	0.11		1004.10	4.2	11.5	11.4	0.8	11.5	1-41	4.7	48	6.6	16.51
17 18	1144.	635. 629.	0.10		1004.23	7.3	12.0	11.9	1.5	11.9	1.53	4.9	1 6 1	3.9	29.23
19	1151.	624.	0.10		1004.57	12.5	12.4	12.3	2 • 1 2 • 8	12.4	1.66	5.1	1.73 1.56	6.6	40.78 51.72
20	1157.	621.	0.08		1004.80	14.3	13.4	12.9	3.3	13.2	1.91	5.5	1.99	7.6	63.46
21	1164.	618.	0.08		1005.08	15.9	13.7	13.2	3.7	13.6	2.04	5.6	2.11	8.4	00.23
22	1170.	614.	0.07		1005.36	17.3	14.0	13.4	4.1	13.8	2 - 16	5 . 6	2.23	9.1	75.11
23 24	1177.	611.	0.07		1005.65	18.5	14.2	13.4	4 . 5	13.9	2.26	5.9	2.33	9.7	81-17
-7	1190.	605.	0.06		1006.25	20.6	14.2 14.2	13.4	4.6 5.0	14.0	2.36	6.1	2.43	10.3	86.48 91.09
ž	1199.	604.	0.06		1006.61	21.1	14.0	13.1	5.0	13.7	2.51	6.3	2.58	11.0	93.93
3	1208.	604.	0.06	-C - 28	1006.98	21.6	13.7	12.7	5.0	13.4	2.56	t • 4	2.63	11.2	46.22
4	1217.	603.	0.06		1007.33	22.0	13.3	12.3	5.0	13.0	2.50	6.3	2.55	11.5	40.15
5	1227.	602.	0.06		1007.67	22.4	12.8 12.2	11.8	4.9	12.5	2 - 36	6.1	2.42	11.6	94.62
7	1245.	601.	0.06		1008.30	23.0	11.0	10.7	4.5	11.3	2.21	5.9	2.20	11.8	92.35
8	1254 .	635,	0.06		1008.53	22.7	11.0	10.2	4.2	10.8	1.90	5.5	1.94	11.6	#5.46
	1262.	610.	0.07		1008.74	22.3	10.5	9.7	4.0	10.2	1.75	5.3	1.80	11.6	81.50
	1271 •	614.	0.07		1008.95	22.0	9.9	9.2	3.7	9.7	1.61	5.0	1.65	11.4	77.00
11	128C ·	616. 623.	J.08		1009.14	21.7	9.3 8.7	8.7	3.4	9 • 1 8 • 5	1.46	4.0	1.50	11.3	72.04
13	1297.	627.	0.09	-0.41	1009.50	21.2	8.2	7.6	2.9	8.0	1.18	4.5	1.36	11.2	64.31
14	1304.	630.	0.10	-0.42	1009.64	21.0	7.7	7.2	2.7	7.5	1.07	401	1.10	10.7	60 • 7 B
15	1311.	634.	0.10		1009.76	20.8	7.2	6.7	2.5	7 - 1	0.47	3.9	1.00	10.8	57.36
16 17	1318.	637. 640.	0.11		1009.88	20.7	6.8	6.3	2.4	6.6	0 . 87	3.7	0.90	10.7	53.98
18	1332.	644.	0.12		1010.00	20.3	6.3 5.9	5.9	2.0	6 • 2 5 • 7	0.78	3.5 3.5	0.80	10.5	50.65 47.37
19	1339.	647.	0.12	-0.47	1010.20	20.2	5.4	5.1	1.9	5.3	0.61	3.1	0.62	10.5	44.15
	1347.	650.	0.13	-0.48	1010.30	20.1	5.0	4.7	1.7	4.9	0.54	y	0.54	13.4	40.02
21	1354+	653.	6-13		1010.40	20.0	4 . 6	4 . 3	1.5	4.5	0.45	4 . 7	U . 46	10.3	37.50
22	1361.	656.	0.14		1010.48 1010.56	19.8	3.8	3.9	1 • 4	4 • 1	ظ و م ن د د	>	€ و د	10.5	34.44
24	1376.	663.	0.15		1610.56	19.6	3.4	3.5	1.2	3.7 3.3	0.32	2.3	• و و ∪ طانون	10.2	21.4.
	1384 •	666.	0.15		1010.69	19.6	3.1	2.5	1.0	3.0	3.22	1.9	0.23	1	45476
															•

WAVE ENERGY IN THE BREAKER ZONE # 0.417E 10 JOULES

TOTAL LONG-SHORE CURRENT ENERGY # 0.273E TO JOULES

TOTAL POSITIVE LONG-SHORE CURRENT ENERGY = 3.2696 10 JOULES

TOTAL NEGATIVE LONG-SHORE CURRENT ENERGY . 0.4375 08 DOULES

HOLLAND: MICHIGAN

RUN BEGINS AT HOUR 13 ON JULY: 18, 1970

STORM - BAROMETRIC PRESSURE AT CENTER OF LOW = 1004.0 MILLIBARS
PRESSURE AT LARGEST ENCIRCLING ISOBAR = 1012.0 MILLIBARS
MAXIMUM PRESSURE INCLUDED IN STORM = 1013.1 MILLIBARS

LENGTH OF MAJOR HALF AXIS = 550.0 KILOMETERS
LETGH OF MINOR HALF AXIS = 300.0 KILOMETERS
ORIENTATION OF MAJOR AXIS = 28.0 DEGREES FROM NORTH

LONGSHORE CURRENT EQUATION FROM FOX AND DAVIS. 1972

SHORE - POSITION COORDINATES - X = 1015.0 Y = 560.0 KILOMETERS

SHORE LATITUDE . 42. ONSHORE AZIMUTH . 90. DEGREES
NEARSHURE SLOPE . 0.033 AVERAGE FETCH . 150. KILOMETERS

HOUR	×	Y	X1	Y1	BARO. PRESS.	WIND ANGLE	SURF.	ONSH WIND	ALSH WIND	EFFECT.	WAVE H	WAVE T	H BŘ	EAKER Angle	LSC VELOC.
	KM	KM	RAD	RAD	мв	DEG	4/5	M/S	M/S	M/5	M	SEC	M	DEG	CM/SEC
13	796.	705.	0.17	0.26	1011.16	202.6	4.7	-4.4	-1.8	1.7	0.03	0.7	د ٥٠٥	-11.8	-11.98
14	805	704.	0.17		1010.95	202.9	5.1	-4.6	-1.9	1.8	0.05	0.9	0.05		-14.43
15	815.	703.	0.17	0.24	1010.73	203.3	5 - 4	-4.9	-2.1	1.9	0.06	1.0	0.00		-16.15
16	824.	702.	0.17		1010.50	203.7	5.7	-5.2	-2.3	2.0	0.07	1.1	0.07		-17.72
17	833.	751.	0.17		1010.26	204.1	6.0	-5.5	-2.4	2 • 1	0.08	1 - 1	0.09		-19.14
18	843.	700.	0.16		1010.00	204.5	6.3	-5.7	-2.6	2 • 2	0.09	1.2	0.10		-20.50
19	852 •	699.	0.16		1009.74	205.0	6.5	-5.9	-2.7	2 • 3	0.11	1.3	0.11		-21.83 -22.71
20	855.	697.	0.16		1009.62	205.1	6.6	-6.0	-2.8	2.4	0.11	1.3	0.12		-23.46
21	858.	696.	0.16		1009.51	205.2	6.7	-6.1	-2.8	2.4	0.12	1.4	0.13	-13.0	-24.12
2? 23	861 ·	694.	0.16		1009.39	205.3	6.8	-6.2 -6.2	-2.9	2.5	0.14	1.5	0.14		-24.71
24	868.	691.	0.15		1009.15	205.5	7.0	-6.3	-3.0	2.5	0.14	1.5	0.15		-25.25
i	871.	689.	0.15		1009.02	205.6	7.1	-6.4	-3.0	2.6	0.15	1.5	0.15		-25.76
ž	885 .	689.	0.15		1008.59	206.7	7.3	-6.5	-3.3	2.7	0.15	1.6	0.16		-27.04
3	899.	688.	0.15		1008.15	208.0	7.4	-6.6	-3.5	2.7	0.16	1.6	0.16	-14.3	-28.38
4	913.	656.	0.15		1007.71	209.5	7.5	-6.5	-3.7	2.0	0.17	1.6	0.17		-29.13
5	928.	68 %	0.15	0.10	1007.28	211.4	7.4	-6.3	-3.9	2 • 8	0.18	1.7	U+18		-51.00
6	942.	66 .	0.15		1006.86	213.8	7.3	-6.0	-4.0	2.8	0.18	2 . 7	0.18		-32.0B
7	956.	686.	0.15		1006.46	216.7	7.0	-5.6	-4.2	2 • 8	0.18	1 • 7	0.18		-32.84
â	966.	683.	3.14		1006.14	218.9	6.7	-5 . 2	-4.2	2.7	0.19	1.7	0.16		-32.92
¥	975.	679.	0.14		1005.84	221.5	6.3	-4.7	-4.2	2.6	0.19	1 • 8	0.18		-32.49
10	985.	676.	0.14		1005.56	224.9	5.9	-4 . 2	-4.2	2.5	0.18	1.7	0.17		-31.13
11	995.	673.	0.13		1005.31	229.1	5.4	-3.5	-4.1	2 • 4	0.16	1.7	0.15	-22.0	-28.97 -25.95
12	1004.	666.	0.13		1005.08	234.7	4.9	-2.8	-4.0 -3.9	2.3	0.15	1.5	0.11	-24.7	-21.81
13 14	10144	663.	0.12		1004.89	242.2 256.9	3.9	-0.8	-3.8	2.2	0.14	1.5	0.07	-30.3	-14-07
15	1041.	661.	0.12		1004.59	276.3	4.1	0.4	-4.0	2.6	0.15	1.5	0.05	-31.7	-9.86
16	1055.	658.	0.11		1004.54	296.5	4.3	1.9	-3.8	3.5	0.18	1.7	6.13		-24.46
17	1068	655.	0.11		1004.55	319.7	4.2	3.2	-2.7	3.9	0.21	1.8	0.20		-30.22
. 8	1081.	653.	0.11		1004.63	340.5	4 . 8	4.5	-1.6	4.7	0.27	2.0	0.28	-10.2	-29.15
19	1095.	650.	0.10	-0.09	1004.78	355.3	5.7	5.7	-0.4	5.7	U.35	2 . 3	0.36	-2.4	-6.80
20	1106.	649.	0.10		1004.95	3.0	6.6	6.6	0.3	6.6	0.44	2.0	U + 46	1.6	6.64
21	1117.	648.	0.10	-0.12	1005.16	8.7	7.5	7.4	1.1	7.4	0.55	2 • 9	0.57	4.0	40.04
22	1127.	646.	0.10		1005.39	12.9	8.3	8 • 1	1.8	8 . 2	0.67	3 . 2	0.64	6.8	32.79
23	1138.	645.	0.10		1005.66	16.2	9.1	5.7	2.5	9.0	0.79	3 • 5	0.82	8.5	43.37
24	1149.	644.	0.10		1005.95	18.8	9 • 8	9 • 3	3.1	9.6	0.92	3 • 8	0.95	9.9	52.5V
1	1160.	643.	0.10		1006-27	20.8	10.4	9 • 7	3.7	10.2	1.05	4.0		10.9	
2	1166.	646 •	0.10	-0.18	1006.44	20.9	10.6	9.9	3.7	10.5	1.16	4.2	1.19	10.9	63.53
4	1178.	653.	0.11		1006.78	20.9	10.7	10.2	3.9	10.6	1.36	4.0	1.40	10.9	58.92
3	1184.	656.	0.11		1006.95	20.9	11.0	10.3	3.9	13.8	1.45	4.8	1.40	10.9	71.05
6	1190.	660.	0.12		1007.13	20.9	11.1	10.3	3.9	10.8	1.53	4.9	1.56	10.9	72.91
7	1196 .	663.	0.12	-0.21	1007.31	21.0	11.2	10.4	4.0	10.9	1.60	5.0	1.64	11.0	74.57
8	1213.	667.	0.12	-0.24	1007.85	22.1	11.3	10.5	4.2	11.0	1.66	5.1	1.70	11.5	78.66
9	1231.	671.	0.13		1008.39	23.0	11.2	10.3	4.4	10.9	1.72	5.2	1.75	12.0	62.20
10	1248.	674.	0.13		1006.93	23.7	10.9	10.0	4.4	10.6	1.76	5.3	1.79	12.3	84.82
11	1265.	678.	0.14		1009.45	24.4	10.4	9.5	4 • 3	10.1	1.78	5.3	1001	15.0	86.55
12	1283.	682.	0-14		1009.94	24.9	9.8	8.9	4 • 1	9.5	1.64	5 • 1	1 467	12.9	84-11
13 14	1300.	686. 6 94.	0.15		1010.40	25.4	9.1 8.5	8 • 2 7 • 7	3.9 3.6	8 • 8 8 • 3	1.47	4.6	1.49	13.1	80.30 76.26
15	1312.	702.	0.17		1010.67	25.0 24.7	8.0	7.2	3.3	7.7	1.33	4.6	1.23	12.9	71.60
16	1335.	710.	0.18		1010.92	24.5	7.4	6.7	3.1	7.2	1.06	4.1	1.08	12.0	66.73
17	134?	719.	0.19		1011.39	24.2	6.9	6.2	2.8	6.6	0.93	3.6	0.94	12.5	61.79
16	1359.	727.	0.20	-0.41	1011.60	24.0	6.3	5.7	2.5	6.1	0.79	3.6	0.81	12.3	56.84
19	1371.	735.	0.21		1011.79	23.7	5.7	5.2	2.3	5.6	0.67	3 . 3	0.69	12.2	51.94
20	1383.	749	0.22		1011.97	23.0	5.2	4 • 8	2.0	5.0	0.56	3.0	0.57	11.7	40.37
21	1395.	762.	0.24	-0.46	1012.13	22,3	4.6	4.3	1.7	4.5	0.45	2.7	0.46	11.5	40.94
22	1407.	776.	0.26		1012.27	21.7	4.1	3.8	1.5	4.0	0.37	2 • 4	38	11.2	36.15
23	1419.	789.	0.27		1012.40	21.0	3.7	3.4	1.3	3.6	0.30	2 • 2	0.31	10.8	31.75
24	1431.	803.	3.29		1012.52	20.5	3.2	3.0	1.1	3.1	0.23	1.9	0.24	10.5	27.57
1	1443.	816.	0.31		1012.63	19.9	2.8	2.6	0.9	2.7	0.18	1.7	0.19	10.2	23.08
2	1460.	839.	0.33		1012.76	10.0	2.2	2.1	0.7	2.2	0.12	1 • •	0.12	9.6	18.42
3	1478.	861.	0.36		1012.87	17.8	1.7	1.6	0.5	1.7	0.07	1 - 1	0.08	9.0	13.83
5	1495.	884. 907.	0.39		1012.95	16.9	0.9	1.2	0.3	1.2	0.04	0.8	0.04	8.5	9.96
6	1512.	929.	0.44		1013.02	16.0	0.6	0.9	0.2	0.9	0.02	0.4	0.02	8.0 7.6	4.57
7	1547.	952.	0.47		1013.10	14.3	0.4	0.4	0.1	0.4	0.00	0.3	0.00	7.1	2.87
'	- , - , ,		007/			,	0.7	V		0	3.00	0.0	3.00	. • 1	2 4 0 1

WAVE ENERGY IN THE BREAKER ZONE . 0-138E 10 JOULES

TOTAL LONG-SHORE CURRENT ENERGY . 0.965E 09 JOULES

TOTAL POSITIVE LONG-SHORE CURRENT ENERGY . 0.962E 09 JOULES

TOTAL NEGATIVE LONG-SHORE CUPRENT ENERGY - 0.263E 07 JOULES

SHEBOYGAN. WISCONSIN

RUN BEGINS AT HOUR 7 ON JULY 13: 1972

STORM - BAROMETRIC PRESSURE AT CENTER OF LOW - 990.0 MILLIBARS
PRESSURE AT LARGEST ENCIRCLING ISOBAR - 1008.0 MILLIBARS
MAXIMUM PRESSURE INCLUDED IN STORM - 1010.6 MILLIBARS

LENGTH OF MAJOR HALF AXIS = 1750.0 KILOMETERS
LENGTH OF MINOR HALF AXIS = 25.0 DEGREES FROM NORTH

LONGSHORE CURRENT EQUATION FROM FOX AND DAVIS. 1972

SHORE - POSITION COORDINATES - X = 1892+0 Y = 1112+0 KILOMETERS

SHORE LATITUDE • 42. ONSHORE AZIMUTH • 270. DEGREES NEARSHORE SLOPE • 0.030 AVERAGE FETCH • 200. KILOMETERS

HOUR	. x	4	×1	71	BARO. PRESS.	WIND	SURF. WIND	mano Uniw	AL SH WIND	EFFECT.	WAVE H	WAVE T	6R H	EAKER Angle	LSC VELOC.
	KM	KM	RAD	RAD	Mø	DEG	M/S	M/S	M/S	M/\$	•	SEC	M	DEG	CM/SEC
7	585.	1794.	-0.25	-0.49	1010.53	49.2	0.1	0.1	0.1	0.1	0.00	0.1	0.00	22.9	1.74
8	621.	1793.	-0.25	-0.48	1010.51	49.3	0.2	0.1	0.1	0.2	0.00	0.1	0.00	23.0	2.17
9	657.	1792.	-0.25	-0.47	1010.49	49.4	0.2	0.1	0.2	0.2	0.00	0.1	0.00	23.1	2.68
10	692.	1791.	-0.25	-0.45	1010.46	49.6	0.3	0.2	0.2	0.3	0.00	0.2	0.00	23.3	3.27
11	728.	1790.	-0.25	-0,44	1010.42	49.7	0.4	0.2	0.3	0.3	0.00	0.2	0.00	23.4	3.96
12	764.	1789.	-0.25	-0.42	1010.37	49.8	0.5	0.3	0.4	0.4	0.00	0+2	0.00	23.6	4.75
13	800.	1768.	-0.25	-0.41	1010.32	50.0	0.6	0.4	0.4	0.5	0.00	0.3	0.00	23.7	5.65
14	842.	1788.	-0.25	-0.39	1010.24	50.2	0.8	0.5	0.6	0.7	0.01	0.4	0.00	23.9	6.80
15	884.	1788.	-0.25	-0.38	1010-14	50.4	0.9	0.6	0.7	0 - 8	0.01	0.5	0.01	24.0	8.13
16	926.	1788.	-0.25	-0.36	1010.02	50.6	1.2	0.7	0.9	1.0	0.02	0.5	0.01	24.1	9.64
17	969.	1788.	-0.25	-0.35	1009.87	50.8	1.4	0.9	1.1	1 • 2	0.03	0.6	0.02	24.2	11.32
16	1011.	1788.	-0-25	-0,33	1009.70	51.1	1.7	1 • 1	1.3	1.5	0.04	0.8	0.03	24.3	13.15
19	1053.	1788.	-0.25		1009.49	51.3	2.0	1.3	1.6	1+8	0.05	0.9	0.04	24.4	15.22
20	1085.	1790.	-0-25		1009.31	51.5	2 • 3	1.4	1.8	2.0	0.07	1.0	0.06	24.5	17.10
21	1116.	1792.	+0.25		1009+12	51.8	2.6	1.6	2.0	2 • 2	0.08	1 - 1	0.07	24.6	18.96
22	1147.	1794.	-0.25		1008.91	52.0	2.9	1.8	2.3	2 • 5	0.10	1.3	0.09	24.7	20.85
23	1179.	1796.	-0.26		1008.67	52.2	3.2	1.9	4.5	2 + 8	0.12	1 • 4	0.10	24.0	22.79
24	1210.	1798 •	-0.26		1008.41	52.5	3.5	2 • 1	2.8	3 • 1	0.15	1.5	0.13	24.9	24.77
1	1242.	1800.	-0.26		1008.12	52.7	3.9	2.3	3.1	3.4	0.17	1.6	0.15	25.0	56.80
2	1280 •	1806.	-0.26		1007.76	53.1	4.3	2 • 6	3.4	3 • 7	0.20	1.8	0.17	25.1	28.91
3	1318.	1813.	-0.26		1007.38	53.5	4.7	2 . 8	3 . 8	4 • 1	0.24	1.9	0.20	25.3	31.08
•	1393.	1819.	-0.26		1006.96	53.8	5.2	3.0	4.2	4.4	0+28	2+1	0.24	25.4	33.27
-	1431.	1826.	-0.27		1006.51	54.3	5.6	3.2	4.5	4 • 8	0.32	2.2	0.21	25.5	35.44
7	1469.	1832.	-0.27			54.7	6.0	3.4	4.9	5.2	0.37	2.4	0 - 31	25.7	37.57
8	1505	1862.	-0.28		1005.53	55.1 55.7	6.4	3.6	5.3	5 • 5 5 • 7	0.41	2.5	U - 35	25.8 26.0	39.63 41.30
Š	1541	1885.	-0.29		1004.92	56.2	6.8	3.8	5 • 5 5 • 7	5.8	0.50	2.8	0.41		42.70
10	1576 .	1907.	-0.30		1004.62	56.8	7.0	3.6	5.8	5.9	0.54	2.9	0.44	26.2	43.87
ii	1612.	1930.	-0.31		1004.33	57.3	7.1	8 • د	6.0	6.0	0.58	3.0	0.47	26.5	44.84
12	1648.	1953.	-0.32		1004.05	57.9	7.3	3 · B	6.1	6.1	0.61	3.1	0.49	26.7	45.62
13	1684.	1976.	-0.32		1003.79	58.4	7.3	3.8	6.3	6.2	0.64	3.2	0.51	26.8	46.22
14	1702 .	1994.	-0.33		1003.76	58.8	7.3	3 . 8	6.3	6.1	0.67	3.2	0.53	26.9	46.82
15	1721.	2013.	-0.34		1003.73	59.1	7.3	3 • 7	6.3	6.1	0.69	3.3	0.55	27.0	47.25
16	1739.	2031.	-0.35	-0.05	1003.72	59.4	7.3	3.7	6+3	6.1	0.70	3.3	0.56	27.1	41.54
17	1757.	2050.	-0.35		1003.71	59.7	7.2	3.6	6.2	6.0	0.72	3 . 4	0.57	27.1	47.70
18	1776.	2068.	-0.36	-0.04	1003.71	60.0	7.2	3.6	6.2	6.0	0.73	3.4	0.58	27.2	47.75
19	1794.	2087.	-0.37	-0.03	1003.72	60.4	7.1	3.5	6.2	5.9	0 . 74	3 . 4	0.58	27.3	47.70
20	1811.	2108.	-0.37	-0.03	1003.78	60.7	7.1	5 . 4	0.1	5.8	0.75	3.5	0.58	27.3	47.57
21	1829.	2128.	-0.38	-0.02	1003.84	61.0	7.0	3 . 4	6.1	5.8	0.75	3.5	0.58	27.4	47.34
22	1846.	2149.	-0.39	-0.01	1003.92	61.3	6.9	3.3	6.0	5.7	0.75	3.5	0.58	27.4	47.04
23	1863.	2169.	-0.40		1003.99	61.7	6.8	3.2	6.0	5.6	0.75	3.5	0.58	27.4	46.67
24	1861.	2190.	-0.41		1004.08	65.0	6.7	3 4 1	5.9	5.5	0.75	3.5	0.58	27.4	46.22
1	1898.	2210.	-0.41		1004-17	62.3	6.6	3.0	5.8	5.4	0 - 74	3.5	0.57	27.5	45.00
3	1920	2231.	-0.42		1004.22	62.8	5 + 5	2.9	5.8	5 • 3	0.12	3 • 4	0.55	27.6	44.57
3	1941.	₹251.	-0.43		1004.28	63.3	6.4	2.9	5.7	5 • 2	0 . 70	3.4	0.53	27.7	43.41
	1963.	2272.	-0.44		1004.34	63.8	0.3	2 . 8	5.7	5-1	0.68	3 • 3	0.51	27.8	42.19
5	1985.	2293.	-0.44		1004.42	64.3	6.2	2.7	5.6	5.0	0.65	3.3	0.49	27.9	40.91
6	2006.	2313.	-0.45		1004.50	64.9	6.1	2.5	5.5	4.9	0.63	3 • 2	0.46	28.0	39.55
7	2028.	2334.	-0.46	0.05	1004.60	65.6	6.0	2 . 4	5.4	4 - 8	0.60	3 . 2	0.44	20.1	18.11

WAVE ENERGY IN THE BREAKER ZONE . 0.168E 09 JOULES

TOTAL LONG-SHORE CURRENT ENERGY = 0-117E Q9 JOULES

TOTAL POSITIVE LONG-SHORE CURRENT ENERGY = 0.117E 09 JOULES

TOTAL NEGATIVE LONG-SHORE CURRENT ENERGY . 0.000E 30 JOULES

SHEBOYGAN. WISCONSIN

RUN BEGINS AT HOUR 7 ON JULY 16: 1972

STORM - BAROMETRIC PRESSURE AT CENTER OF LOW = 1002.0 MILLIBARS
PRESSURE AT LARGEST ENCIRCLING ISOBAR = 1016.0 MILLIBARS
MAXIMUM PRESSURE INCLUDED IN STORM = 1018.0 MILLIBARS

LENGTH OF MAJOR HALF AXIS = 1050-0 KILOMETERS
LENGTH OF MINOR HALF AXIS = 700-0 KILOMETERS
ORIENTATION OF MAJOR AXIS = 20-0 DEGREES FROM NORTH

LONGSHORE CURRENT EQUATION FROM FOX AND DAVIS: 1972

SHORE - POSITION COORDINATES - X = 1892.0 Y = 1112.0 KILOMETERS

SHORE LATITUDE - 42. CNSHORE AZIMUTH - 270. DEGREES NEARSHORE SLOPE - 0.030 AVERAGE FETCH - 200. KILDMETERS

HOUR	x	Y	X 1	71	BARO. PRESS.	WIND	SURF. WIND	ONSH UND	ALSH WIND	EFFECT.	WAVE H	WAVE T	ВR Н	EAKER ANGLE	LSC VELOC.
	KM	KM	RAD	RAD	мв	DEG	M/S	M/S	M/S	M/S	M	SEC	м	DEG	CM/SEC
7	176 •	286.	0.52	-1.08	1018.02	29.6	0.0	0.0	0.0	0.0	0.00	0.0	0.00	0.0	0.00
8	212.	296.	0.51		1018.02	29.5	0.0	0.0	0.0	0.0	0.00	0.0	0.00	0.0	0.00
9	247.	306 •	0.51	-1.04	1018-02	29.3	0.0	0.0	0.0	0.0	0.00	0.0	0.00	0.0	0.00
10	283.	315.	0.50	-1.02	1018.02	29.1	0.0	0.0	0.0	0.0	0.00	0.0	0.00	0.0	0.00
11	319.	325.	0.49	-0.99	1018.02	29.0	0.0	0.0	0.0	0.0	0.00	0.0	0.00	0.0	0.00
12	354.	335.	0.49		1018.02	28.8	0.0	0.0	0.0	0.0	0.00	0.0	0.00	0.0	0.00
13	390.	345.	0.48		1010.02	28.6	0.0	0.0	0.0	0.0	0.00	0.0	0.00	0.0	0.00
14	425.	356.	0.48		1018.02	28.4	9.0	0.0	0.0	0.0	0.00	0.0	0.00	0.0	0.00
15	459.	367.	0.47		1018.02	28.3	0.0	0.0	0.0	0.0	0.00	0.0	0.00	0.0	0.00
16	494.	377.	0.46		1018.02	28.1	0.0	0.0	U.O	0.0	0.00	0.0	0.00	0.0	0.00
17	529 •	388.	0.45		1018.02	27.9	0.0	0.0	0.0	0.0	0.00	0.0	0.00	0.0	0.00
18	563.	399.	0.45		1018.02	27.7	0.0	0.0	0.0	0.0	0.00	0.0	3.00	0.0	0.00
19 20	598.	410	0.44		1018.01	27.5	0.0	0.0	0.0	0.0	0.00	0.0	0.00	0.0	0.00
21	630 •	424.	0.43		1018.01	27.4	0.0	0.0	0.0	0.0	0.00	0.0	0.00	0.0	0.00
22	663.	438.	0.42		1018.00	27.1	0.1		0.0	0.1	0.00	0.0	0.00	0.0	0.00
23	728.	~>2.	0.41		1017.97	27.0	0.2	0.1	0.0	0.1	0.00	0.1	0.00	13.7	1.82
24	760.	480.	0.46		1017.94	26.9	0.2	0.2	0.1	0 • 2	0.00	0.1	0.00	13.6	4.59
1	793.	494	3.39		1017.91	26.8	0.3	0.3	0.1	0.3	0.00	0.2	0.00	13.6	3.49
2	841.	518.	0.37		1017.83	26.7	0.5	0.5	0.2	0.5	0.00	0.3	0.00	13.7	5.09
3	888.	542.	0.36		1017.72	26.5	0.8	0.7	0.3	0.8	0.01	0.4	0.01	13.7	7.13
4	936.	565.	0.34		1017.55	26.4	1.2	1.1	0.5	1.2	0.02	0.6	0.02	13.0	9.62
5	984.	589.	0.33		1017.33	26.3	1.7	1.5	0.7	1.6	0.04	0.8	0.04	13.6	12.58
6	1031.	613.	0.31	-0.54	1017.03	26.1	2.2	2.0	1.0	2 • 2	0.06	1.0	J.06	13.5	15.97
7	1079.	637.	0.30	-0.51	1016.64	26.0	2.9	2.6	1.3	∡ • 8	0.10	1.2	0.10	13.5	19.75
8	1130.	689.	0.26	~C.48	1016.07	27.0	3.8	3.4	1.7	3.7	0.15	4.5	0.15	14.0	24.65
9	1102.	740.	0.23		1015.37	28.2	4 . 8	4 • 2	2 • 2	4.6	0.22	1.8	0.22	14.6	30.55
10	1233.	792 •	0.20		1014.53	29.5	5.7	5.0	2.8	5 • 5	0.30	5 + 5	0.30	15.4	36.64
11	1284.	844.	0.17		1013.58	31.0	6.7	5.7	3.4	6.4	0.40	2.5	0.40	10.0	42.93
12	1336.	895+	0.13		1012.51	32.0	7.6	6.3	4 - 1	1.2	0.51	2.8	0.50	10.8	49.12
13	1387.	947. 989.	0.10		1011-38	34.9	8.3	6.8	4.7	7.8	0.61	3.1	0.60	17.8	59.59
14 15	1456.	1030.	0.07		1010.61	37.C 39.4	8.7 9.7	6.9	5.2	9 • 1 8 • 3	0.71	3.5 3.5	0.76	18.7	63.31
10	1490.	1072.	6.00		1009.11	42.3	9.2		6.1	8.4	0.88	3.7	0.82	20.9	65.93
17	1524.	1113.	-0.30		1008.41	45.0	9.3	6.6	0.0	5	0.94	3.8	0.87	22.1	67.20
18	1559.	1155.	-0.02		1007.77	48.0	9.3	6.2	6.9	8.3	1.00	3.9	0.89	43.3	67.34
19	1593.	1196.	-0.05		1007.20	51.0	9.2	5.8	7.2	8.1	1.04	4.0	0.91	4403	66.36
20	1636.	1233.	-0.07		1006.44	53.8	3.0	5.3	7.2	7.8	1.06	4.1	0.70	25.5	64.55
21	.680	1270.	-0.10		1005.79	56.4	8.5	4.7	7.2	7.3	1.07	4.1	0.89	26.	04.01
22	17234	1307.	-0.12	-0.10	1005.28	59.0	8.2	4.2	7.0	6-8	0.98	3.9	J.78	26.9	56.33
23	1766.	1343.	-0.:-		1004.92	61.6	7.7	3 . 6	6 + 8	6 • 3	0.84	3 . 7	0.67	27.6	50.14
24	1810.	1340.	-C-17		1004.73	65.1	7.2	3.0	6.5	> · B	0.75	3.5	0.55	48.5	43.14
1	18:3.	1417.	-0.19		1004.71	70.9	6.6	2 • 1	6.2	5	0.63	3.2	0.40	24.0	31.34
2	1882.	1504.	-0.24		1005.80	77.3	6.6	1 • 4	6.5	4.9	0.61	3.2	0.32	30.5	c4.02
3	1911.	1590.	-0.30		1007.09	83.1	6.5	0.7	6.4	4.5	0.55	3.0	0.22	30.8	17.92
5	1940.	1764	-0.35		1006.52	87.5 90.9	6.4	0.2	6.2	4.2	0.48	5 • 8	0.11	30.9	9.95
6	1999.	1850	-0.44		1010.00	93.6	5.8	-0.0	5.6	3 • 8	0.40	2.4	0.04	30.4	10.03
7	2026.	1937.	-0.52		1012.87	95.8	4.0	-0.4	5 • 2 4 • 6	3.4 2.9	0.24	2.4	0.09	30.1	11.15
9	2037.	1986.	-0.55		1013.65	45.8	4.3	-0.4	4.2	2.7	0.20	1.9	0.07	29.9	10.35
9	2045	2039.	-0.58		1014.37	95.9	3.8	-0.4	3.8	2.4	0.17	1.8	J.U6	29.8	7.45
10	2054.	2040.	-0.62		1015.03	95.9	3.4	-0.5	3.4	2.1	0.13	1.6	0.05	29.6	8.44
ii	2063.	2141.	-0.65		1015.61	95.9	3.0	-0.3	3.0	1.9	0.10	1.4	U.03	29.3	7.4.
12	2071.	2192.	-0.68		1016.12	95.9	2.5	-0.2	2.5	1.0	0.07	1.2	0.02	29	6.34
13	2080.	2243.	-0.71		1010.56	96.0	2.1	-0.2	2.1	1.3	0.05	1.0	0.02	26.9	5.30
14	2099.	2268.	-3.74		1010.69	97.0	2.7	-0.2	1.7	1.1	0.03	0.8	0.01	28.7	4.70
15	2117.	2334.	-0.71	0.14	1017.16	98.0	1.4	-0.1	1.4	0.8	0.02	0.0	0.00	28.6	4.01
16	2135.	2379.	~0.80	0.15	1017.39	98.9	1.1	-0.1	1.1	0.6	0.01	U . 5	0.00	28.4	3.30
17	2154.	2424.	-0.83		1017.58	99.7	0.8	-0.1	0.8	0.5	0.00	0.4	0.00	28.3	2.61
1.0	2172.	2470.	-0.84		1017.72	100.5	0.6	-0.1	0.6	0.3	0.00	0.2	0.00	20.2	1.98
19	2191.	2515.	-0.89	0.18	1017.83	101.2	0.4	-0.0	0.4	0.2	0.00	0 • 2	0.00	28.2	1.46

MAVE ENERGY IN THE BREAKER ZONE . 0.224E 09 JOULES

^{*}CTAL LONG-SHORE CURRENT ENERGY = 0.195E 09 JOULES

TOTAL POSTTIVE LUNG-SHORE CURRENT ENERGY = 0.195E OF JOULES

TO AL NEGATIVE LONG-SHORE CUPHENT ENERGY . 0.300F 00 JOULES

SHEBUTUAN, WISCONSIN

RUN HEGINS AT HOUR 7 ON AUGUST 1. 1972

STURM - BAROMITRIC PRESSURE AT CENTER OF LOW . 1002.0 MILLIBARS
PRESSURE AT LANGEST ENCIRCLING ISOBAR . 1012.0 MILLIBARS
MARIMUM PRESSURE INCLUDED IN STURM . 1013.4 MILLIBARS

LENGTH OF MAJOR HALF AXIS = 1400+U KILOMETERS
LENGTH OF MINOH HALF AXIS = 450+0 KILOMETERS
ORIENTATION OF MAJOR AXIS = 12+0 DEGREES FROM NORTH

LONGSHORE CURRENT EQUATION FROM FOX AND DAVIS: 1972

SHORE - POSITION COORDINATES - x = 1892-0 Y = 1112-0 KILOMETERS

SHORE LATITUDE = 42. ONSHORE AZIMUTH + 270. DEGREES
NEARSHORE SLOPE = 0.030 AVERAGE FETCH = 200. KILOMETERS

HOUR	x	Y	X.	71	BARO. PRESS.	WIND ANGLE	SURF.	ONSH	ALSH Wind	EFFECT.	HAVE	WAVE	8H H	ŁAKEH ANGLE	LEC VELOC.
	KM	KM	RAD	RAD	мв	DEG	M/5	M/5	M/5	M/5	м	SEC	M	DEG	CM/SEC
7	598.	1911.	-0.38	-0.61	1013.45	57.5	0.0	0.0	0.0	0.0	0.00	0.0	0.00	0.0	0.00
	639.	1793.	-0.32	-0.59	1013.45	57.4	0.0	0.0	0.0	0.0	0.00	0.0	0.00	0.0	0.00
9	680.	1675.	-0.24	-0.57	1013.45	57.2	0.0	0.0	0.0	0.0	0.00	0.0	0.00	23.9	0.00
10	721.	1557.	-0.21		1013.44	57.0	0.0	0.0	0.0	0.0	0.00	0.0	0.00	23.6	0.00
11	763.	1439.	-0.15		1013.44	56.8	0.0	0.0	0.0	0.0	0.00	0.0	0.00	23.8	0.00
12	804.	1321.	-0.09		1013.44	56.6	0.0	0.0	0.0	0.0	0.00	0.0	0.00	23.7	0.00
13	845.	1203.	-0.04		1013.44	50.3	0.0	0.0	0.0	0.0	3.00	0.0	0.00	23.6	0.01
14	875.	1273.	-0.07		1013.44	56.5	0.0	0.0	0.0	0.0	0.00	0.0	0.00	23.7	0.01
15	906.	1344.	-0.11		1013.44	56.7	0.0	0.0	0.0	0.0	0.00	0.0	0.00	23.7	0.02
70	934.	1414-	-0.14		1013.44	56.9	0.0	0.0	0.0	0.0	0.00	0.0	0.00	23.6	0.04
17	966.	1484.	-0.17		1013.44	57.1	0.0	C.O	0.0	0.0	0.00	0.0	0.00	23.8	0.04
16	997.	1555.	-0.21		1013.44	57.3	0.0	0.0	0.0	0.0	0.00	0.0	0.00	23.9	0.05
19	1027.	1625.	-0.24		1013.44	57.5	0.0	0.0	0.0	0.0	0.00	0.0	0.00	24.0	0.09
20	1055.	1615.	-0.23		1013.44	57.5	0.0	0.0	0.0	0.0	0.00	0.0	0.00	24.0	0.16
21	1093.	1605.	-0.23		1013.44	57.5	0.0	0.0	0.0	0.0	0.00	0.0	0.00	24.2	0.27
22	1111.	1596.	-0.23		1313.44	57.5	0.0	0.0	0.0	0.0	0.00	0.0	0.00	24.5	0.43
23	11-0-	1586.	-0.22		1013.44	57.6	0.0	3.0	0.0	0.0	0.00	0.0	0.00	24.9	0.69
24	1168.	1576.	-0.22		1013.43	57.6	0.0	0.0	0.1	0.1	0.00	0.0	0.00	25.3	1.07
1	1196.	1566.	-0.21		1013.42	57.6 57.6	0.1	0.0	0.1	0.1	0.00	0.1	0.00	25.6	1.49
3	1221.	1565.	-0.21		1013.39	57.7	0.2	0.1	0.2	0.2	0.00	0.1	0.00	25.9	2.03
- :	1270.	1563.	-0.21		1013.37	57.7	0.3	0.1	0.2	0.2	0.00	0.1	0.00	24.1	2.72
3	1295	1562.	-0.21		1013.34	57.8	0.4	0.2	0.4	0.4	0.00	0.2	0.00	26.3	3.58
6	1320.	1561.	-0.21		1013.30	57.9	0.0	0.3	0.5	0.5	0.00	0.3	0.00	26.4	4.62
7	1345.	1560.	-0.21		1013.25	57.9	0.0	0.4	0.7	0.7	0.01	0.4	0.00	26.5	5.88
8	1396	1574.	-0.22		1013.11	58.1	1.3	0.7	1.1	1.1	0.02	0.5	0.01	26.7	8.44
š	1447.	1588	-0.22		1012.89	58.3	2.0	1.0	1.7	1.7	0.04	0.8	0.03	26.9	11.82
10	1498 .	1602.	-0.23		1012.57	58.6	2.9	1.5	2.5	2.4	0.07	1.1	0.06	27.0	15.91
ii	1549.	1616.	-0.24		1012-11	58.9	4.0	2 - 1	3.5	3.4	0.12	1.4	0.10	27.2	20.56
12	1600.	1630.	-0.24		1011.51	59.3	2.3	2.7	4.6	4.4	0.20	1.7	0.16	27.3	25.53
13	1651.	1644.	-0.25	-0.11	1010.75	59.7	6.5	3.3	5.7	5.5	0.29	2.1	0.23	27.5	30.43
14	1697.	1670.	-0.26	-0.09	1010.05	60.3	7.4	3.6	6.4	6.1	0.38	2.4	0.30	27.6	34.44
15	1744.	1670.	-0.27	-0.07	1009.30	61.1	8.0	3.8	7.0	5.6	0.47	2.7	0.36	27.8	37.47
16	1790.	1722.	-0.29	-0.04	1008.55	62.1	8.1	3.8	7.2	6.7	0.54	2.9	0.41	28.1	39.29
17	1837.	1748 .	-0.30	-0.02	1007.86	63.5	7.9	3.5	7.0	6 • 4	0.59	3.0	0.44	28.4	39.56
18	1883.	1774.	÷0 • ≥ .		1007.29	66.1	7 - 1	2.0	6.5	5.7	0.61	3.1	0.43	28.9	37.66
7.0	1930.	1800.	-0.32	0.01	1006.88	71.5	5.6	1.8	5.5	4 - 4	0.53	3.0	0.34	29.3	30.18
20	1969.	1831.	-0.34	0.03	1006.81	81.6	4.3	0.6	4.3	3.1	0.27	2 • 2	0.12	30.0	14.08
21	2007.	1801.	-C.35		1006.90	97.8	3.0	-C • 4	3.0	1.8	0.10	1.4	0.04	24.5	8.40
22	2046 •	1691.	-0.37		1007.13	118.5	2 • •	-1.1	2.1	1 • 2	0.04	0.9	0.03	25.3	11.08
23	2084.	1922.	-0.38		1007.51	139.4	2 • 4	-1.8	1.0	1.0	0.03	0.7	0.02	18.3	10.97
24	2123.	1952 .	-0.40		1008.01	150.2	2.6	-2.4	1.0	0.9	0.02	0.1	0.05	11.2	8.33
7	2161.	1983.	-0.41		1008.60	107.9	2.9	8	0.6	1.0	0.05	0.	0.02	6.1	5.24
2	2205.	2020.	-0.43		1009.35	176.7	3 • 1	-3.1	0.1	1-0	0.02	0.0	0.03	1.6	1.52
3	2249.	2056.	-0.44		1010.12	182.7	3.1	-3.1	-0.1	1.0	0.03	0.7	0.03	-1-4	-1.28
•	2292.	2093.	-0.46		1010.86	186.9	2.5	-2.9	-0.3	0.9	0.02	0.7	0.01	-3.3	-2.86
5	2336.	2130+	-C-48		1011.52	190.0	2.5	-2.5	-6.4	0.8	0.02	0.	0.02	-4.8	-3.58
٠	2380.	2166.	-0.50		1012.08	192.4	2 • 1	-2.0	-0.4	0.7	0.01	0.5	0.01	-5.9	-3.62
7	2424.	2203.	-0.51	0.25	1012.53	194.2	1 - 6	-i.5	-0.4	0 - 5	0.00	Ų • •	0.01	-6.6	-3.19

WAVE ENERGY IN THE BREAKER ZONE - 0.204E OB JOULES

TOTAL LONG-SHORE CURRENT ENERGY . 0.118E 08 JOULES

TOTAL POSITIVE LUNG-SHORE CURRENT ENERGY - 0.118L 08 JOULES

TOTAL REGATIVE LONG-SHORE CURRENT ENERGY = 0.963E 02 JOULES

MAGDALEN ISLANDS - EAST SIDE

RUN BEGINS AT HOUR 2 ON NOVEMBER 26: 1974

STURM - BARUMETRIC PRESSURE AT CENTER OF LOW - 976.0 MILLIBARS
PRESSURE AT LARMEST ENCINCLING ISCHAM - 1008.0 MILLIBARS
MAXIMUP PRESSURE INCLUDED IN STORM - 1012.6 MILLIBARS

LENGTH OF MAJOR HALF AXIS = 860+0 AILOMETERS
LENGTH OF MINOR HALF AXIS = 860+0 AILOMETERS
ORIENTATION OF MAJOR AXIS = 0+0 DEGREES FROM NORTH

LONGSHORE CURRENT EQUATION FROM FCX AND DAVIS. 1972

SHORE - FUSITION COORDINATES - X + 1665-0 Y + 1110-0 KILOMETERS

SHORE LATITUDE . 48. UNSHORE 42 MUTH . 270. DEGREES NEARSHORE SLUPE . 0.026 AVERAGE FETCH . 200. KILOMETERS

TILES - SPRING TIDE HANGE + 0.80 NEAF TIDE RANGE + 0.97 METERS SLUPE AT LOW TIDE = 0.026 SLUPE AT HIGH TIDE - 0.030

DIURNAL TIDE - FURM NUMBER IS 5.00

HOUR	*	*	хĮ	71	BARO. PRESS.	#IND Angle	SURF.	0N5H	ALSH WIND	EFFECT.	#AVE	#AVE	. 64	IL AKEN AHULL	VELUC.	1151
	KM	k.M	RAD	RAD	мв	DEG	M /5	M/5	4/5	M/5	м	SEC	•	DEG	CM/StC	•
2	1154.	466.	0.49		1006.74	23.6	8.3	7.6	3.3	8.0	0.37	2.5	6.0	12.5	33.44	• •
3	1185.	486.	2.48		1305.86	23.0	9.1	8.4	3.6	8 • 9	0.58	2.9	V • 7 E	12.1	40.73	
4	1206.	507. 527.	0.45		1004.89	22.4	10.0	9.3 10.1	3.8	9.8 10.7	0.76	3 • •	U + 78	11.5	49.40	0.50
6	1232.	947.	0.43		1003.83	21.8	11.8	11.0	4.0	11.5	1.14	4.2	1.10	11.1	24.23	
7	1234.	568	0.42		1001.46	20.2	12.7	11.9	4.4	12.4	1.33	4.5	1.30	10.7	55.35	U.4.
9	13.0.	588.	0.40		1000.16	19.4	13.5	12.7	4.4	13.2	1.53	4.8	1.57	10.4	50.97	
9	1341.	651.	0.35	-0.25	997.02	20.3	15.0	14.1	5.2	14.7	1.76	5.2	1.00	1007	43.00	
10	1373.	714.	0.36	-0.22	993.60	21.6	16.0	14.9	5.9	15.7	2.01	5.5	2.35	11.4	11.05	V.**
11	1404.	71.	0.25	-0.20	990.04	23.2	16.4	15.0	6.4	15.9	2.22	5 • 8	2.26	140.	70.04	U . + b
à.	1436.	840.	0.20	-0.17	986.53	25.5	15.6	14.3	6.8	15.3	2 • 38	6	2 - 4 1		86.84	54
19	1467.	903.	0.16	-0.15	983.27	28.8	14.4	14.6	6.9	13.8	2.46	6.2	2.46	1	44.74	u.e.i
14	1449.	966.	0.11	-0.12	980.48	34.2	12.1	10.0	6.6	14.4	2.30	6.0	4.45	27.2	>0	
15	1508	997.	J • V 8	-0.12 -0.11	979.51	39.5	10.9	8.4	6.9	10.1	1.90	5 • 5	1.8.	14.1	42.20	!
10	1517.	1060	. • C &	-0.13	978.08	46.3 55.3	9.6 8.7	4.9	7.1 7.1	8 • 8 7 • 4	1.48	408	1 - 5 -	45.7	6	
19	1530.	1091.	0.01	-0.10	977.65	66.9	7.8	3.0	7.2	6.2	3.77	1.5	0.55	29.1	40.75	4.72
19	1545	1123.	-3.00	-0.09	977.41	81.2	7.3	1.1	7.2	5.2	0.47		0.21	31.0	100	
20	1554.	1154.	-0.03	-0.08	977.38	96.8	7.2	-0.8	7.1	4.5	0.40	4.5	0.15	3	(:	
21	1559.	1165.	-0.04	-0.08	977.37	102.8	7.2	-1.5	7.0	4.2	U . 37	204	0.20	300-	14.00	
22	1565 .	1176.	-0.65	-0.07	977.39	108.7	1.2	-2.5	6.8	4.0	3.35	2.5	40.4	. 7.7		
2.3	157ۥ	1187.	-0.00	-0.07	971.45	114.5	7.3	-3.0	6.7	3.9	0.33	2.5	0.24		. 0 . 0 4	4 * 'c
24	1576.	1179.	-3.06	-0.06	977.53	120.0	7.5	-3.7	0.5	3.7	0.32	6.0	0.25	. 1.0	4 * * * * *	:
1	1561.	1210.	-0.0!	-0.06	977.64	125.2	7 + 6	-4.5	6.3	s • ?	0.31	2 • 2	0.20	25.3	364	
2	1537.	1221•	-0.08	-0.06	911.77	130.1	0 + 1	-5.2	6.2	3.6	0.31	2.2	0.27	41.0	11.50	• * •
2	1545	1215.	-0.48	-0.0:	977.53	132.1	7.5	-5.0	5.6	3 - 3	0.40	2 • 1	0.24	44.0	1	
3	1615.	1204.	-0.07	-0.03	977.11	134.5	7.0	-4.9	5.J	2.7	0.22	1.7	U-20	21.	46.58	• • 5
,	1624.	1199.	-3.06	-3.03	976.93	140.6	5.7	-4.0	3.0	2 • 4	0.15	1.6	0 . 1 4	19.	£ 4 + 1/4	
÷	.534.	1193.	-0.30	-0.02	976.77	144.6	5.4	-4.4	3.1	2.1	0.12	1.4	U+11	17.4	41006	
8	1543.	.189.	-0.36	-0.01	976.64	149.4	5.0	-4.3	2.5	1.9	0.09	1.2	0.07	15.4	47.64	
į.	100-	1202.	-0.07	-0.00	976.83	101.8	5 . 0	->.4	1.7	1.9	0.09		Jelu	y . 3	12.90	
10	16 '6 .	1217 .	8	0.00	977.14	171.2	5.5	-6.4	0.9	2.2	0.10	1.3	0.10	4.5	7.14	
11	16-30	1231.	-0.09	0.32	977.49	178.2	7.4	-1.4	0.2	4 . 4	J.11	1.3	Uela	J . Y	4.5	
12	1710.	1245.	-0-10	3.03	971.95	163.4	8 . 4	-4.4	-0.5	4.5	0.1	1 • •	0.14	-:.×	-3.35	. • * .
13	1720.	1260.	-3.11	6.64	, H.50	187.4	9.4	-4.4	-1	9 • 1	0.15	1.07	3.10	~ , , ,	~1.4.	* 6
14	17-3-	12744	-3.14	0.00	979.12	190.6	10.4	-10.2	-1.4	3.5	0.18	1 . 7	0.14		-12	
. 5	1761 -	.28	-3.43	J.07	979.82	193.7	1	-1:00	-2.7	3.4	0.54	1	0.24		-16.44	
16	178.	1301.	-:-14	0.08	980.59 981.44	146.2	12	-11.7 -12.5	-:	4.5	0.29	. • •	نام يون دو و د	-y.		
17	1798.	131	15	3.10	982.34	200.1	13.7	-12.9		4.8	0.27	2 • 1	U a 34	-10.5		
19	1635	1341	17	0.13	983.33	201.6	14.4	-13.4	-5.5	5.1	0.37	204	66.58	-ii.		
20	1854.	1354.	-5-18		984.32	202.9	15.0	-13.2	-5.8	5.4	3.41	2.5	0.44	-1		
21	1869.	1386.	-0.21	0.15	985.96	701.6	15.0	-15		5.6		, ,	3041	-11.5		
22	1844.	1417.	-0.3	0.16	981.68	233.6	16	-15.1	-9.0	5.7	3.49	2.0	Jaha	-1 -0		6 5
23	.8 -8 -	1444.	26	€.16	1 14.44	144.	16.3	-15.4	-5.5	5 . 2	0.53	2 + 4	64.4			•
24	19:3.	1-80.		4.1 *	• • 2 •	. AA	10.5	-15.4	-5.3	5.7	0.56	پ • و	4.5		- ** - 7 -	4.15
1	1920.	1512.	3;	6.2	2 . 49	1	18		•	5.6	10	100				
2	19.3.	1543.	-3.13	0.21	4473	147.9	15=8	- : 5 • .	-4.8	5.5	5 9	3 . 1	war i	~7.1		•
3	1993.	1606.	-2-18	0.25	998.54	. +8.6	14.5	****	-4.0		L+59	3 - 1	4.64.2	*9 **		• **
4	2043.	1009.	+0.43		1001.92	. 19.2	12.4	-11.	-4.0	4.5	0.50	2.4	0.53		• : . • • .	• • • •
6	2093.	1731.	-0.48		1034.77	200.0	7.9	-1.5	-3.4	3.5	3.35 G.2.	4.0	11631			
,	2192.	1857	-0.57		10038.8.	200.0	5.9	-5.6	-2.5	2.8	J.12	1	0.13		- 1 - 0 - 1	
	.242.	1920	-4.52		1010.11	200.6	4.3	=4.0	-1.5	1.5	0.06	1 4 4	د با واد			
									,	,			0 .			•

WAVE INTEREST IN THE BREAKER . INC . 0-178E 10 JOULES

¥

T TAL L MG-SHORE CURPENT ENERGY + 0+152E 17 a GUIS

TOTAL HISTITYE CUNGHISHORE CURRENT ENERGY + C. 150E 10 JOUGES

TOTAL NEGATIVE CONSENSE COMMENT ENERGY & GAZHBE OF COURSE

MAGDALEN ISLANDS - WEST SIDE

RUN BEGINS AT HOUR 2 ON NOVEMBER 26: 1974

STORM - BAROMETRIC PRESSURE AT CENTER OF LOW . 976.0 MILLIBARS
PRESSURE AT LARGEST ENCIRCLING ISOBAM . 1008.0 MILLIBARS
MAXIMUM PRESSURE INCLUDED IN STORM . 1012.0 MILLIBARS

LENGTH OF MAJOR HALF AXIS = 860 = C KILOMETERS LENGTH OF MINOR HALF AXIS = 860=0 KILOMETERS DRIENTATION OF MAJOR AXIS = 0.0 DEGREES FROM NORTH

CONSCHORE CURRENT EQUATION FROM FOR AND DAVIS: 1972

SHURE - POSITION COORDINATES - A = 1865-0 Y = 1110-0 KILOMETERS

JMORE LATITUDE • 48. ONSHORE AZIMUTH • 90. DEGREES NEARSHORE SLOPE • J.026 AVERAGE FETCH • 200. KILOMETERS

TIDES - SPRING TIDE RANGE = 0.86 NEAP TIDE RANGE = 0.37 METERS SLOPE AT LOW TIDE = 0.026 SLOPE AT HIGH TIDE = 0.030

DIURNAL TIDE - FORM NUMBER IS 5.00

430 8	x	۲	×1	٧1	BARO. PRESS.	#1ND ANGLE	SURF.	ONSH	ALSH WIND	EFFECT.	WAVE	WAVE T	8R M	EAKER Anglè	LSC VELUC.	TIDE
	K.99	KM	HAD	RAD	48	DEG	M/5	M/5	M/S	M/5	M	SEC	H	DEG	CM/SEC	
2	1154.	445.	-0.49		1006.74	203.6	8.3	-7.6	-3.3	3.0	0.08	1.1	0.08		-15.90	0.56
3	1180.	480.	-0.48		1005.86	203.0	9.1	-8.4	-3.6	3.3	0.13	1.4	0.14		-18.97	U.53
•	1236 -	507.	-0.46		1004.89	202.4	13.0	-9.3	-3.8	3.6	0.16	1.6	0.17		-21.09	0.50
5	1232.	527.	-0.45		1003.83	201.8	10.9	-10.1	-4.0	3.9	0.20	1.0	0.21		-22.73	0.47
•	1250.	547.	-0.43		1002.69	201.0	11.6	-11.0	-4.2	4.2	0.24	1.9	0.25		-24.05	0.44
7	.284.	568.	-0.42		1001.46	200.2	12.7	-11.9	-4.4	4.5	0.28	5.1	0.29		-45.65	0.42
8	1310.	588.	-2.40		1000.16	199.4	13.5	-12.7	-4.4	4.7	0.32	2.2	0.38		-28.85	0.42
•	1341.	651	-0.35	0.25	997.02	200.3	15.0	-14.9	-5.2	5.3 5.7	0.42	2.0	0.43		-32.31	U.44
10	1373.	714.	-0.30	0.22	990.04	203.2	16.4	-15.0	-6.4	5.9	0.47	2.7	0.48		-36.00	0.48
12	1435.	840.	-0.20	0.17	986.53	205.5	15.8	-14.3	-6.8	5.8	0.51	2.8	0.52		-37.95	U.54
13	1467.	903.	-0.16	0.15	983.27	208.8	14.4	-12.6	-6.9	5.4	0.53	2.9	0.53		-43.86	v. 61
1.	1499.	966.	-0.11	0.12	980.48	214.2	12.1	-10.0	-6.8	9.7	0.50	2.8	0.50		-45.71	0.69
15	1508.	997.	-0.08	0.12	979.51	219.5	10.9	-8.4	-6.9	4.4	0.44	2.7	0.43		-44.47	0.77
lo	1517.	1029	-0.00	3.11	978.71	226.3	9.4	-0.7	-7.1	4.2	0.39	د. ډ	0.30		1.24	U.85
	1520	1960.	-0.03	0.1.	978.08	235.3	8.7	-4.9	-7.1	4.1	0.35	2.3	0.29		-35.20	0.90
1.0	1536 -	1091.	-3.01	0.10	977.65	246.9	7.6	-3.0	-7.2	4.2	0.31	2.4	0.22		-45.94	0.93
19	1945.	1123.	3.00	0.09	977.41	261.2	7.3	-1.1	-7.2	4.5	0.34	2.3	0.15	-31.0	-15.36	4.94
20	1554.	1154.	0.63	0.08	977.38	276.8	7.2	3.6	-7.1	5.1	0.37	2.4	0.14	-31.5	-14.11	0.91
21	1559.	1165	0.04	0.08	977.37	282.8	7.2	1.5	-7.0	5.3	0.41	2.5	0.22	-31.2	0 . 67	U . 86
22	1565.	1176.	0.05	0.07	977.39	288.7	7.2	2.3	-6.B	5.6	0.46	2.7	0.29	-30.4	-26.91	0.80
23	1570.	1187.	3.06	0.07	977.45	294.5	7.3	3.0	-6.7	5.4	0.50	2.0	0.36	-28.9	- 53. 06	0.72
24	1576.	1199.	3.00	0.06	977.53	300.0	7.5	3 • 7	-6.5	6+3	0.55	2.9	0.43	-27.3	-39.05	4.67
1	1581.	1210.	3.37	0.00	971.64	305.2	7.8	4.5	-6.3	6.7	0.61	3.1	0.51		-44.67	0.59
2	1587.	1221.	3.08	J.06	977.77	310.1	8.1	502	-6.2	7.1	0.67	3.2	0.59		-47.67	0.53
3	1596 .	1216.	0.0	0.05	977.53	312.1	7.5	5.0	-5.6	6 • 7	0.71	3.3	0.64		-51.74	0.49
•	1000.	1210.	0.37	0.04	977.31	314.5	7.0	4.9	-5.0	6 • 3	0.73	3.4	0.67		-53.14	0.45
•	1615.	1204.	0.07	0.33	977.11	317.3	0.5	4.7	-4.4	5.9	0.71	3.4	0.67		-52.94	0.43
•	705	1199.	0♠	0.03	976.93	320.6	5.9	4.6	-3.8	5.5	0.63	3 . 2	0.60		-49.87	0.41
7	1634.	2193.	0.00	0.02	976.77	324.6	5.4	***	-3.1	>-1	0.55	3.0	0.54		-46.02	0.40
	1643.	1166.	0.00	0.01	976.64	329.4	9.0	4.3	-2.5	4.7	0.46	5.0	0.48		-41.03	0.59
٧	1000.	1232.	2.07	0.00	y76.83	341.8	5.6	5.4	-1.7	5.5	0.51	3.0	0.53		-30.70	0.40
70	1676.	1217.	3.38	-0.00	977.12	351.2	6.5	6.4	-0.9	6.5	0.57	3.0	0.59	-6.9	-3.81	0.45
11	1693.	1231.	3.09	-0.03	977.49	350.2	7.4	7.4	-0.2	7.4	0.65	3.2	U. 79	1.8	8.14	0.45
12	1710.	1245.	2.10	-0.03	978.50	3.4 7.4	3.4 9.4	9.4	0.5	9.4	0.75	3.4	0.74	3.4	18.96	U.58
13	1726.	1260.	0.11	-0.04	979.12	10.6	10.4	10.2	1.2	10.3	1.02	4.0	1.67	5.0	28.99	V.67
15	1761.	1287.	3.13	-C.07	9 '9 . 82	13.7	11.3	11.0	2.7	11.2	1.18	4.3	1	7.5	39.91	V. 76
10	1780.	1301.	0.1.	-0.08	984.59	16.2	12.4	11.7	3.4	12.0	1.35	4.6	1 . 35	8.6	50.00	0.85
17	1798.	1314.	0.15	-0.10	951.44	18.4	13.0	12.3	4.1	12.8	1.52	4.0	1.57	9.7	54.20	J. 93
10	1017.	1327	0.16	-0.11	982.34	20.1	13.7	12.9	4.7	13.5	1.70	5.1	1.74	10.6	67.60	0.98
13	1835	1341	0.17	-0.13	983.30	21.6	14.4	13.4	5.3	14.1	1.87	5.4	1.71	11.4	74.92	U.99
ن 2	1850.	1354.	5.18	-0.14	964.32	22.9	15.0	13.8	5.8	14.6	2.05	5.6	2.48	12.6	81.15	J. 98
21	1869.	1300.	0.21	-0.15	985.96	21.6	15.6	14.5	5.8	15.3	2.23	5.4	2.21	11.4	80.87	U. 93
22	1884 .	1411	3.43	-0.16	988	20.6	16.1	15.1	5.6	15.0	2.40		6.46	13.8	Bu . 1 7	U. 85
23	1898 -	14474	3.26	-3.10	989.44	19.7	16.3	15.4	5.5	16.0	2.57	6.05	2.04	10.4	79.10	0.77
24	1913.	1480.	26	-0.19	991.22	19.0	16.3	15.4	5.5	16.0	2.71	6.5	2.79	10.0	77.75	U.08
1	1928 .	1512.	3.31	-0.20	945.94	18.4	10.1	15.3	5.1	15.9	4.63	6.6	2.46	9.7	76-15	0.59
2	1943.	15+3+	0.33	-0.21	994.73	17.9	15.0	15.0	4 . 8	15.5	2.43	6.8	3.52	4.4	74.50	0.52
,	1993.	1606.	5.30	-0.25	994,54	10,6	14.3	13.6	4.6	14.1	2.96	6.0	3.05	4.7	70.40	J. 40
•	2043.	1669.	2.43		1001.92	19.2	12.4	11.7	••0	15-1	4.51	D . 5	4 + 5 8	40.0	11.24	0.04
•	2093.	1731.			1004.77	19.7	10.1	9.5	3 . 4	9.9	1.90	5.5	1 + 96		h !	0.40
6	21-2-	.794.	0.53		10:100	20.0	7.9	1.5	2 . !	• 8	1.34	4 . 6	1 . 3 7		52.57	V . 38
7	2192.	1857.	0.57		1004.62	20.4	5.9	2.0	4.0	5 • 8	0.84	3 . 7	0.67		41.67	J. 97
	22-2-	1920.	3.62		1010.11	20.6	4.3	4.0	1.5	4 • 2	0.47	∠ • 8	0.47	.0.4	10.44	4.57

WAVE ENERGY IN THE BREAKER ZUNE + 0.583E TO JOULES

TOTAL CONGESHORE CORRENT ENERGY # 0.2316 11 DOUCES

TOTAL POSITIVE CONUMSHOUSE CURRENT FRENCH - GAZZAE ... JOUCES

PLUM ISCANDA MAUSACHUSETTS

HUN BEGINS AT HOUR IN UN APPRIL HE 1975

JIONN - HARCHETHIC SHESSORE AT CENTER OF COM . SPECIO MICCIBANS
PREESORE AT LANCENT ENCINCETING ISSUEM = 1012-0 MICCIBANS
MAXIMOM SHESSORE INCLUDED IN STORM . 1010-3 MICCIBANS

LENGTH OF MAJOR THAT ARES - 120100 RELOMETERS LENGTH OF MINOR THAT ARES + 60100 RELOMETERS LIBERTAL, VIOL MAJOR ARES + 1540 LEGISES FROM NORTH

CON SHOTE CURRENT EQUATION FROM FOR AND CAVIDS 1972

SHORE - PUSTITION COORDINATES - A . 1980-0 T . 834-0 ALLOMETERS

SHORE LA ITULE + 42. ONSHURE AZIMUTH + 255. DEGREES NEAMSHURE LEUME + 3.0300 ALFRADE FETCH + 4000. KILUMETERS

TIDES - SPRING TIDE HANDS + 3.50 NEAR TIDE MANCE + 2.00 WETERS S. PE AT USE TOLE + USING SCHEAT HIGH TIDE + 2.00

SEMIUL WHAL TILE - FORM NUMBER IN USEL

HÜLR		,	٠.	*1	HESS.	WINE ANGLE	5UMF .	MINC	WIND	EFFECT.	RAVE	***	b#	ANULL	, \$ c veru	1106
	K.M.		RAC	RAD	₩Đ	Dto	M/5	u . 5	M/5	M. 5	•		•	Des	CM SEC	
1	11	139.	2+16	-1.57	1616.35	37.	0.0	3.0			4.00	9.9		0.0		
ž	1 ***	206.	3.13	-1.0-	1016.35	37.4	3.0	J • D	0.0	0.0		0.0	L . JU	5.0	00	1.45
3	3'	278.	J.10	98	1016.34	37.9	0.0	0.0 3.0	0.0	0.0	0.00	0.0	0.00		0.00	
3	444	-10.	V+34		1010.34	38.8	3.3		U.J	U . U		Jec		6.0	0.00	
•	527.	486.	0.12	8.	1016.34	39.4	5•7		3.0	V-9	4.00	0.0	U. ()		U+60	
•	613 •	555.	-2000	-0.17	1016.34	40.0 39.9	0.0	3.0	0.5	0.0	0.00°	0.6	U = 0U			
	770.	573.	-5400	-0.66	1010.33	39.4	2.0	3.0		J.0	0.00	U · U	U + U U	18.0		25
i.	850.	582.	36		1010.30	39.	0.1	0.1	0.1	(· • 1	0.00	0.1	0.00	. 4 . 4		U . 5 %
11	911.	592.	04.4	-0.5	1010.22	39.4	1.0).; J.a	3.6	0 • • 0 • ¥	3.0.	U # #	0.00	19.6	2.90	U + 1 U
11	1011.	601.	0 0.(1		1015.64	39.2	2.0		,	1.9	0.04	U . B	4.04		9	
14			-2.0	-0.46	1015.53	*5.0	2.5	1.5	1.6	2+3	0.47	1	06	2000		• • •
1	11-1-	J V .	-000		1115.34	****	د . و			2-7	0.10	1.2	0.99 0.13		7.47	1.07
17	117	9.	-6.5 -6.65		10.5414	+1.c	3.5			3.2	U-17	1.6				
14	121.	65.	-19	-0.40	1.14.00	43.7	4.5	3 . 1	1.2	4.2	v * 4 d	1.8	40.00		7	
19	1258.	906.	-0+.2	-0.36	1 14.40	44.5	5 . 4	3. /	1.7	• • 1	3.27	• • •		22.1	4	
20 21	1335.	655.	-0.13		1012.80	****	4.00	9.4	5.7 H.B	•• 7	0.78	3.4	20/1	2200	,	
21	1412.	84:			1010.5	*3.5				15.0	1.2		1	1700	5. 5 . 2 5	
23	1500 .	623.		-0	1002.26	+2 . B	41		14.3	14.2	1.8,	500				u. · · .
2 %	.0-3.	7.7	-2.31	• i e	99'151	-1.6	49.4	, , , ,	11.0	42.02	2.44	6.0	4 + 1/2		15.448	0.47
,	1 20	802	-0,00 -0,0		19	-1.1	24.4	16.3	3	14.9	3.13		£ . + -			40.00
•	1763.	826.	-C.	-C.i.	309	42.7	21.1	4	14.5	14.2	3.55	7	****		V0	
•	.705+	0.1	+3 , u s		9 * B • • •	****	23.2		1 + + 2	18.5	3.14	6.7			101.40	3
•	16.7.	676. 970.	-0.04	-2.08	487.55	***	.8.2			10.1					yy. 2/	
	1050	9.			¥ ML	5	17.0	:3.7	4504	7	4				74+" 1	
	1869.	925.		-0.05	983.44	51.0	15.4	4.5	12.1			• • •			61. f.	4001
	1905.	925.	-0.05	-3.04	984.5; 983.67	52.9	13.	8.2	9.0	9	1.46	4.9	1.00			
- 11	.924.	925	-0.07	-3.51	183.33	55.8	4.4		8.4	F . 4	0.95	1.8			****	· · ' ·
12	1942.	425.	-0.05	+0.0¢	982.9.	57.9	٠.9	٠.,	L.,	1.00	0.00	5 4 4	4 5.5	21.		
13	1961.	925.	-0.	0.00	982.35	67.9	3.7	2 - 6	3.4			1.8	0.14		. 5. 27	
	1980.	914.	-0.04 -0.04	0.30	982.25	103.3		-3.3		1.3		3.7	Laus			
16	2017.	90e	-0	5	982.29	151.4	2.0	-1.7		. • *	. د ه د	0.5	0.01	13.0	****	4.
1.7	2035.	930.	-0.32	وبول	942.48	173.2	3.4	-9.1		1.0		U • 7	U + 0 2	-1.3		40.5
10	2054+	111.	-0,.2	2.00	983.27	182.5	3.8	-5.8	-0.7	1.9		L . Y		-4.0	-4,71	4 4 3 4
20	2015	891.	-1.662	3.05	983.35	167.2	0.0	-5.9	-0.1	2.0	7	100	1.4%		,4	
21	2078+		-02	06	943.43	187.0	6	-6 + 1	-0	2.0	UNUN	1 • 4	14 UP	~ 3 · .	-7.14	4102
23	2081.	900.	-0.02	.50	983.52 983.60	186.7	9.4	-6.2	-0.7	2.1	J . U.Y	4 • 4		- 3, 4		
2.	2087.	903.	- 02		941.69	180.3	6.6	-6.5	-0.7	4 . 2	4.10	1.5		- 3 . 3	-4.2.	
1	2090.	9 26 .	-0.02	0.00	983.19	100-1	• •	-6.7	-0.1	2 • 2	0.1.	1.5	****	- '		
?	2109.	912.	-0.02	0.01	985.19	187,6	7.9	-4.8	-1.5	2 • 6 3 • U	3		0.1	-4.5	-6.4.	****
	2144.	9:5.	-0.02	0.09	986.03	189.4	7.9	-4.7	-1.0	3.3	0.17	1.0	16	-5.0	-t . 2 i	
9	2165.	919.	-0.02	C-11	986.94	190.0	10.7	-10-6		3.6	0.20	4 • 7		-5.3	-10.05	1.50
•	2202.	922.	-0.03	0.13	987.93	190.9	11.5	-11.5	-2.3	3.9	0.23	2.0	0.21		-1 '6	4010
- 4	2727	900.	-2.50	14	990.67	193.5	13.1	-12.0	-3.0	4.9	0.24	2	4.5.	- " • .	•	4
i	2451 .	Ø 76 ·	0.00	5.19	992.5.	199.5	13.6	-13-3	- 3 - 0	• •		4.0			-,3,77	1.69
10	2274.	851.	0.02	3.16	994.46	197.0	14.1	-13.5	1		5.36 0.39		4 1	- 7.5	-, 3. 77	4 0
12	2325.	0.2.		4.19	979.48	199.3	4.0	-13.2	-4.6	4.0		2.6		-10.4	*. ** 1.	v + 65
1.3	. 350 .	.,	27		1000.47	200.2	13.5	-1		••1	3	0	U = 45	- 4 . • 5		••••
1.0	43714	70.		0.21	10.2.09	200.4	.2.0	-12-1		***	0.45	2.7	U . 4 t	0.7	~3	
19	2943.		.39		1-1-11		,								*****	••••
1.	1430.	1.30	1.12	2 3	13 (.5)	101.45	1 6			9.5			. + :			
4.0	2457		13	24	1007.78	202.5	9.5		-1.	3.2	3.35	2.5	× + 21	-1 -	*. t . J .	
1,	24.79.		• • •	• • • •	13) .61	2-2-4	7	-6.8	-2.0	2.0	0.7.			-1-17	~. 2 4	
٠.				1.29	1-11-99			-5.4	3		V	1 . 0	1 4	-16.0	-11.41	
22	. 0.00		U-10	4.34	1.101		• • •	-2.4		1.*		,	0.0		-11.63	
23	2680.	91.	2.23		1013.98	262.8			-1 - 4					-10.7		
	2 7 7 7	522.	2 3	1	. 11.15	2000	2.5		*	• •	0.0.	• 5		-16.6	-4437	
	2 33.	50%				2000	1.	-1	-0.4	U • *	0.00	• *	- U U	-11.5	-2,46. -1,197	
	284	, '*.	• • • •		1.15.98	2.2.1		- 16 '						-10.	-1037	
÷		500.				2-1-9		- / -	~~ * ~	- • •		3.0		÷ •		• 16
6	3607.	557.	•• '		1 1445	2 4.0	~ • ·	• . :			* N*	U . V			* 1 * 2 *	***
	3100.	****	0.28	2.5	1.16.1.	4-1-7			-0.0	. • •			0 + 10 t			
•	1211.	211	5.1.	1.66	1016.34	7			-0.			1.0				
4.	39:0.	525.	2.34	6.64	1.14.3-	4	4.0	****	•			• •			* - 4 - 1	
• 1	136 .	51.7	0.19	1.11	1.15.34	20102	3.0					• •		- : - • :		
ندد	14.33		2.15		1.15.34	2.1.4	1.0	-0.0	-			• • • • • • • • • • • • • • • • • • • •		- /	-0.00	• • • • • • • • • • • • • • • • • • • •

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HUR BESTING AT HOUSE IT ON SUCH 14: 1973

STORM - HARDMETRIC PRESSURE AT CENTER OF LOW - 1010-0 MICLIBARS - 4555046 AT LANCEST ENGINEETING ISOBAR - 1016-0 MICLIBARS WAXIMUM PRESSURE INCLUDED IN STORM - 1016-0 MICLIBARS

CENTR OF MAJOR HALF ARIS = 74860 KILOMETERS
CHAITH OF MINOR HALF ARIS = 22460 KILOMETERS
CREENTAITH OF MAJOR ARIS = 40.0 DIGHEES HROW MORTH
LUNGSHORE CURRENT EMUATION FROM CERRES, 1/73

SHORE - PUSITION COORDINATES - A . 2287.0 Y . 1537.0 KILOMETERS

SHORE LATITUDE . 38. ONSHORE AZIMUTH . 285. DEUREES NEARSHUNE SUCHE . 08/037 AVERAGE FETCH . + 5000. KILUMETERS

TIDES - SPMING TIDE RANGE * 1-72 NEAR TICE RANGE * 0.90 METERS SCHE AT LOW TIDE * 0.035 SECHE AT HIGH TIDE * 0.039

MINED SEMICICANAL TIDE - FORM NUMBER IS - DENG

HOU:		٠	×1	*1	HARO. PRESS.	#IND ANGLE	SURF.	0N5H #1ND	AL SH	EFFECT.	WAVE H	#Avt		EAREK ANULI	işt Veruti	1;
		**	нар	HAU	Mf4	DEG	M/5	4/5	w/5	M/5	M	Stv	M	DŁG	CM/SEC	м
7	195.	548.	1.33		1016-87	29.2	0.0	0.5	J . U	0.0	0.00	U. W	0.00	0.0	J.00	
8	253. 311.	597.	1.27		1016.87	29.6	0.0	0.0	0.0	ن• ن	0.00	9.5	0.00	0.0	0.00	4.45
10	368.	696.	1.22		1016.87	30.0 30.3	0.0	0.0	Ú.U	0.0	0.00	Ú.U	U. UU	0.0	J. 30	4.24
11	420.	745.	11		1016.87	30.7	J.U	3.0	U.U	0.0	3.00	0.0	0.00	6.0	y.00	
12	484.	795.	2.05	-1.38	1 14.67	31.1	0.0	0.0	0.0	U+U	0.00	J.U	0.00	Jev	U. UU	
13	542.	844.	0.99		1016.87	31.5	0.0	3.0	0.0		0.00	4.0	U . U1.	J + C	0.00	w.2t
15	672.	984.	0.94 J.89		1016.87	31.8	J.J	3.0	0.3	J.J	0.00	ن ۾ ن نڌ ۾ ن	U.UU	U = U	0.50	U.98
16	737.	979	3.53		1010.87	32.5	0.0	3.0	5.0	U+U	0.00	0.0	0.00	3.0	4.40	1044
17	802 .	.025 •	3.78	-1.10	1010.07	32.8	3.0	0.0	3.0	0.0	0.00	U.0	0.00	3.0	U. U.	4 . 38
1.6	867.	1070.	v = 7 4		1916.87	33.2	0.0	¥ • U	0.0	J = U	0.00	ن ۽ پ	بال و ب		U = U ~	1.37
20	912.	1115.	0.63	-1.06	1016.87	33.6	0.0	J.0	0.0	U . U	0.00	0.0	0.00	3.0	U+00	1-20
21	1362.	1184.	J.58	-0.97	1010.6	34.1	0.0	0.0	0.0	0.0	0.00	0.0	0.00	0.0	0.00 0.00	U.55
22	1125.	1219.	0.54	-0.92	1016.87	34.3	0.0	9.0	0.0	0.0	6.00	0.0	0.00	0.0		
23	1193.	1254+	0.49	-0.87	1016.67	34.6	0.0	0.0	0.0	0.0	3.00	U.U	بان و ∪	U • C	J. U.	-0.01
1	1258.	1288.	3.40		1016.87	34.9	3.C	3.0	0.0	0.0	0 + QU	U . U	0.00	200	0.00	-0.09
- 2	1396.	1355.	3.36	-C.74	1010.87	35.5	3.0	3.0	3.0	0.0	0.00	0.0	J.J.	۔ • د	0.00	-(
3	1468.	1386.	2.31	-0.67	1516.80	35.4	0.0	0.0	U.3	Ų.J	0.00	Ü. ü	J . JU	5.5	0.00	V.43
•	1540.	1417	3.27		1016.86	36.2	0.0	0.0	0.0	0.0	0.00	0.0	0.00	200	0.00	
5	1613.	1480.	0.23		1-16.86	36.6	0.0	3.5	0.0	0.0	3.00	0.0	0.00	10.	U + 1 *	
,	1758.	1512.	3.14	-0.44	1010.05	3 - 7	0.1	0.1	0.0	0.1	4.00	0.1	0.00	10.4	3.50	V. 92
	1785.	1506.	7-14	-0.44	1016.83	37.5	0.9	U.2	0.2		0.00	J.1	V. UU	10.5	0.74	
4	1813.	1499.	3.14	-3.39	1016.83	37.3	0.5	0.5	فيان	0.6	V.U0	ووق	0.00	19.6	12.15	7
12	1840.	1493.	Jel4 		1016.73	31.0	2.0	3.9	0.7	1.1	0.01	0.5	0.01	18.6		
12	1895.	1460.	3.13		1016.43	36.4	3.2	4.5	1.2	1.8	U.U4 U.U9	1.5	9	18.5	4	-0.02
13	.422.	1676.	0.13	-0.29	1016.14	35.0	4.7	3.8	4.7	4.4	0.18	1.6		18.3	28.83	
1.4			4-19	-0.27	1015.73	35.5	>	٠.,	3.0	6 + 1	0.31	2.4	U . 3L	14.4	11.53	U. 18
15	1976.	1461.	0.13	-0.24	1015.19	34,5	8.4	6.9	• •	7.9	0.48	2 - 7	0.47		97.52	v • 72
17	2031.	1449.	0.13	-0.22	1014.51	34.0	10.1	b.4	5.6	4.5	0.68	302	Uahe Uahe		110.44	1
16	2016.	1-42-	3.13		1012.89	32.1	11.4	y.7	5,9	10.0	1.05		1.34		137.44	1.39
19	2009.	1436 .	00.3	-C - 15	1012.08	28.2	10.4	**1	4.4	9.9	1.15	406	1 - 12	. * . 6	132.15	34
20	2111.	1420.	0-13	-0.12	1010.88	22.2	8.3	7.4	٠.،	7 - 8	1.16	4.5	1.14			1.09
21	21000	1411.	0.13	-0.07	1010.00	3-5-6	2.8	1.0	0.5 6.5-	2.4	0.61	1.7	U.65		-08.00	
23	2190.	1403.	2.13	-0.05	1010.75	438.0	3.1		-2.7	1.5	0.07		J. U.		-44.84	****
2 *	2217.	:394.	3.13		1911-13	215.0	5.1	~4 . 1	-3.0	2 . 3	U • QB	1 • 1	J. U. 7	-18.0	- 37. 44	
ž	42/20	1386.	0.13	-C.JU	1-12-29	208.2	7.4	-5.8	-3.1 		0.11	1 • 4	0.09	-14-5	- 15.6 3 7	
í	.302.	1390.	0.12	0.02	1-12-90	202.6		-/-:		2.0	0.13	1 . 5	0.12		- 13.77	0.04
	2331.	1392.	0.11	0.07	1-13-55	20000	7.4	- 1 - 0	• t		0.14	1.5	V	-10.6	- 5 1 . 5 .	
•	2360+	1395 -	0.10		1414-17	149-2	6 . 7	-6.5	-2.2		6.14	1.00	- • • 5	-4.7	-34+4-	
ţ	2390.	1397.	0.09		1014.75	148.3	5.1	-5.7	-1.9	1.0	0.13	1.5	4.15		6. 47	U . V .
	2441.	1435.	0.05		1-15-02	VA.5	5.6	-5.9	-1.5	1.4	J. U.	104	J 4 J Y			. '
9	2463.	14 '00	- i	9.16	1-14-40	175.2	5.9	7	-1.5	2.0	0.09	1		- '. 9		
10	4485 .	1506.	-c.o:		1-14-59	143.8	6	-6.4			3.10	1 . 3	* * * k	- ' • •		
11	2530.	1544.	-0.09		1014.41	192.2	6	-0.4	-1.5	4 • 1		1 • 2				-0.00
13	2552.	40130	2	3.21	1:14.17	168.3	0.4	-0.3	-3.4		11	1.9	• • •			
1.	2564.	.64 .	-0.15	2.21	1013.84	105.0	6 . 5	٠٠.,			7+11			- 1	26	****
15	2575.	166.	-58		1413.64	182.0	6.	6.1	-0.2		0.433			-:•;		× • 5 =
16	4548.	17144			1013.37	178.4	, , ,		200		0 4 2 1 0 4 0 H			•	** **	
10		17e.	-3.26		36.6161	1.4.8					1000				• • •	
; 9	2021.	1814.	-5.31	22	1013.44	153.	3.1	- 5 . 5	. • *		5	1				
2	2632.	1892.			1013.61	11112	1.	*4.	• • •	1 + 2					16.4	
::	4072.	195;			13.9. i - 14.3.	04.0		-0.8	• • •		0 0 0 P			3		
- 23	20134	19 "		2.	1 14.24		4,8	6			0415		• •	1	11.00	
24	26'1.	23.90		6.22	1-15-46	6446	5.~	4 + 3	* * *						4.6.	
	4689.	2044.	-0.56		1 - 11 - 76	61.4	• •	4 = 1	***	• 5				• `• •	****	
,	2674.	4117			16.35	54.,	3.4	• •					• • •		•	
•	2100.	2148.	+6/		16.55	20.3					2009					
*	21.5	2184 .	5		16.12	97.5					0000					
•		2215.				5000	• •	*	• *		3.01	• 5	• •	4.748	. * . * *	
				•••		2.64	• •	- * *	٠,	!			• •		** *	

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TOTAL CONSTRUCTE CHARACT EXEMPT A LATERED A SOURCE

TOTAL POLICE CONSTRUCTOR SWIFTERS CONTROL OF A SAME OF A SAME

I TAL MELATICAL COMPANIES COMMENT THEN THE BOOK TO THE HOLD

SAFELS TOLANDO E MOTA

RUN BEUTNE AT MUCH . . IN FEBUARY 14+ 1969

STIRM - BANGMETHIC PMESSORE AT CENTER OF CON . MOGGO MICCIEANS FAI SSURE AT CANDEST ENCINCUING ISOBAR . 1012-0 MICCIEANS VARIMUM MRESSORE INCLUDED IN STORM . 1015-1 MICCIEANS

LENGTH MAJUH HACE ARES # 9200 RECOMETERS
LENGTH OF MENON HACE ARES # 9200 RECOMETERS
LENGTH OF MENON HACE ARES # 020 DEGREES FROM NORTH

U. NUSHUNE - UNITER ENUATING FROM - SERVICE 1973

SHURE - FUSITION COOKDINATES - X + 1575-0 T + 525-0 KELOMETERS

SHIRE LATITULE = 350 UNSHORE AZIMUTH = 2900 DEGREES NEARSHINE SEUPE = 00:15 AVERAGE FETCH = 30700 RELUMETERS

TIMES - SPHING TIME HONDE + 2004 NEAP TIME HANGE + 1083 METERS ACCES AT MIGHTING FOR COURS

SEMPOTORNAL FILE - FORM NUMBER 15 1412

¥يون⊶	•	٠	۸.	•1	MARC. PRESS.	=1ND ANULE	SURF.	UNSH #150	AL SH #IND	EFFECT.	#AVE	WAVE	bk M	E AREH Anule	15C	TIDE
	. •	4,4	HAL	HAD	MB	Jt.u	٠,	w · 5	M 5	4/5	*	SEC	•	DEG	CM/St C	•
	814.		3.26	-0.48	1008.77	02.6	8.7	4.0	7.1	7.1	3.31	2.1	0.40	28.7	30. 15	6.30
ž	# 3 B .	4 4.	1.26	-06		04.4	7 4 4	• • 2	8.4	7.4		2.6	0.00	46.4	40.55	
,	25 .		3.26	-0.4	120 .90	62.1	Y . *	4.4	0 . 3	7.7	0.57	٥. د	44	25.4	56.80 64.65	10.5
:	A .	3 * * •	- 1 (5		1.07.44	01.7	17.3	4.E	H.C	F	3.00	3.2	0.5x	20.1	12.14	4.01
7	846. -15.	3 3.	U+25		1.76.96	01.4		5.0	4.4	H . 6		3.7	U.67	27.4	77.64	2.17
,	*14.	307.	3.25		1075.98	61	10.7	5 . 4	9.4	b • Y	3.45	3.0	U . 74	27.8	80.59	2.09
8	***	,,,,	2.24		10.5.18	60.9	11.1	5.4	9.1	4.2	1.03	4.0	0.80	47.7	81.05	. 78
9	y '8.	301.	0.44		1:04.76	υ'	11.4	5.0		9.4	1.11	406	0.87	47.6	14.52	
• •		3 '9	• 4		100-13	٠.٠	11.7	5.4	10.4	9.7	1.17	4.5	1.00	41.6	77-11	U . 84
. 1	1.21.	3 '5.	ويون ويوز		1003.50	59.7	11.9	6.1	10.5	16	10.7	4.6	1.06	27.4	7 78	0.22
	1305.	, , ,			1012.19	59.6	12.3	٥.,	.0.0	3	1.42	4.7	1.14	47.5	76.85	0.23
14		, , , ,	3.44			59.3	12.5	0.3	10.7	3.4	1.45	4.8	1.18	47.4	81.45	4.47
. •			~			59	12.5	6.4	1 6	40.00	1.55	4.9	1.25	47.1	87.43	U-67
٠. د	36.	373.	3.4.	-2.25		>0.	14.0	6.5	1003	10.6			1.26		95.08	1.73
	1100.	9 7 9 e	2000	-0.24	444	73.4	14.5	0.5	1 . 6	.0.6		5	1033		105.4.	1.72
	12.	3'3.		-6.21		51.		6.0	1	. 3 . 4	1.13	5	1.44		166.34	1.75
		٠٠٠ و	0.18	-0.19		57.0	4.1	0.0		4	1.76	5.3	1.44		103.40	
	7.	3034	• • •			***	1.00	6.3	9.7	9.9	1.17				98.91	4 . 56
	14.46.	36.	3.16	~(··	7.5.08	5: . *	1	* • 1	٠. ٥	9.5	1. ! ?	5 . 3	1.47	40.1	72.63	vet4
4.5		1924	15	-0.13			1000	5.9	8.7	7.0	1.76	5.0	1.47	45.8	78.12	- • • •
2.	. 3	3 4 7 .	-12	11	993.5.	54.9	* . 8	5.3		7.7	1.76	7.0	1.14	42.2	71.68	- • 13
	. 4 . 2	4021	-12	-CB	492.47	53.0	4.5	- 1	5.0			4.4	1.24	24.8	20.34	0.15
,		-28.	3.11	-0.07	994	51.4	8	4.7	5 . 3	7.0	1.04	4.6	45	24.4	10.48	w + 75
-		•11·		-0.00	991.84	50.4	• 5	4.	5 • 8	6.6	41	400	0.00	23.7	lu.E.	4.76
•	145.	* ; * *	4.4.4	5	407.56		0.9	***	٠.	5.4	U . A 6		U • 7.7 U • 7.7	23.3	10.05	7 8
Ċ.	1-10	•1'•	9	-04	991.10	40	5.4	3	* + 6 * + 0	• • •	- • 76 J • • *		0 0 0	22.5 20.8	6/405 6/405	4 - 4 3
	2084	427.	3.38	-0.02	990.01	37.0	5.1	•••	3.1	* • •	·;		U 0	10.0	49.8.	6010
	15.	411	7	-6.3	990,58	36.7	4.3	3.7	4 4 4	4.1	3.41	: . 6	3.41	:5.5	36.25	73
	15-6-	440 .	3.00	0.00	990,40	2-64	1.5		1.5	1.0	0.31	2 . 1		10.1	41.47	5
1.1	150.00		4		0029		3.1	1.6	. 5	1 • 4	4		45	3.6	6 · U !·	U. C.
• •	1540.		0.003	'4		348.4	2.8	4.4	-1.5	2 - 8		1.0		-0.0	-8.50 -19.57	U. J.
1.5	161	4044	e e e e e e e e	و د ډر په ټون	6 9	3	2.6			4.				-/3.0	-25.34	
.5	.0344			0.05	790.19	2,1.5		6	- 9 . 4						-47.44	. 4. 9
lo		5.1.	33		490.54	200.0	3.9		-3.1	2.9	J + . 6	1.6	0.04		-44.07	444
•	.666.	7400	,31		990.75	279.4	***	2.6	:	2.4	4.47					4 4 4 3
1.0		52-4	-2002	3.09	991.01	2/3.3	*	- 1	-4.1	2.9	U . 1 H	1.	0 a 125 a 124			4.04
1.0		524.	-0.00		991.34	200. ·	•••	-0,4	-4,	4 • ¥		1.			6.74	1.96
21				1.	341.44	26		-0.8	-4.5	3			U 1 U Y		9 . 3 .	110?
	. 7 5 .	355.		0.17	994.30	25 '	5		- 5	9					-44.84	
. 3	. 742 .	56	•	4.13	492.09	254.4	٠.,	-:.3	-5.4	9 · ·	ueż.		U 4		***>	- 4 7 1
44		>72.			493.55	254.5	5.7		- ' • •	5 • 1	2	1.7			-26.64	
•	.03>	56	- 1.11	J.15	V73.89	250.7	5.9	-2.1	-: • ;	3	J	1.7	30.1		-47.00	
•	.8.77	344.		1.16	994.25	/46.6	6.0		- >	302	0.23				-31.75	
		·	2		194.54		0.1	t	- 5	5+2			1		-34.51	#8
,		. 6 •	-0.13	27	494.47	243.2	5.1			3 . 4	3.24		U B		-37.57	3
•	. * * *		14	- 10	995.15		• • •	-3	- 1 7	3.2	2025	2.0				
	. 9	633.	-3., s		995.75	244.2	6.0	- 1 - 1	->.,	3 . 3 3 . 3	0.25		نايەت ئايەن	-25.6	-44+63	4 + 2 3
- 7				و د د	9 - 0 - 4 -	23.0	6.7	- 2. *	- > .				U	-26.1	-41.74	4,0
			-	3.14	+ +5 . 1 *	235.	0.0	-3.7	-5.7	2 . 3	10.0	2.4	0.74	-23.4	- 34.6	
1.				3.30	99	237.7		- 3 . 0	-5.1	1.9	2.46	4 + 1	- • 6 4	5.	- 5 7	
- 2		• • •	6	0.2	997.44	294.	6.9	1	-	3.3	0.27				-34,43	
. 3	1452.	6'4. 6:1.		2.2.	998	331.8 V.Et.			-5.6	9 . 9 3 . 1	0.67				-34.55	0.15
				0.2.	*****	433.1	٠			3.1			3		-34.84	
	* : .	25.	÷0		199,9,		*		1	1.1	• 4	4 4 4		4. 4	- 24 . / 6	4.51
	. 90				454.15	234	1.4			1.3					- 17. 15	8000
. 9	. 9 +			. 43	454.44	234.	7.1		-: • •	***			• • •		-,,,,,	
. *		· ••		····	.,	43	`••	- · · ·		3.3	• 4 0		• • •	-, -,	,	

make them in the mean R & Section 5.

الأعلمان والمراكبة المراجع والمستوا والمسترجون

عمليم الراف المراجع والمجور المراجعين المسترام الأري المراجع

وياصد فالأرادي يواريع فيعهوا كالمعقد المصروبي مروايي مرواها

MUSTANO ISLAND, TEXAS

HUN BEGINS AT HOR IT ON JANUARY 23% 1972

STORM - BANDMETHIC PRESSURE AT CENTER OF COW = 998.0 MILLIBARS
PRESSURE AT CARGEST ENCIRCLING ISOBAR = 1016.0 MILLIBARS
MAXIMUM PRESSURE INCLUDED IN STORM = 1018.0 MILLIBARS

LENGTH OF MAJOR HALF AXIS = 1800+0 KILUMETERS LENGTH LF MINOR HALF AXIS = 500+0 KILUMETERS ORIENTATION OF MAJOR AXIS = 33+0 DEGREES FHUM NORTH

LUNGSHURE CURRENT EQUATION FROM FUX AND DAVIS: 1972

SHORE - POSITION COURSINATES - x = 2415.0 Y . 0.0 KILOMETERS

SHORE CATITUDE + 28. ONSHORE AZIMUTH = 280. DEGREES NEARSHORE SLUPE + 0.018 AVERAGE FETCH + 1000. KILUMETERS

TIDES - SPRING TIDE RANGE + 0.85 REAP TIDE RANGE + 0.80 METERS SUPPE AT LOW TIDE + 0.016 SUPPE AT HIGH TIDE - 0.020

MIAED DIGRAAL TIDE - FORM NUMBER IS 2.50

норн	. a	,	۸.	Y 1	HARU. PREJS.	NINU ANGLE	SURF.	0N5H	AL SH CP1#	EFFECT.	#4.E	*AVE	## B#	EARUR ANGUE	LSC VELUC.	TIDE
	ζ.		RAD	RAD	₩8	SEG	w · 5	M. 5	M75	M/S	•	SEC		DEC	(M/5t(•
7	1777.	031.	-0.26	-U.28	1018.60	47.6		0.0	0.0	0.0	0.00	ن و ن	J. 00	2:02	2.25	0
	1820.	828.	-0.26	-0.27	1018.59	47.7	0.0	0.0	0.0	0.0	0.00	U • O	0.00	21.0	0.45	0.23
9	1004.	824.	-0.26	-0.29	1018.58	47.8	0.1	0.0	0.0	0.1	0.00	0.0	0.00	22.1	0.75	U.28
1.	1967.	821.	-4.26		1018.57	47.9	0.2	0.1	0.1	0.1	0.00	0.1	0.00	22.5	1 8	0.37
11	1951.	818.	-0.24	-0.22	1018.54	48.1	0.3	0.2	0.4	0.3	0.00	0.1	0.00	22.8	2.04	0.45
12	1994,	• 1 • •	-6.24	-c.20	1018.50	48.2	0.5	0.3	0.4	0.4	0.00	0.2	0.00	23.1	3.04	0.52
13	2038.	811.	-0.27		1018.45	49.4	J.8	0.5	0.6	0.7	0.01	0.4	0.00	23.4	4.47	U.51
14	2014.	810 •	-0.21		1018.38	48.5	1 • 1	0.7	J.8	1.0	0.01	0.5	0.01	23.3	5.92	0.57
15	2111.	839.	-07		1018.28	48.7	1.5	1.0	1.1	1.3	0.03	0.7	0.02	23.4	7.57	v.55
16	2.47.	808.	-0.27		1019.16	40.6	2.0	1.0	1.5	1.6	0.04	0.8	0.04	23.5	40-3	J • 48
17	2185.	809.	-0.27		1018.00	49.0	2.0	1.7	2.0	2.3	0.07	1.0	0.06	23.0	11.49	J + 4 J
10	2220.	807.	-0.28		1017.78	***5	3.4	2.2	2.5	3.0	0.11	1.3	0.10	23.7	13.74	1 و ه ت
19	2256.	635.	-0.28		1017.52	49.4	4.3	2.8	3.2	3.8	0.16	1.6	0.14	6 3 . 9	10.10	0.21
20	2315.	834.	-C.28		1016.93	49.8	0.1	3.9	4.0	5.3	0.20	2.0	0.23	24.1	19.97	0.13
22	2435.	803. 801.	-0.29		1016.13	50.3	8 . 2	5 . 2	6.3	7 • 2	0.41	2.5	0.36	24.4	24.53	0.07
23	2494.	801.	-0.29		1015.09	50.8 51.4	10	0.6	0.1	9+1	0.61	3.0	0.53	24.6	29.47	0.04
24	4554.	748	-6.30		1012.36		12.6	8	9.8	11.0	0.86	3.6	0.74	24.9	34.34	0.03
- 1	2613.	97	-0.30		1010.79	52.2	14.4		11.3	12.5	1.12	4 • 1	0.96	25.4	38.61	0.02
į	4853.	811.	-0.31		1010.13	53.8	15.2	9.2	12.3	10.3	1.36	4 . 5	1.16	. 5 . 5	42.22	J. 03
3	2893.	92	-0.31		1009.52	54.6	14.8	8.6	12.3	13.1	1.55	4.9	1.31	45.0	44.5	0.04
	2733.	630.	-0.32		1006.98	55.6	14.1	8.0	12.1	12.7	1.69	5 • 1 5 • 3	1.44	25.8 26.0	45.90	0.04
•	2774	852.	-0.33		1008.54	56.7	13.2	7.2	11.0	11.2	1.84	5.4	1.49	26.3	46.16	0.05
6	2814.	865.	-0.34		1008.20	58.3	12.0	6.3	10.2	10.1	1.85	5.4	1.49	26.8	45.11	0.05
- 5	2854.	879.	-3.34		1007.99	60.8	13.6	5.2	7.5	5.8	1.83	5.5	1.43	27.2	42442	0.07 0.10
В	2894.	880.	-0.35		1307.67	67.0	8.7	3.4	8.0	7.0	1.30	4.7	0.92	.8.3	31.50	0.16
9	2935.	881.	-0.35		1007.56	81.9	6.0	0.8	0.0	4.3	0.53	3.1	0.23	44.8	11.57	0.23
10	2975.	881.	-0.35		100 . 68	103.6	4.0	-0.9	3.5	2.3	0.16	1.7	0.09	28.3	8.05	W+32
11	3015.	882.	-0.36	0.16	1008.02	126.5	3.7	-2.2	3.0	1.7	0.09	1.3	0.08	22.9	11.20	
12	3355.	863.	-2.30	1 7	1008.58	144.8	4.2	-3.5	2.4	1.0	0.06	1.4	J . Ob	10.4	11.30	U-51
13	1098.	374.	-0.30	0.19	1009.31	157.2	4.9	-4.5	1.9	1 • 7	J.08	104	J. Q.	11.4	10.15	U-58
1.4	3127.	45 .	-3.35	3.20	1309.98	103.6	5 . 3	-5.1	1.5	1.8	0.08	1.1	0.08	8.4	8.21	J. 61
15	3:	9 ·	-C.37	3.21	1010.70	168.3	5.6	-5.5	1.1	1.9	0.08	1.2	0.09	6.0	6.20	0.01
16	3:40.	84.	-0.37	0.22	1011.46	172.0	5 . 8	-5.7	0.8	1.9	0.09	1.2	U . U Y	4.1	4.36	26
1.	3221.	897.	-0.37		1012.22	174.9	5 . 8	-5.8	0.5	1.9	0.09	1.2	0.39	2.6	2.78	0.47
18	3253.	900.	38		1012.99	177.2	5.7	-5.7	0.2	1.9	0.09	1 . 2	0.16	1.4	1.40	7600
19	3284.	903.	38		1013.73	174.1	5.5	-5.5	0.0	1.0	J. U9	1.3	0	0.4	01	4026
20	33	907.	-0.38		1914.54	180.9	5.1	-5.1	-0.0	1 • 7	0.08	1 - 1	0.09	-0.4	-0.41	10
21	3357.	911.	-0.39		1015.28	182.4	* . t	6	-0.2	1.5	0.07	1 • 1	0.07	-1.1	-0.93	J. UB
22	3343.	915.	-0.39		1015.94	183.7	4.1	l	-2.5	1.3	りょじり	1.0	0.06	-1.7	-1.2.	U-03
23	3423.	919.	-3.43		14151	184.7	3.5	-3.5	-0.2	1-1	0.04	0.9	0.04	-2.2	-1.32	0.00
2 •	14500	923.	-3.40		1017.00	185.6	2.9	-2.9	-0 · Z	0.9	0.64	0.7	0.03	-2.7	~1.30	00
1 2	35:2.	927.	-3.40		1017.40	186.4	2.3	-2.3	-0.2	. • 8	0.01	٠٠٥	0.05	-3.1	-1.21	0.31
	3505.	938.	-0.41		1017.40	196.2	2.3	-4.3	-0.2	0.8	0.41		0.01	- 3 - 0	-1.73	0.03
3	3508	95	-0.41		1017.39	185.9	2.3	-2.3	-0.2	○•8	0.01	0.5	0.04	9	-1-19	U = 1) N
•	3511. 3515.	961.	-0.42		1017.39	185.7	2.3	~2.03	-0.2	0.8	2.01	0.5	2.01	-2.8	-1-15	
,	3518.	984.	-3.42		1017.39	105.4	2 • 3	-2 - 3	-0.2		0.01	0.5	0.01	-2.1	-1-10	0.05
7	3521.	995.	-3.43		1017.38	184.9	2.3	-2.3	-0.2	4.7	0.01	0.5	0.04	-4.6	-1.05	U - 06
	~ >		0.73		4 V 1 · 4 2 0	.04./		-2.3	-0.4	V • 7	0.01	0.5	L • (+)		-1.00	U = 38

MAVE ENERGY IN THE BREAKER ZONE . 3.601E 39 JOULES

* *

TOTAL LONG-SHURE CURRENT ENERGY + 9.799E 99 JOULES

TOTAL POSITIVE LONG-SHORE CURRENT ENERGY . 0.399E 09 JOULES

TUTAL NEGATIVE LUNG-SHURE CORREST ENERGY + 0.1615 33 JOULES

MUSTANG ISLANDE TEXAS

RUN BEGINS AT HOUR 7 ON JANUARY 12, 1972

STORY - HAROMETRIC PRESSURE AT CENTER OF LUM = 940.0 MILLIBARS
FRESSURE AT LANGEST ENCIRCLING ISOBAR = 1018.0 MILLIBARS
MAXIMUM PRESSURE INCLUDED IN STORM = 1022.0 MILLIBARS

LENGTH OF MAJOR HALF AXIS = 2500.0 KILOMETERS
LENGTH OF MINOR HALF AXIS = 400.0 KILOMETERS
UNIENTATION OF MAJOR AXIS = 25.0 DEGREES FROM NORTH

LONGSHORE CURRENT EGUATION FROM FOX AND DAVIS: 1972

SHORE - POSITION COORDINATES - X = 2415.0 Y . 0.0 KILOMETERS

SHURE LATITUDE = 28. ONSHORE AZIMUTH = 280. DEGREES AEARSHORE SLOPE = 0.018 AVERAGE FETCH = 1000. KILOMETERS

TIDES - SPRING TIDE RANGE = 0.85 NEAP TIDE RANGE = 0.30 METERS SLOPE AT LOW TIDE = 0.316 SLOPE AT HIGH TIDE = 3.320

MIXED DIGRNAL TIDE + FORM NUMBER IS 2.50

HOUR	*	٧	*1	*1	BARU. PRESS.	WIND ANGLE	SURF.	ONSH WIND	AL SH WIND	EFFECT.	WAVE	WAVE T	вк н	EAKER Angle	LSC VELUC.	TIDE
	KM	KM	RAD	RAD	мв	DEG	4/5	M/S	M/5	M/5	M	SEC	м	DŁu	CM/SEC	×
7	2381.	1593.	-0.41		1022.04	52.5	0.1	0.0	0.0	0.0	0.00	0.0	0.00	23.5 24.4	0 6 5 5 0 6 6 1	0.04
8	2436.	1583.	-0.41		1022.02	52.6	0.2	0.1	0.4	0 • 2	0.00	0.1	0.00	24.0	2.70	0.05
10	2492 •	1574.	-0.41		1021.96	52.8 52.9	0.6	0.8	1.1	0 • \$ 1 • 2	0.02	0.6	0.02	25.0	5.70	0.14
11	2603	1554.	-0.41		1021.57	53.1	2.9	1.7	2.3	2.5	0.07	1.0	0.06	25.2	10.06	0.23
12	4658+	1545.	-0.41		1021.06	53.3	5.2	3.1	4.2	4.5	0.17	1.6	0.15	45.5	10.44	0.34
13	2714.	1535	-0.41		1020-19	53.6	8.7	5.1	7.0	7.5	0.38	2.4	0.34	25.7	44.95	U+46
14	2759.	1535.	-0.41	0.01	1019.25	53.9	11.7	6.9	9.5	10.1	0.65	3.1	0.55	25.9	33.27	U.56
15	2604.	1535.	-0.42	0.03	1018.04	54.3	14.9	8.7	12.1	12.8	1.00	3 - 6	0.84	<0.1	41.55	0.63
16	48 +9 +	1535.	-0.42		1016.55	54.7	17.7	10.2	14.5	15.2	1.39	4.5	1.16	26.2	48.90	0.00
17	2894 •	1535.	-0.42		1014-67	55.3	19.7	11.2	16.2	10.0	1.78	5.2	1 - 4 5	20.4	54.48	U. 63
16	2939.	1535.	-0.42		1613.11	56.1	20.1	11.2	16.7	17-1	2.11	5.6	1 . 74	20.6	57.25	U-56
19	2 184 .	1535.	-0.42		1311.46	57.4 59.4	18.7	10.0	15.6	15 · 6 13 · 4	2.30	6.0	1.67	27.4	55.67	0046 6600
20 21	3023.	1538.	-0.43		1010.35	64.4	16.1	5.3	13.8	10.0	2.34	6.2	1.74	28.4	46.09	0.21
22	3100.	15-5	-0.43		1009.19	82.4	7.0	0.9	7.0	5.0	0.71	3.0	0.30	24.4	12.61	0.11
23	3139.	15-8	-0.44		1309.25	120.9	3.9	-2.0	3.3	1.9	0.11	1.4	0.09	24.5	10.37	U . U 4
24	3177.	1552.	-0.44		1009.76	157.4	4.9	-4.5	1.8	1.7	0.09	1.3	0.09	10.6	8.20	0.01
1	3215.	1555.	-0.44		1010.67	175.9	6.6	-6.6	0.4	2.2	0.09	1.2	0.10	4.1	2.36	0.00
2	3264.	1564.	-0.44		1012-19	186.2	8.3	-8	-0.8	2 • 7	0.12	1.5	0.12	-3.2	-3.52	0.02
3	3311.	1573.	-0.45		1013.94	191.6	9.0	-8.9	-1.8	3.0	0.14	1.5	0.15	-6.1	-7.07	V. U4
4	3358.	1501.	-0.45		1015.73	194.9	5.9	-8.6	-2.2	3.0	0.16	1.6	0.17	-7.0	-4.37	5
5	3406	1590.	-0.46		1017-40	447.2	7.9	-7.5	-2.3	2 • 7	0.16	1.6	0.17			0.00
6	3453.	1599.	-0.46		1018.81	199.0	6.5	-6 • 1	-2.1	2 • 2	0.15	1.7	0.16	-9-1	-7.67	U.U5
8	3501.	1608.	-0.47		1019.93	200.4	4.9	-4.6 -3.1	-1.7	1.7	0.08	1.3	U+04	-9.7 -10.2	-5.55	L.U5
9	3548.	1611.	-0.47		1021-32	202.6	3.4 2.1	-2.0	-0.8	1+2	0.01	0.0	0.01	-10.0	-3.07	0.06
10	3643.	1617.	-0.48		1021.67	203.4	1.2	-1.1	-0.5	0.4	0.00	0.3	0.00	-11.0	-4.20	0.16
11	3691 .	1621.	-0.48		1021.86	234.0	0.7	-0.6	-0.2	0 • 2	0.00	0.1	U . OU	-11.3	-1.28	V.16
12	3738.	1624.	-0.48		1321.97	204.6	0.3	-0.3	-0.1	0.1	0.00	0.0	0.00	-11.5	-0.07	4.20
13	3786.	1627.	-6.49	0.28	1022.02	205.1	0.1	-0.1	-0.0	0.0	0.00	0.0	0.00	-11.6	-0.32	0.39
14	3827.	1633.	-0.49	0.29	1022.04	205.4	0.0	-0.0	-0.0	0.0	0.00	J.0	3.00	-11.9	-0.16	0.50
15	3868.	1640.	-0.49		1022.05	205.7	0.0	-0.0	-0.0	0.0	0.00	0.0	0.00	-12.0	-0.01	5.59
16	3909.	1646.	-0.50		1022.05	500.0	0.0	-0.0	-0.0	0.0	0.00	0.0	0.00	-12-2	-0000	0.04
17	3951.	1653.	-0.50		1022.05	206.2	0.0	-0.0	-0.0	0.0	0.00	9.0	0.00	-12-2	-0.0.	0.54
18	3992 •	1659.	-0.50		1022-05	206.5	0.0	-0.0	-0.0	0.0	0.00) • U	0.00 0.00	-14.4	-0.00	0.51
19 20	4033.	1666.	-0.51 -0.51		1022.05	206.7	0.0	-0.0	-0.0	0.0	3.03	5.0	0.00	-14.44	-0.50	96.0
21	4030	1698.	-0.52		1022.05	206.5	0.0	-0.0	~6.0	0.0	3.00	0.0	0.00	-12.4	-0.30	J. 47
22	4028	17140	-0.52		1022.05	206.4	0.0	-0.0	-0.0	0.0	0.00	0.0	J. GJ	-1	-3.50	0.10
23	4026+	1730.	-3.52		1022.05	206.3	0.0	-0.0	-0.0	U • U	0.00	0.0	0 + 0 €	-12.5	- 3•00	C+07
24	4C25.	1746.	-0.53		1022.05	236.2	0.0	-0.0	-0.0	0.0	0.00	U . U	بات ۽ ن	-12.2	-0.53	
1	4023.	1762.	-0.53		1022.05	206.1	0.0	-0.0	-0.0	0.0	0.00	0.0	U • QO	-1	-0.00	U . L .
2	4111.	1769.	-0.54		1022.05	206.6	0.0	-0.0	-0.0	U • U	3.00	U • U	0.00	-12.4		0.01
3	4199	1775	-0.54		1022.05	207.0	0.0	-0.0	~0.0	0.0	0.50	0.0	J • U L	-12.0	-0.00	
4	4286.	1781.	-0.55		1022.06	207.3	0.0	-0.0	-0.0	0.0	0.00	0.0	0.00	-12.H	-0.00	
5	43744	1788.	-0.56		1022.06	207.6	0.0	-0.0	~ (• 0	0.0	0.00	0.0	U+00		ವಾಕವರ ಚಿಕ∪ಬ	U.U8
6	4462.	1794.	-0.57		1022.06	208.1	0.0	-0.0	-0.0 -0.0	0 ∪	3.00	0.0	U + U U			7
8	4597.	1837.	-0.57		1022.36	208.2	0.0	-3.3	-0.0	0	J.30	0.0	0.00	, •. U•U		U.L
ě	4643.	1512.	-3.57		1022.06	208.3	0.0	-0.0	-0.5	(• 0	0.00	0.0	U • U()	U.U		7
10	·690 ·	1818.	-0.58		1022.06	208.4	0.0	-0.0	-0.0	0.0	0.00	0.0	2000	5.0		40.0
ii	4737.	1824.	-0.58		1022.06	208.5	0.0	-0.0	-0.0	ن و ن	0.00	J . J	3000	0.0	ل ل و ب	0.14
12	4783.	1829.	-0.59	0.53	1022.06	208.6	0.0	-0. 0	-0.0	0.0	0.00	0.00	0.00		U + U U	
13	4830 .	1835.	-0.59	0.54	1055.00	208.6	0.0	-0.0	- 0•0	0.6	0.00	0.0	0.00	0.0	3 ⊕)∪	0 . 32

MAVE ENERGY IN THE BREAKER ZONE + 3.721E 09 JULIES

TOTAL LONG-SHORE CURRENT ENERGY + 0.593E 09 JOULES

TOTAL POSITIVE LONG-SHORE CURRENT ENERGY . 3.5406 09 JOULES

TOTAL NEGATIVE CONG-SHORE COMMENT ENERGY # 0.846E 05 JOULES

MONTEREY. CALIFORNIA

RUN BEGINS AT HOUR 4 ON FEBUARY 13 1967

STORM - BARJMETRIC PRESSURE AT CENTER OF LOW - 998.0 MILLIBARS
PRESSURE AT LANGEST ENCINCLING ISOBAR - 1012.0 MILLIBARS
MAXIMUM PRESSURE INCLUDED IN STORM - 1014.0 MILLIBARS

LENGTH OF MAJOR HALF AXIS = 700.0 KILOMETERS
LENGTH OF MINOR HALF AXIS = 390.0 KILOMETERS
ORIENTATION OF MAJOR AXIS = 50.0 DEGREES FROM NORTH

LONGS-ORE CURRENT EQUATION FROM LONGUET-HIGGINS: 1973

SHORE - POSITION COORDINATES - X = 817.0 Y = 813.0 KILOMETERS

SHORE LATITUDE = 37. ONSHORE AZIMUTH = 120. DEGREES NEARSHORE SLOPE = 0.0057 AVERAGE FETCH = 50000. KILOMETERS

TIDES - SPRIAG TIDE RANGE - 2.07 NEAP TIDE RANGE - 0.91 METERS SLOPE AT HIGH TIDE - 0.058 SLOPE AT HIGH TIDE - 0.058

MIXED SEMICILARAL TIDE - FORM NUMBER IS 1.00

HOUR	*	٧	X L	٧1	BARU.	#IND	SURF.	ONSH	ALSH	EFFECT.	WAVE	MAVE	ø.	LALLK	LSC	
					PRESS.	ANGLE	WIND	WIND	WIND	WIND	н	7	n	ANGLE	VELUL.	TIDE
	K.44	KM	RAD	RAD	MB	DEG	M/S	M/S	M/5	M/S	м	SEC	м	DEG	CM/SLC	M
	895.	1235.	J + 38	0.14	1010.24	239.3	6.7	-3.4	-5.7	3.3	3.10	1.2	U.08	-27.4	-31.06	1.14
5	906.	1217	0.37		1009.45	243.0	7.4	-3.3	-6.6	3.8	0.16	1.5	0.12		-46.36	1.00
6	927.	1200.	3.37		1008.60	247.5	8.1	-3.1	-7.5	4.4	0.21	1.8	0.15		-53.42	1.03
1	948.	1184.	0.36		1007.7.	252.6	8.6	-4.0	-8.4	5.0	0.28	2.1	0.17		-56.55	1.08
8	970.	1164.	2.36	0.04	1006.85	259.2	9.5	-1.7	-9.4	5.8	0.36	2.3	0.17		-00.23	1.19
9	991.	1147.	0.35	3.01	1006.02	266.8	10.4	-0.5	-10.4	6.7	0.46	2.6	0.12		-50.87	1.35
10	1012.	1129.	0.35		1005.28	275.3	11.4	1.0	-11.3	7.9	0.58	3.0	0.20	-34.2	-65.53	1.52
1:	1038.	1138.	0.34		1004.57	286.1	12.4	3.4	-11.9	9.4	0.75	3.4	0.44	-31.0	-95.45	1.68
12	1064.	1087.	0.34		1004.11	298.9	12.7	6.1	-11.1	10.5	0.93	3.8	0.72	-28.0-	113.99	1.50
13	1090	1066.	0.33		1003.94	314.2	12.2	8.5	-8.7	11.0	1.09	4.1	0.99	-22.5-	114.56	1.66
14	1116.	1044.	0.33		1004.09	330.3	12.2	10.6	-6.0	11.7	1.26	4.4	1.25	-15.3		1.86
15	11-2-	1023.	0.32		1004-55	344.9	12.7	12.3	-3.3	12.6	1.44	4.7	1.49		-54.18	1.00
16	1108+	1002.	0.32		1005.26	356.8	13.4	13.3	-0.7	13.3	1.62	5.0	1.70	-1.7	-12.54	1.0
17 18	1189.	989.	0.32		1005.98	3 . 2	13.7	13.7	0.7	13.7	1.79	5 • 3	1.48	1.7	13.45	1.60
19	1210.	976. 963.	3.32		1006.76	8.5	13.8	13.6	2.0	13.7	1.94	5.5	2.03	4.5	36.07	1.50
20	1252.	950.			1007.58	12.7	13.5	13.2	3.0	13.4	2.05	5.7	Z + 14	6.7	56.34	1.45
21	1273.	937.	0.31		1008-41	16.2	13.0	12.5	3.6	12.9	2 - 13	5.8	2.21	6.5	70.53	1.43
22	1294	924	3.31		1009.22	19.1 21.6	12.3 11.3	11.6	4.0	12.0	2.18	5.9	2.25	10.0	64.80	1.40
23	1322.	906	0.31		1010.91	24.3	9.7	10.5	4.1	11.0	2.17	5.9	2.23	11.2	92.07	1.56
24	1349.	868.	0.31		1011.70	26.6	8.0	7.2	4.0	7.7	1.67	5.1	1.70	12.6	89.55	1.54
1	1377	870.	0.31		1012.35	28.5	6.4	5.6	3.6	6.1	1.20 0.80	4 • 4	1.21	13.7	61.00	1.50
ž	1405	853.	0.31		1012.86	30.0	4.8	4.2	2.4	4.6	0.49	3.6 4.6	J.80	14.5	70.43	1.54
3	1432 .	835.	0.31		1013.24	31.4	3.5	3.0	1.6	3.3	0.48	2.1	J-28	15.7	44.57	1.37
4	1460.	817.	0.30		1013.52	32.5	2.4	2.0	1.3	2.3	0.14	1.5	0.14	10.1	32.64	1
5	1505 .	799.	0.31		1013.78	33.5	1.3	1.1	0.7	1.2	0.04	3.9	0.04	16.2	18.39	1.13
5	1551 .	761.	0.32	-0.62	1013.92	34.3	0.6	0.5	0.3	0.6	0.01	0.4	0.01	10.3	9.21	1.65
7	1596.	763.	0.33	-0.66	1013.98	35.0	0.2	0.2	0.1	0.2	0.00	0.2	U. OU	16.0	4.15	1.03
6	1641.	746.	J.33	-0.71	1014.01	35.6	0.1	0.0	0.0	0.1	0.00	0.0	0.00	J. U	0.00	1.09
9	1687.	728.	0.34	-6.75	1014.02	30.1	0.0	0.0	0.0	J.0	0.00	0.0	0.00	0.0	0.00	1.22
10	1732 .	110.	0.35	-0.80	1014.02	36.6	0.0	٥.٥	0.0	0.0	0.00	J. U	0.00	5.0	0.00	1.41
11	1789.	694.	0.36		1014.02	36.9	0.0	0.0	0.0	0.0	3.00	0.0	0.00	0.0	0.00	1.04
12	1845.	5/8.	0.37	-0.91	1014.02	37.2	0.0	0.0	0.0	0.0	0.00	0.0	0.00	0.0	0.00	1.61
13	19.2+	661.	0.39		101+.02	37.4	0.0	0.0	٥.٠	0.0	0.00	0.0	0.00	0.0	4.00	1.95
14	. 439.	645.	0.40		1014.03	36	0.0	0.0	0.0	0.0	0.00	0.0	0.00	0.0	J. 00	2.02
15	2315.	679.	3.41		1014.03	37.8	U . O	0.0	0.0	J+0	0.00	U . J	U . ∪€	ن ۽ ن	J. QU	2.01
16	2072.	515.	0.43		1014-03	38.0	0.0	3.6	0.0	0.0	0.00	0.4	0.00	3.0	0.00	1.92
17 18	2116.	592.	0.43		10: 03	38.3	0.0	0.6	0.0	٥.٥	3.40	0.0	0.00	ن. ن	J. C.	1.70
19	2160.	571.	0.43		1014.03	38.6	3.0	0.0	0.0	0.0	0.00	0.0	0.00	0.0	0.00	1.01
20	2203.	550. 529.	0.44		1014-03	30.0	0.0	0.0	0.0	0.0	0.00	0.6	0.00	0.0	₩• @८	
21	2291.	508.			1014.03	39.1	0.0	0.0	ن • د	0.0	J-11	U.U	0.00	0.4	0.00	** 36
22	2335.	487.	0.45		1014.03	39.3 39.5	0.0	0.0	0.9	0.0	0.00	0.0	0.0 0	U . U	0.00	1.0.
23	2392.	4 1	0.46		1014.03	39.5	0.0	0.0	0.0	U.U	0.00	<i>U</i> • •	0.00	يدون	0.30	1 .
24	2449.	454.	U.48		1014.03	39.6	3.0	⊍•⊍ ⊍•⊍	0.0	J.J	3.00	0.0	0.00	2.0	7.30	4 . 17
1	Z505 .	438	0.49		1014.03	39.7	0.0	0.0	0.0	0.0	0.00	J	0.00	3.0	9.40	1.44
ż	2562.	422	0.50		1014.03	39.7	0.0	0 ± (1	0.0	0.0	0.00	C.U	0.00	3.0	6.00	
3	2619.	405.	0.52		101-03	39.8	0.0	0.0	0.0	0.0	0.00	0.0	0.00	J. U	0.00 0.00	1.47
	2676.	389.	0.53		1014.03	39.8	0.0	U • U	0.0	0.0	0.00	J.0	0.00	3.0	0.00	4.43
										0.0	~~~	~ • •	3		U . U U	4

MAVE ENERGY IN THE BREAKER ZONE = 9.155E 10 JOULES
TOTAL LONG-SHURE CURRENT ENERGY = 0.530E 09 JOULES

TOTAL POSITIVE LONG-SHORE CURRENT ENERGY . 0.389L DY JOULES

TOTAL NEGATIVE LONG-SHORE CURRENT ENERGY + 0.141E 09 JOULES

SOUTH BEACH. OREGON

RUN BEGINS AT HOUR & ON DECEMBER 14. 1973

STORM - BAROMETRIC PRESSURE AT CENTER OF LOW . 990.0 MILLIBARS
PRESSURE AT LANGEST ENCIRCLING ISOBAR . 1024.0 MILLIBARS
MAXIMUM PRESSURE INCLUDED IN STORM . 1028.9 MILLIBARS

LENGTH OF MAJOR HALF AXIS = 500+0 KILOMETERS
LENGTH OF MINOR HALF AXIS = 500+0 KILOMETERS
ORIENTATION OF MAJOR AXIS = 0+0 Degrees from North

LONGSHORE CURRENT EQUATION FROM FOX AND DAVIS. 1972

SHORE - POSITION COORDINATES - X = 4400.0 Y = 1520.0 KILOMETERS

SMORE LATITUDE = 43. ONSHORE AZIMUTH = 90. DEGREES NEARSHORE SLOPE = 0.020 AVERAGE FETCH = 5000. KILOMETERS

TIDES - SPRING TIDE RANGE = 3.81 NEAP TIDE RANGE = 1.85 METERS SLOPE AT LOW TIDE = 0.017 SLOPE AT HIGH TIDE = 0.023

MIXED SEMIDIURNAL TIDE - FORM NUMBER IS 0.90

HOUF	l g	Y	XI	Yl	BARC. PRESS.	w I ND ANGLE	SURF.	ONSH	ALSH WIND	EFFECT.	WAVE H	mAVE T	ы Н	EAKER ANGLE	LSC VELUC.	TIDE
	K M	K.M	RAD	RAD	мв	DEG	M/S	M/S	M/S	M/S	м	SEC	м	DEG	CM/SEC	M
4	3173.	1085.	-0.58	1.63	1028.92	211.3	0.0	-0.0	-0.0	0.0	0.00	0.0	0.00	-14.5	-0.00	27
5	32+3+	1062.	-0.61		1028.92	208.6	0.0	-0.0	-0.0	0.0	0.00	0.0	0.00	-1200	-0.00	
6	33:3.	1040.	-0.64		1028.92	235.8	0.0	-0.0	-0.0	0.0	0.00	0.0	0.00	-14.1	-0.00	U . 84
6	3353.	994	-0.67 -0.70		1028.92	202.9	0.0	-0.0	-0.0	0.0	0.00	0.0	0.00	-10.8	-0.00	0.71
9	3523.	972	-0.73		1028.92	197.0	0.0	-0.0	-0.0	0.¢	0.00	0.0	0.00	-9.5 -8.1	-0.01	0.73
10	3593.	949	~0.76		1028.91	193.7	0.0	-0.0	-0.0	0.0	0.00	0.0	0.60	-0.0	-0.05	1.29
11	3049.	963.	-0.71		1028.88	193.4	0.1	-3.1	-0.0	0.0	3.00	0.0	0.00	-6.5	-0.16	74
1.2	3736.	1017.	-0.07	0.92	1028.82	1 73 - 1	0.3	-0.2	-0.0	0.1	0.00	0.0	0.00	-0.5	-0.44	2.47
13	3752	1050.	-1.62		1028.67	192.6	0.6	-0.6	-0.1	0.2	0.00	C • 1	0.00	-6.5	-0.90	ن 5 <u>ه م</u>
14	3819.	109.	-0.58		1028.35	192+1	1 • 3	-1 - 3	-0.2	0 • 4	0.00	Uek	0.00	-0.3	-1.69	2.04
15	30 '5 • 3 4 3 2 •	1118.	~0.53 ~0.49		1027.74	191.6 190.8	2.5	-2.5	-0.5	0.8	0.01	J • 4	0.01	-6	-2.75	2.50
.;	3952	11900	~		1020.64	190.8	7.0	-4.3	-0.8	1.5	0.03	1.0	10.0	-5.1	-3.77	20
	4.31.	1229.	~ 38		1022.20	190.7	13.2	-10.0	-1.9	3.4	0.12	1.4	د نون	-5.7	-7.29	1.01
19	*0sl.	.267.	-0.33		1018.57	190.6	13.8	-13.5	-2.5	4.6	0.41	1.0	0.21	-5.0	-8.98	U.50
20	-1:1:	1330.			1613.82	190.5	17.0	-10.7	-3.1	5.7	0.31	2.2	0.34		-10.50	
21	4100	1344.	-2.25		1308.24	190.4	19.0	-18.7	- 5 . 4	6.4	0.40	2.5	0.42		-11.07	U 7
42	423 y •	1333.	-0.18	0.22	.002.34	190.1	19.0	-18.7	3 . د –	6 • 4	0.48	2 • 7	U.50		2.37	4.0
23	4269.	1414.	-0.12 -0.07	0.17	997.45 993.49	192.6	16.8	-10.5	-3.7	5 • 7	0.51	2.0	0.53		-14.00	0.50
1	4346	1505	-0.02	0.07	990.98	217.0	12.5	-5.6	-3.9	2.8	0.37	2.4	U • 13		-18.55	1.12
ž	4394.	1546.	J. J3	C+32	994.28	312.4	6.2	4 . 2	-4.0	5.5	0.26	2.0	0.23		-23.33	1012
3	4423.	.586.	3.08	-0.03	991.49	3.0	14.0	14.0	0.7	14.0	0.89	3.6	0.92	1.6	6.14	1.51
4	4401.	1027.	3.14		994.44	14.0	22.7	22.0	5.5	22.5	1.94	5.2	1.98	7 . 7	34.35	1.50
5	4527	1674.	0.20		1000.62	23.6	30.2	27.7	1 2 • 1	29.4	3.21	6.7	3.21	12.5	14.39	1.37
7	~592 ·	1722.	0.26		1007.99	27.9	31.2	27.6	14.6	30.0	4.13	7 • 7	4.07	1400	84.00	1.15
ė	4723	1816.	0.39		1.43.56	3043 3148	26.7 19.6	23·1 16·7	13.5 10.3	25.5 18.6	4.20	8.3 8.1	4.49	15-8	93.44 89.73	leve
ě	4788.	1864.	0.45		1024.40	32.8	12.5	10.5	6.8	11.9	2.29	6.0	2.26	16.3	64.67	U. 76
10	4854.	1911.	0.52		1326.72	13.6	7.0	5.8	3.9	0.6	0.89	3 • 7	0.87	17.0	42.25	1.15
11	4913.	1988 .	0.62		1026-10	31.9	3.0	2.5	1.6	2.9	0.21	1.8	0.21	16.0	24.38	1.44
12		2065.	0.72		1028.66	30.7	1.1	0.9	0.5	1.0	0.03	0.7	0.03	14.9	7.68	1.77
13	5031.	2142.	C - 82		1028.85	29.7	0.3	J • 3	0.1	11 • 3	0.00	0.2	0.00	1403	6.30	4 + 4 4
14 15	5389. 5148.	2219.	0.93		1028.91 1028.92	28.9	0.0	9.0	U . U	J.O	0.00	0.0	0.00	3.0	0. 00	
16	5207	2373.	1.03		1028.92	20.3	J.O	J.U U.U	0.U	0.0 0.0	0.00	0.0	0.00	0.0 J•0	0.00	4 1 3 0
17	5250.	23.30	1.16		1023.92	28.5	0.0	0.0	3.0	J • 5		0.0	J.00	U • U	0.00	1043
. 8	5293.	2414.	1.19		10.8.92	29.3	0.0	0.0	J.U	⊍•3	0.00	0.0	0.00	3 6 4	L	1.52
19	5335.	2434.	1.21		1320.92	30.0	0.0	0.0	2.0	0.0	0.00	0.0	U a UU	J. U	6.00	
20	5378.	24544	1.24		1028.92	30.6	0.0	2.2	0.0	0.0	J • UU	0.0		3.0	0.00	V . 68
21	5421.	2475.	1.27		1028.92	31.2	0.0	3	0.0	4.0	0.00	0.0		ن ۽ ن		
22	5520.	25434	1.30		1028.92	31.8	0.0	0.0 0.0	0.0	0.0 0.0	0.00 00	ن و در	0.00	ن و ر دون	پ≎ ډل	4 0 2 4
24	5577.	2570	1.42		1028.93	32.3	3.0	3.	3.0	U • U	0.00		U • U C	0.00	3.3 0 2.00	0.39
ī	5633.	2637.	1.49		1028.93	32.1	0.0	2.0	0.0	V.0	J. UU	. • •		2.0	0.00	
2	2047.	2685.	55		1020.93	32.2	0.0	3.0	0.0	0.0	0.00				3.00	
3	5740.	2734.	1.01		1958*33	32.3	→ 0	u • €	0.0	0.0	300		0.00			
5	5853. 5866.	278	1.03		1028.93	32.4	0.0	4 • 4		0.0	9 € 6 5		0.00	0.0	- · · ·	
0	>955.	2821.	1.73 1. E		1028.93	32, '	0.0). 3.		U • U	0.00	0.0		. • -	5° • • •	. • * •
Ÿ	5443.	2902.	1.54		1.28.93	33.4	3.0	3.		U • U	3.00	() (()	0.00	200		
8	6056.	2442.	1.87		1:28.73	33.		J • .	6.4	(1.0	J. UJ	Leu		- • •		~
9	6120.	2982 .	1.75	-2.29	1028.95	33.9		0.0	الأوان	0.0	3.00	17 . U				
10	c183.	3023.	2		1:29493	34.2	:.3	2 •		Casi	- • - ١		0.00			
11	62.42.	3066.	6		1-29.93	34.3	- • •	J.	51 € 57	0	4.40	0.0	0.00			
13	6333. 635 9.	3152.	2.11		1025.93 1028.93	34.4	3.0	J•0 3•5	U + J	i. • U	0.00	J • J	0.00	. • •	. • • •	. • • `
14	c 418 .	3195	2.23		1028493	34.6	2.0	. • .	∪ • ∪	U • U		0.00	U . U . U . U .			
15	64.6.	3233.	,		1029.73	3 ?	3.0	9.1	0.0	0.0	J. 00		0.00			
15	45 15.	3281.	2.34		1329.93	* • • b	J	0.00	0.0	5.5	3.00	3.0	0.00		• • •	

MAYE ENERGY IN THI BREAKTH ZUNE x DeHUSE TO DOULES

FUTTL ESAU-SHURE LUNGENT ETERSY + LAWINE TO DOUGES

TOTAL + STITLE CONSTRUCT COMMENT ENERGY + LAWORE IS SOCIES

TOTAL NEWATIVE CONSESSION COMMENT ONE HOW A CONTROL OF SCHOOL

CINCULAR NO AN TEST

RUN REUTY AT MOUR IT ON LOUR HE TYPE

THE PROPERTY OF THE PROPERTY O

CENOTH OF MAJOR HALF HAIS # 5. .. NICOMETERS CENOTH OF MINOR HALF HAIS # 30 .. NICOMETERS ORICENTATION OF HAUGH HITCH - CAL DEGREES FROM NICOMETERS

STORM VECTOTTS = ACT REFERENCES STORM AZIMUTH = 90.

SHOPE - POSITION COUNDINATES - C = 1000. Y = .. KILOMETEN.

SHOME CATITUD: = 43. INSHIRE AZIMOHH = 90. DEGREED NEARSHORE DEGREE K 0.233 AVERAGE METOH = 10 . KILOMETO .

п. "н		٠	Αż	61	6416. PRESS.	#1 ND ANGLE	JURE:	سو. ن ن ا ا	ALSH AINU	SPECT.	^/.t	##+E		EARER Avuet	vt.vi.
	٨.		₩A;	HAC	MB	DEG	M/5	4/5	My.	475		sti	•	دعو	. 4/26.
	22	500. 600.	1.1	1.15	1022.8	-46.2	9.3	~ • •	-0.	3 • 3				· · · ·	
3	560	530.	1.11	1.06	1022.8	105 2.3.3	3.J	J.0 J.0	-0.0 -0.0	J.U	er∎ili. U∎ili.	. ∎ ∪ • ∪	•0.	-2-05	
4	.00.	5.00	1.1	0.83	4 5	345.6		i .		200		U . U	0.00		
>	5-0·	500.	1 * 4 4	0.80	1022.6	31.4 €	0.0	0.0	-0.0	~ • ~	بالآه ل			-66.5	74.44
6	650. 720.	500.	1.11	C + 71	1 22.8	311.7	3.9	3.0	-0.0	J • U	ب د د	- • ∪	بال و ب		•
ė	To C.	5.00	1.11	0.62 0.53	1322.8	315.1 318.7	0 • 1 0 • 2	0.0 0.1	-0.0 -0.1	∪•i •i	0.00		0 0 0 00 0 0 00	-21.0	/
9	800.	5. (•	1.11	3.44	1202.0	322.5	3.2	2	-0.1	• .	- • J .			-1:	-:-
10	540.	5000	1.1.	0.35	1022.8	326.6	0.3	وون	- 146	V • 3	3.00		v • 1.		,
11	660.	5 √ ∴ •	1.11	0.46	1022.8	333.5	3.4 3.5		-004		♦ 12 m	- • 4	- • U .		-:::
1.2	924	500.	1.1.	C+17	1 22.8	2.55.2	2.5	0.5	-0.2	• 5	1.00		V •	= t ₂	
13	950. 1600.	5u. • 5u. •	1.11	3.08 1.00	102.00	آ. دو: و ۱۹۰۰	2.0	J., 3 J. 6	-0,2 -0,1) • 6	0.00 0.01	- • "		-11.5	***** ****
15	1040.	500.	1.11	-C.De	132	5	€	2.6	1	Ü.0	0.01	,		-5.5	-
10	10-0.	5.00	1.1.	-6.17	1022.8	252.4	5	3.5	-3.0	U • 5		J			
1.7	1:2.	5.0.	1.11	-0.20	1932.8	35/.8	→ • **	9 9 4	~~ J	U 4 **	< * UC	U + 2		-1	7 - 4 -
18	1150.	500	1.11	-C • 35	1.722.8	4	٥.٥	فوت	O		• • •	J • 3	-• -		
19 20	1200.	500.	1.41	-0.44 -0.52	1322.8	9. y	0 • ± 3 • 3		0.0	• 2		2.4	200	4.5	
21	1283.	600	1.11	-0.62	1022.8		0.1	2.1	0.0	.1	• • • •	•	2120		
	1300.	500.	1.1.	-0.71	1027.8	10.5	5.3	J . J	5.0	. • 5	0.00		2.00	5.	
23	13.0.	500.	1.11	- u • " y	.022.6	20.1	J • U	0.0	4.0	€*3				• •	- • -
24	1400.	5	1	-0.35	1024.5	23.0	J	3.0	V	₩	- +		- • . •	• • • •	• •
2	14-0.	6 0.	1.1.		102.48	25.7	0.0	2.5	0.0	€ * ()	3.0			٥.	٠
3	152	5 601.	1	-1	1 2	30	0.	•	0.0	₩.j	3.1.		3 • 7 € 6 • 7 •	•	J.,
	_							-							
ì	4434	4/04	0.00	1.15	8.3171 8.3171 8.1571	291.9	J.~	3.6	-0.0	,	- •	• •	40.00	-25.00	- • •
4	-200	ية بال 14	2.68€	1.16	1716.8	294.1	್ಕಾರಿ		-L.J	5.4.2	s a dis		~ · ·	0.~	
3	3 h . •	40.0	2.68	0.97	102114	276.6	2	4 • €	-0.0	5.44	4.4	- • •	Se # 4 4		
5	600 + 847+	400€ 400€	0.es	3.88 3.83	1142.4	294.3	3.1	2 6 . 2 6 a	-0.1 -0.2	. +1 . +2	- • • -	• •	ب. • . پ مان • د	-, -, -,	
	6.10.	400	. 86	5.71	1322.4	305.6	3.5		-11.4	J 4 4	2 • C 2 C •	::		-	
7	1.1	40.	2.50	6.4	1.127	344.3	2.4	2.5	***	U . 7	1	. 4	- 0.74		
	1000	ه ټاټ په	Jode	4.53	1027.7	111.3		2 . 8		1 • 1	• • •	0.40		-22.5	
	8∵≎•	400.	V • 5 F	2.4	1-22.6	317.7	•				4.0	. • 1	و ۽ د د		
13 11	843.	400 e	3.3H	3.26	1524.5	3.2.5	2.9	1.00	-1.4	2.42	ა•მნ მ•:	. • 4		0.0	7.4.1 7.1.4.1
12	920.	430.	3.35	0.17	1022.1	3. 1.6 33.0	3.9	3.0	-1.5	3 4 3	1.			-11.9	-11
13	960.	400.	0.98	Jaco	1227.12	137.€	4.1	3.5		3.6		• 5			- 17 . 3
14	1-32.	400.	೦. ಕಟ	0.00	1022.2	344.3	ظ و و	2.1	-i.v	2.8	4961	. • "	· • • •	- × .	
15	1040.	400.	0.89	-0.08	1022.2	35000	3 • 7		0	***	0	1 • 7	~ + ¿ ··	-: •	
16 17	1080.	430. 433.	∪•88 J•¤8	-0.17 -0.26	10274.	355.6 1.0	3.4	2.9	****	3.4	3.4		• • •	-2.	
18	1160.	4000	9.48	-0.15	1022.5	5	ذ ،	ز و و	V		•				
19	1200.	4J0.	0.65	-0.44	1022.7	14.9	. • 3 1 • 7	1.1							71 - 44
20	1240.	40J.	J.88	-0.53	1722.7		4 4 4	3.7	→ * *		S	• 7		. 4	
21	1280.	400.	J. 88	-0.62 -0.1	1022.7	19.3	J • 8	2 • 3		2 + 6	~• ·	• *	∪ 4	7.0	* • .
22	1320. 1360.	430. 430.	0.68 0.88	-0.79	1022.6	43.		3 4 4 4 9 2	0.1		V. UU	• 6	J • Ut	1007	3.0
24	1400	430	0.88	-C.88	1322.0	εÿ.3	J					••		.5.	3.7
1	1440.	47J.	J.88	-5.97	1072.8	3	. • -	~ · .	~ · .						
2	1480.	• . 0 •	0.88	-1.65	1:2	3	5.0	· ·	· • .		• •	. • •	U •	•	- • ·
7	1527.	400.	88	-1+15	1022.4	6.			- • -	. • •	• •		• • •	~•	- • -
1	44.		. • 66	1.1		. 54.4	. •								
2	425.	3000		2 - 2 -		300.3		.:		• •			• • • •	-	
	?; ":	3 .	2 1 94			. 54.	2.03			• .		• •		4.4.5	-,
	5 . ·	3-20	20.54				•	• -				• -		v	
5	550.	3	+56 -+56	• 9 t	1 4		. •	• •	7.4	• •	• .	•	• •	• •	-
	223	3				1-1-1	3.			•	• • •		•		-,:
5	`6 •) ·		•53		1.5	~ • •		-						
÷	400 ·	34. •	4.55	5.44		3		4			• • *	: :		-	- **
•	840.	3	• • • •	• 35	1 41.1	316.2	5 + E	• •		" • .		. • *		1.	
	88 ·	3 .	. • * *	• 2	112.45	3.20.						•		÷ · · · ·	' • •
15	**		2 + 5 t	•	12.12	336.							•		
		3.2	1.55	0.00			• • •					•			-
.•		;				41								-	
	• • •			- • •						• . • •	. • *	•			• . •
	12.	1	• • •	-0.2	. 6	٠							•		. * .
: .				-1.44				•		•	•		•		
				• · ·			4.7	٠.			•				
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CIRCULAR STORM TEST

RUN BEGINS AT HOUR I ON JULY 4. 1776

STORM - BAROMETRIC PRESSURE AT CENTER OF LOW = 1000.0 MILLIBARS

PRESSURE AT LARGEST ENCIRCLING ISOBAR = 1020.0 MILLIBARS

MAXIMUM PRESSURE INCLUDED IN STORM = 1022.9 MILLIBARS

LENGTH OF MAJOR HALF AXIS = 300.0 KILOMETERS
LENGTH OF MINOR HALF AXIS = 3JO.0 KILOMETERS
ORIENTATION OF MAJOR AXIS = 0.0 DEGREES FROM NORTH

STOPM VELOCITY = 40. KILOMETERS/HOUR STORM AZIMUTH = 90. SHORE - POSITION COORDINATES - X = 10000 Y = 0.0 KILOMETERS

SHORE LATITUDE = 43. UNSHORE AZIMUTH = 90. DEGREES NEARSHORE SLUPE = 1,033 AVENAGE FETCH = 1000. KILUMETERS

HGUR	×	Y	× 1	¥1	BARO. PRESS.	WIND ANGLE	SURF.	UNSH WIND	ALSH w1ND	EFFECT.	wAvt M	#AVE T	H BH	LAKER ANGLE	LSC VELUC.
	KM	KM	RAD	RAD	мв	DEG	4/5	4/5	M/5	M/5	м	SEC	м	DES	CM/St C
1	480.	200.	0.44	1.15	1022.8	273.9	0.1	0.0	-0.1	11 .]	0.00	0.0	JeUu	-30.5	-0.7
2	520 •	200.	0.44	1.06	1022.8	275.9	0.3	0.0	-0.3	9.2	0.00	0.1	3.30	-31.2	7.09
3	560 •	200.	0.44	0.97	1022.5	278.2 260.7	0.7	0.1 (.2	-0.7 -1.5	0.5 1.1	0.00	0.5	J.00	-31.·	-4.2 -8.2
5	640.	200.	0.44	0.80	1522.3	263.3	2.6	0.6	-2.8	2.1	0.05	0.9	J • Q 5		-14.4
6	680.	200.	U-44	0.71	1021.9	260.3	5.0	1 . 4	-4.8	3.8	0.13	1.4	J. Ub	- 9 9	-23.4
7 3	720 • 760 •	200.	0.44	0.62 0.53	1021.2	239.8	8.0 11.9	2.7	-7.5 -10.9	6 . 3 9 . 6	0.57	2.9	0.15	-30.4 -29.6	-35.8 -51.3
3	800.	200.	0.44	0.44	1019.0	299.3	16.3	7.9	-14.2	13.5	3.99	3.6	V. 17	-28.3	-60.3
10	840.	500.	0.44	0.35	1017.5	305.6	20.0	11.8	-15.5	17.5	1.55	4.0	1.30	6.3	-64.3
11	880 ·	200.	0.44	0.26	1016.J	313.3	23.1	16.3 20.8	-17.2 -15.9	21.2 24.4	2.22	5 • 7	2079	-23.3	-75.0 -77.4
13	960.	200.	0.44	0.08	1013.8	333.0	27.0	24.6	-12.5	26.0	3.64	7.3	3003	-14.9	-54.9
14	1000.	200.	0.44	0.00	1013.4	344.3	28.0	2/.0	-1.5	27.6	4.25	8.0	4+37	-6.4	-56.5
15 16	1040 •	200.	0.44	-0.08 -0.17	1013.8	355.5	27.6	20.0	-2.0 4.0B	27.5	5.13	8.5	5.30	-203	-16.9
17	1120.	200.	0.44	-0.26	1016.0	15.3	23.7	22.9	6.2	23.4	5.30	9.1	5.49	5.1	61.0
19	1160.	200.	0.44	-0.35	1017.5	23.0	20.3	18.7	7.9	19.7	4-51	8 • 4 6 • 9	4.59 3.02	14.0	31.03
19 20	1200 ·	200.	0.44	-0.44	1019.0	29.3 34.5	16.3	10.0	7.9	15.6	3.03	5.3	1.75	17.5	70.0
21	1230.	200.	0.44	-0.62	1021.2	38.8	8 • 4	6.5	5 • 2	7.6	0.71	8 • و	v•57	19.4	54.0
22	1320.	200.	0.44	-0.71 -0.79	1021.9	42.3	5 · " 3 · 2	2.2	3.6 6.6	4.9 2.9	0 0 3 B	2.4	36 	20.9	31.5 22.5
23	1360.	200.	0.44	-0.88	1022.5	45.2 47.5	1.7	1.2	1.3	1.5	0.00	0.3	0.00	6404	5.7
1	1440.	200.	0.44	−ĉ.97	1022.7	49.3	0.9	0.5	0.5	0.9	0.00	0.0	U . UU	.	ټ و د
2	1480.	200.	0.44	-1.06	1022.8	50.7	0.4 0.1	U • 2	. • 3 (- 1	0.1	0.00	0 • 6 J • 6	0 • 0 0 0 • 0 0	0.0 0.0	0.e∪ 0.e∪
3	1520.	200.	0.44	-1.15	1022.5	51.8	0.1	0.1	0.1		J • C J	3.0	0.00	0.5	
1	480.	1000	0 • 4 2	1.15	1022.8	259.1	3.2	-0.0	-0.2	10 a.k	0.00	J e U	0.00	-30.2	
2	520. 550.	100.	0.22	1.36	1022.7	260.5	3.5 1.1	-0.1	-0.5 -1.1	U • 3 U • 7	0.00	0.02	0.00	-31.1 -31.6	-2.0 -5.7
4	600.	100.	0.22	83.C	1022.3	254.0	2.2	-0.2	-206	1.4	J.C.		0.01	- 5:. 9	~5.6
5	643. 680.	100.	3.22	0.80	1021.8	266.2	··•1	-0.2	-4.1 -6.4	4.0	10.07	1.0	J. U.	-36.4	-12.6 -14.3
7	720 •	130.	0.22	0.62	1319.6	272.0	11.0	0.4	-1,	7.4			J. U.S	-36.6	-24.4
8	760.	130.	5.22	0.53	1017.8	275.4	16.0	1.5	-15.0	41.02	0.75	3 . 3	Ject	-36.0	-43.0
10	800. 840.	100.	0.22	0.44	1015.3	200.7	21•? 26•9	4.0	=20.8 =20.8	10 e 8 21 e 5	1.20	و . به ۲ و ي	1041	-34,4	-60.0 -90.0
11	880.	100.	C • 2 2	0.26	1009.5	294.1	30.5	14.9	~2 8	c = . =	2.10	*** *	1.96	-29.9	-113.0
12	920.	100.	0.22	0.17	1006.9	305.€	30.3	1/-	-44.6	45.44	5.41	• •	7		-126
13	950.	100.	0.22	0.08 0.00	1035.2	322.5	28.3 27.5	20.5	-17.2	21.00	4	•	3.657	-170,	-114.U
15	1040.	100.	0.22	-0.08	1005.2	6.1	29.3	2300	3	. 5 . 2	5 • Se	+ • î	う・25	3 . 5	24.1
15 17	1040.	100.	3.22	-0.17 -0.26	1006.9 1009.5	43.0 34.5	30 • ± 30 • 9	21.9	11.6	44.1	5.98 5.98	7.5	7.654	11.00	91.4 130.6
18	1150.	195.	3.22	-0.35	1012.5	42.3	29.1	210	12.0	26.5	5.71	7.4	2001	41.6	144.7
19	1200.	:00.	3.22	-0.44	1015.3	47.5	25	16.9	10.5	440-	4	5.4	3 4 9 4	43.5	434.6
20 21	1240.	100.	3.22	-0.62	1017.8 1019.6	50.7 52.7	19.7	8.0	15.2 11.5	17.3	2.45	5.0	58	24.4	11.00
22	1320.	100.	3.22	-0.71	1021.0	54.1	9.4	5.5	7.0		0.83		L . 70	23.5	4140
23	1360.	100.	0.22	-0.79	1021.6	55.1	5.7	3.2	4.7	4 . 7		ومع	0020	25.0	10.7
24	1440.	100.	0.22	-0.88 -0.97	1022.6	50.5	3.2	1.0	2 a 5	e • 1	J = 1/9 . = J()	104	·• UE	26+1	20.e
2	1480 .	100.	0.22	-1.00	1022.7	57.3	0.8	J.4		1 e 5	ن ۽ د ن	U . L	J • UC	V • J	J : U
3	1520.	100.	L.22	-1.15	1022.8	57.3	3.3	• 1	J • Z	J + 3	J • U ·	,	U • UU	Ÿ••	0.0
1	480.	٥.	0.05	1.15	1022.6	241.4	. • 4	-1.		J • 1	0.0.	· • ·	0.00	-20.6	-1.4
2	520. 560.	0.	0.00	1.06	1022.7	241.4 241.4	1.2	-0.2	-0.5 -1.0	(• 3 (• 6	0.00 1.00	0.1	J = 00	-21.6	-5.7
4	.003	0.	3.00	0.88	1322.2	241.4	2 . 3	-1.1	-2.0	1 • 2	0.52	Uet	J . U .	-21.1	-1/01
5	640.	.	0.00	0.80	1921.6	241.4	4.2	-2.	1 و د 🖚	4.1	0.05	0.9	9.614	-21.0	0.0
6	€83. 723.	0.	0.00	0.71	1020.5	241.4	10.2	-3.2	-6.9	3.4 5.1	0.14	1 • 5 1 • 8	J.17	-27.9	-23.6
8	765.	э.	3.00	0.53	1016.5	241.4	13.9	-6.6	-1646	7.6	0.30	2.4	2069	-20.1	-46.66
9	8JO.	0.	0.00 0.00	0.44 0.35	1013.4	241.4	17.1 18.6	-8	-15.0 -16.5	F • 6	0.57	249	0.57	- 46 - 4	:
10	840.	٠. د.	0.00	1.26	1005.2	241.4	18.1	-5.	-15.9	9.5	U . Bc	3 . (,	U . 1. t.	-28.1	-0.1.0
12	920.	ه ز	0.00	J.17	1003.0	241.4	14.4	-6.0	-1247	1.3	0 . 17	3 . 5	0.54	-21.8	-54.5
13	960.	٠.	0.00	0.06	1000.8	241.4	8.0	-3.8	ر . : -	پ و به	0.27	1 • 6	0.17	-27.0	-32.0
14 15	1000.	0.	-0.00 -0.00	-0.08	1000.3 1000.7	61.4	13.1	O e d h e d	11.5	10.8	0.05 0.56	2 4 9	0.44	2003 2001	0.0 52.44
16	1080.	0.	-0.06	-4.17	1.53.6	61.4	43.5	11.3	20.7	19.5	1.51	4.6	1.15	28.7	54 e 5
17 18	1120.	0 •	-0.01 -0.01	-0.24 -0.35	1006.2	61.4	29.6 30.8	14.2	26.3 47.1	24.65	3.67	5.9 6.9	1 e 6 9	28.4	100.7
19	1200.	٥.	-6.60	-5.44	1013.4	61.4	29.0	13.4	14.6	4304	3063	1.4	2014	2803	167.
20	1240 .	Ú.	-6.0C	-1.53	1115.5	61.4	22.7	13.5	19.9	16.	2.65	f + 5	2.4		110 - 7
21	1286 •	i. •	-,.00 -0.00	-3.62 -0.71	1020.5	61.4	16.7 11.4	7.9	14.6 9.8	14.2	1.64 J. P.5	1.6	10.0 U+68	48.0	#6 . ¥
23	1360.	1. •	-0.00	-3.79	1021.6	6 4	5.5	3	6.0	5.46	~ • 3 h	4	40.00	11.7	4 u . N
24	1400.		-0.00 -0.00	-0.88 -0.97	172.02	6 6 4	3.8	1.6	3,4	3 + £	9434	1 4 3	94.7		5 5 4 5 5 4 3
į	145	.7.	+ U + U	-1.06	1 42.5	61.4			• 8	• •		• •	1864		
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Light Cart of the State of

KIN ST INC AT HIGH I LIN DELY 44 1776

STORY OF THE PRODUCT OF THE OF CIA * 1000+0 MILLIBANS OF THE PRODUCT AT LEMEST ENGINEETH OF CIA * 1000+0 MILLIBANS OF THE PRODUCT PACKAGE OF THE PRODUCT OF

THE MINER COURTS IN A NO. FILL METERSCHOOK STERM AZIMOTH # 90.

HOLD SHOULD BE CONCINCTED - X & LUDGOU K & TO USU KICCHETERS

TO THE STATE OF TH

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		- 100		- 1	و در	51		3.6	1.4	1	0.00	0.0	3.00		• ` • •
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*	• •		-, +		1 44.	20404	ÿ.5	- 3.5	-0.46	0.2	4.00	. 1	3.00	-:	-1.2
		 <u></u>	-3:44 -:444 -:444	e. 0.e.	1:22.4	202.5 144.8	1.7	-1.8	-0.A	0 + 5 0 + 6	0.00	2	بائن وال بال وا	1.6	
				• 1	19	146.4	3.2	-3.1	-0.9	1+1	0.02			-8.8	
		-2.00		. t	10.1.2	193.3	5.1	-4.9	-1.:	1.7	0.04	0.0	- 114	- : - :	
		-	44	+5.2	1/20.5	189.1	7.3	-,.2	-1 • i	2 - 4	0.67	. • .			
•	84.	-6		35	1019.0	183.4	. 2 . 3	-9.8	e 0.5	3 • 3 4 • 1	U.1.		0417		•••
· .	89.	-2			1010.1	59.8	. 4 . 4	-1442	2.5	4.9	1.25	4.5	0		0.,
	4200	- LUJ.	-0.44	1	1214.9	160.7	15.9	-15.0	5.2	5 • ٤			0 . 5 .		. 6 . 7
	*** •	- 20.	70.44 70.444	. • • •	13.6	150.3	10.7	4.5 -18	1	6 • a 7 • ∪	0.442	4.0	\$ # # # - # # 25	.5.5	, . '
					1713.8 1713.5 1713.8	27.8		712.0	13.2	1.7	C	3 • •		7	1
					5	11'-3	15.0	-12.2	14.1	t • 1	0.7.	1.0	. 7.4	, B + 7	10.7
		* i = + +		20	1	8 . 2			14.1	F + 14				5 Ú a 14	• •
	•			-0.44					13.2		3.40		0.20	34.5	****
				-0.53	44.0	39.0		• •	×			7.4		3	25.46
- 1				-3.62	1721.02	74.7		J.6	6.		0.00	4		3.00	. 7 .
			~ . • • •	-v•":	1 21.47	81.1	4.5	. • b	• • •	2 • 4	50 4 <i>I</i>	••	. •	1.0	4.7 • 5
	1300+	ه څاري او او او	44	-0.79 -0.88	1722.6	78.2	1.6		1.5	2.0	0.00				440
		-	~	97	. 122.1						VIV.		0.00	J.,	. •
i	144	-2		6			4	. • 1			• 373		4.00	0.0	. •
,	152	*2.00		-1.15	42.4	• •		- • ·	****	• •			0.00	V • V	- 1 -
	L .	- 3000	6	15	1		- • ·					• .	• .		
:	52	-1 :-	-4.56	1.06	1 24.8	. 76. 7	3.0					• •			
4		- 1		0.91	22.4	194.6	• •	- :							• .
	<i>t</i> -					142.0	وود	- ·:		* 4	• •	• •	• •	- t · .	4
	64 ·	- 3 J 1 4 - 3 J 1	-1.66	. 8	1 4. 4	16	í			• •		• 1	0 € U € U € -1 €	:	~. • ^ ~ . • ^
		-33.0	6.0				2.5	-2.0		• 6			2.6		
		• • •		0.53	116600	177.6		+	• •		• • •	• '		. • •	. •
•	8 - J	 		. 45	1 21.1	1 2.6	• • •	1	. •	. • 2	U * v *	•	•		٠
1.	81.	-3	(6	6	2	160.1	5 . s e . s	- • 4		. • 5		. 5		6 • "	
	92. •	~ 3.00		•1.7	1 6.04	. > > . 4		-			•				
	ve .	-3		3.38			* • •	+		٠	• •				• •
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1.4	1500	- 1	-116 -116	55	1 71.1	111.	5	4							
		- 3 . •	-: : 6				4 . 1		4	2.	6	•	• •		
			+ 6 t + 6 t	5 3 6 .			200		•••	1.9		• • •	• • •	•	• * •
		-3000	-0.66	1	1124.5								•	•	•
23		- 3000	-0.06	-0.19	1 22.7	48.7	3.9	7.	- • M	. • 6					
. 4	14. 3.	-9:00	-1.66	-0.88	1255.48	16.	و و ا		. • *	• *	• •	• •		14.4	•
i i	1440.	- 1000 - 3	-0.66 -0.66	-0.01	1.2.08	73.4 81	3	•	. • •	• •	• -	• -	. •	- •	•
ή.	17.3	-30	66	-1.15	1 2 4 . 8	7	• •	:	•	•	:.			•	•
										-				-	

CIRCULAR STORM TEST

RUN BEGINS AT HOUR 1 ON JULY 4. 1776

STORY - BAROMETRIC PRESSURE AT CENTER OF LOW = 1000+0 MILLIBANS PRESSURE AT LANGEST ENCIRCLING 150BAR = 1020+0 MILLIBARS MAXIMUM PRESSURE INCLUDED IN STORM = 1022+9 MILLIBARS

LENGTH OF MAJOR HALF AXIS = 300.0 KILOMETERS
LENGTH OF MINOR HALF AXIS = 300.0 KILOMETERS
ORIENTATION JF MAJOR AXIS = 0.0 DEGREES FROM NORTH

STORM VELUCITY = 40. KILOMETERS/HOUR STORM AZIMUTH = 90. SHURE = PUSITION COURDINATES = X = 1000.0 Y = 0.0 KILOMETERS

SHORE LATITUDE = 43. ONSHORE AZIMUTH = 90. DEGREES NEARSHORE SLOPE = 0:033 AVERAGE FETCH = 1000. KILOMETERS

ноця	×	Y	A1	٧1	BARO. PRESS.	WIND ANGLE	SURF.	ONSH	AL SH WIND	EFFECT.	#AVE	WAVE	H R×	EAREK Anult	LSC VELUC.
	KM	KM	RAD	RAD	MB	DEG	M/5	M/5	M/5	M/5	M	SEC	M	DŁu	CM/SEC
1	483.	-400.	-0.89	1.15	1022.8	191.4	0.0	-0.0	-0.0	0.0	0.00	0.0	0.00	->.4	-0.0
2	520.	00.	-0.89	1.06	1022.8	189.1	0.0	-0.0	-0.0	0.0	0.00	0.0	0.00	-4.4	-0.0
3	560.	-400.	-0.89	0.97	1022.8	186.7	0.0	-0.0	-0.0	0.0	0.00	0.0	0.00	-3.2	-0.0
•	600.	-400.	-0.89	0.88	1022.8	163.9	0.0	-0.0	-0.0	0.0	0.00	0.0	0.00	-1.9	-0.0
5	640.	-400.	-0.89	0.80	1022.8	180.9	0.1	-0.1	-0.0	0.0	0.00	0.0	0.00	-0.4	-0.0
6	680.	-400.	-0.89	0.71	1022.8	177.6	0.3	-0.3	0.0	0.1	0.00	0.0	0.00	1.1	0.0
7	720.	-+00.	-0.89	0.62	1022.7	173.9	0.5	-0.4	0.0	0.1	0.00	0.1	U . 00	3.0	0
8	760.	-400.	-0.89	0.53	1022.7	169.9	0.7	-0.7	0.1	0.2	0.00	0 • 1	0.00	5.4	U /
10	840.	-400.	-0.89	0.44	1022.6	165.5	1.0	-1.0	0.2	0.3	0.00	0.4	0.00	7.5	1.4
-1	880.	-400.	-0.89	0.35	1022.5	160.8	1.4	-1.3	0.4	0.5	3.00	0.3	U • 0C	9.9	2.4
	920	00	-0.89	0.26	1022.4	155.7	1.7	-1.6	0.7	0.6	0.00	0.3	0.00	12.5	3.7
13	960.	-400	-0.89	0.08	1022.2	144.7	2.2	-1.8	1.0	0.7	0.01	0.4	0.01	15.1	5.5
14	1000.	-400.	-0.89	0.00	1022.2	139.0	2.3	-1.8	1.3	0.8	0.02	0.5	0.01	2014	0.7
15	1040.	-400.	-0.89	-0.08	1022.2	133.3	2.2	-1.5	1.0	0.9	0.02	0.0	0.02	22.4	8 . 4 7 . :
. 6	1080.	-400.	-0.89	-0.17	1022.3	127.7	2.0	-1.2	1.6	0.9	0.02	0.6	0.02	24.1	10.0
17	1120.	-400.	-0.89	-0.26	1022.4	122.3	1.7	-0.9	1.5	0.8	0.02	J.6	0.01	2402	9.4
18	1160.	-400.	-0.89	-0.35	1022.5	117.3	1.4	-0.6	1.2	0.7	0.01	0.5	0.01	45.5	8.0
19	1200.	-400.	-0.89	-0.44	1022.6	112.5	1.0	-0.4	1.0	0.5	0.01	0.4	0.00	26.5	0.3
23	1240.	-400.	-0.89	-0.53	1022.7	108.1	0.7	-0.2	0.7	0.4	0.00	0.3	0.00	47.5	4.0
21	1280.	-400.	-0.89	-0.62	1022.7	104.1	0.5	-0.1	0.5	0.3	0.0	0.2	0.00	27.9	3
22	1320.	-400.	-0.89	-0.71	1022.8	100.4	0.3	-0.0	0.3	0 • 2	0.00	0.1	0.00	28. 1	
23	1360.	-400.	-0.89	-0.79	1022.8	97.1	0.2	-0.0	0.1	0.1	0.00	0.0	0.00	∠8.6	1.0
2.	1400.	-400.	-0.89	-0.88	1022.8	94.1	0.1	-0.0	0.1	0.0	0.00	0.0	U • 00	31.6	٠,4
1	1440.	-400.	-0.89	-0.97	1022.8	91.4	0.0	-0.0	0.0	0.0	0.60	0.0	0.00	0.0	0.0
2	1480.	-400.	-0.89	-1.06	1022.8	88.9	0.0	0.0	0.0	0.0	0.00	0.0	0.00	0.0	0.0
3	1520.	-400.	-0.89	-1-15	1022.8	86.6	0.0	0.0	0.0	0.0	0.00	0.0	0.00	0.0	٥.٠
1	483.	-500 ·	-1.11	1.15	1022.8	185.1	0.0	-0.0	-0.0	0.0	0.00		4.00	-2.4	-0.0
2	520.	-500.	-1.11	1.00	1322.8	192.8	0.0	-0.0	-0.0	نون	0.00	0.00	0.00	-1.0	-0.0
3	560.	-500.	-1.11	⊍ 497	1022.8	180.3	0.0	-0.0	-0.0	0.0	0.00	0.0	U . OU	-0.1	-0.0
4	600.	-500.	-1.11	0.88	1022.8	177.6	0.0	-0.0	0.0	0.0	0.00	U • U	□ • U □	1	Cau
5	640.	-500.	-1.11	J. 80	1022.8	174.7	0.0	-0.0	0.0	0.0	3.00		C 4 (1)	4.5	(• C
6	600.	-500.	-1.11	0.71	1022.8	171.6	0.0	-0.0	0.0	0.0	2.00	0.0	4.00	4.6	J
7	720.	-500.	-1-11	0.62	1022.8	168.2	0.0	-0.0	0.0	0.0	U . U U	ن و ن	0.00	5.0	∂. ∪
8	760.	-500.	-1-11	0.53	1022.0	164.6	0.1	-0.1	0.0	() • U	J. UU	0.0	4.00	I • •	0.1
10	840.	-500	-1 - 11	0.44	1022.8	160.8	3.1	-0.1	0.0	0.0	0.00	0.0	C+00	9.3	و ۽ ن
11	880.	-500.	-1-11	0.35	1022.8	156.7	0.2	-0.2	0.0	0.0	0.00	ن د ل	0.00	11.4	0.5
12	920.	-500.	-1.11	0.26	1022.8	152.5	0.3	-0.2	0.1	0.1	0.00	0.0	0.00	13.4	J .
13	960.	-500	-1.11	0.08	1022.0	143.6	0.5	-0.2		0 • 1		3.0	0.00	13.5	
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15	1040.	-500	-1.11	-0.08	1022.8	134.4	0.3	-0.2	0.2	0.1	0.00	0.1	0.00	20.7	4.5
16	1080.	-500	-1.11	-0.17	1022.8	129.9	C . 3	-0.2	0.4	0.1	0.00	0 - 1	0.00	11.6	4.01
17	1120.	-52U.	-1.11	-0.26	1022.8	125.5	0.2	-0.1	0.2	0.1	0.00	0.1	0.00		4 • •
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A mathematical simulation model of a conto forecast or hindcast wave and longshore	
site. Storm parameters for the model are be	ased on the size, shape intensi
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used to model the size and shape of the storing the length and orientation of the major	
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20. Abstract continued

inverted normal curve through the storm center. The storm path is based on actual storm positions for the hindcast mode, and on projected positions assuming constant speed and direction for the forecast mode. The location, shoreline orientation and nearshore bottom slope provide input data for each coastal site.

For each storm position, the geostrophic wind speed and direction are computed at the shore site as a function of barometric pressure gradient and latitude. The geostrophic wind is converted into surface wind speed and direction by applying corrections for frictional effects over land and sea. The surface wind speed, fetch and duration are used to compute the wave period, breaker height and breaker angle at the shore site. The longshore current velocity is computed as a function of wave period, breaker height and angle and nearshore bottom slope.

The model was tested by comparing hindcast output with observed data for several coastal locations. Forecasts were made for actual storms and for hypothetical circular and elliptical shaped storms.